

Probing the interactions between resolved and parameterized waves

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Questions

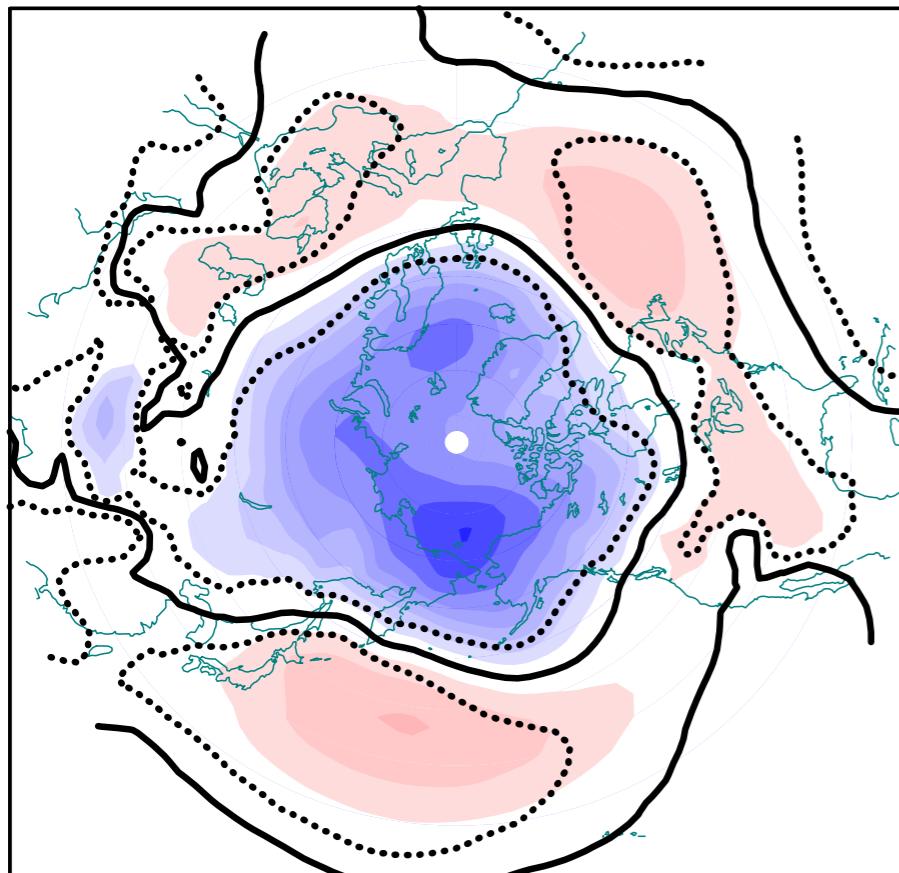
- Can we quantify the uncertainty in climate change integrations associated with parameterized gravity waves?

Questions

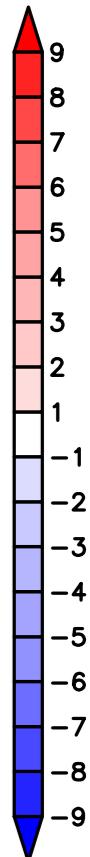
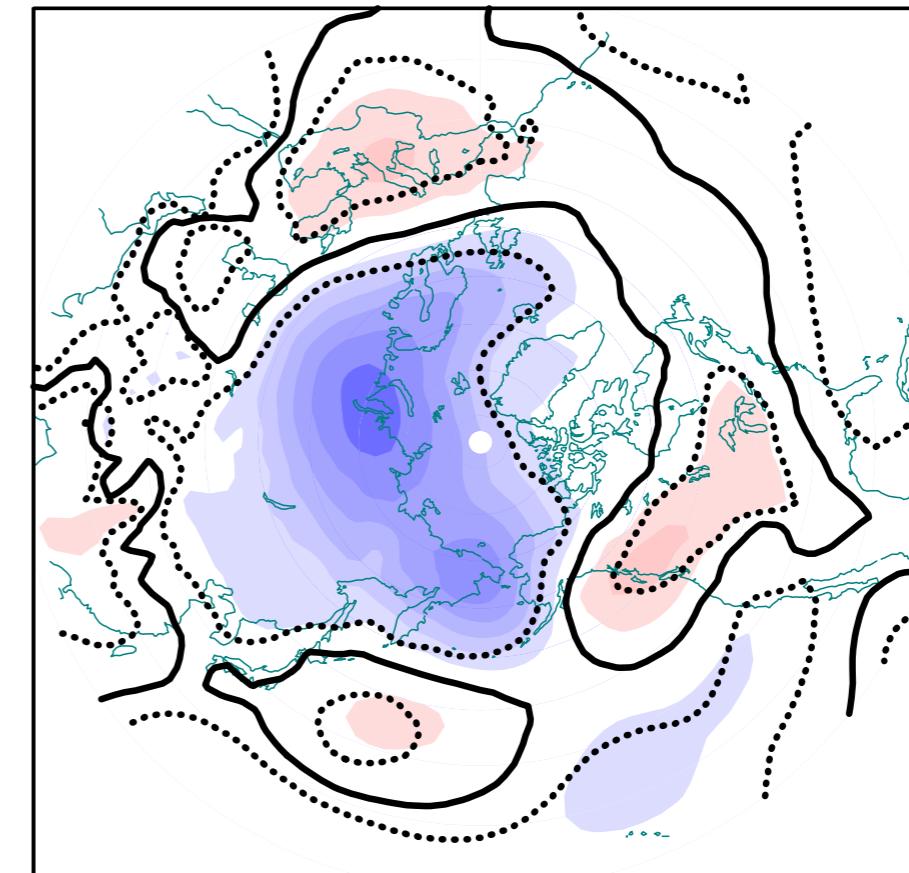
- Can we quantify the uncertainty in climate change integrations associated with parameterized gravity waves?

SLP response to doubled CO₂, CMAM

(c) RESPONSE LOWERED



(d) RESPONSE LOW-G



[Sigmond et al. 2008, Sigmond and Scinocca, 2010]

Questions

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 - ➔ develop an idealized GCM framework to investigate interactions between resolved and parameterized waves

Questions

- Can we quantify the uncertainty in climate change integrations associated with parameterized gravity waves?
 - ➔ develop an idealized GCM framework to investigate interactions between resolved and parameterized waves
- How predictable is the response to perturbations of the parameterized gravity waves?
 - ➔ experiments designed to maximize interaction with resolved waves

An idealized Atmospheric GCM

- dry primitive equations on the sphere
- Newtonian relaxation of temperature to radiative-convective equilibrium profile

- Variable strength polar vortex
[Polvani and Kushner, 2002]

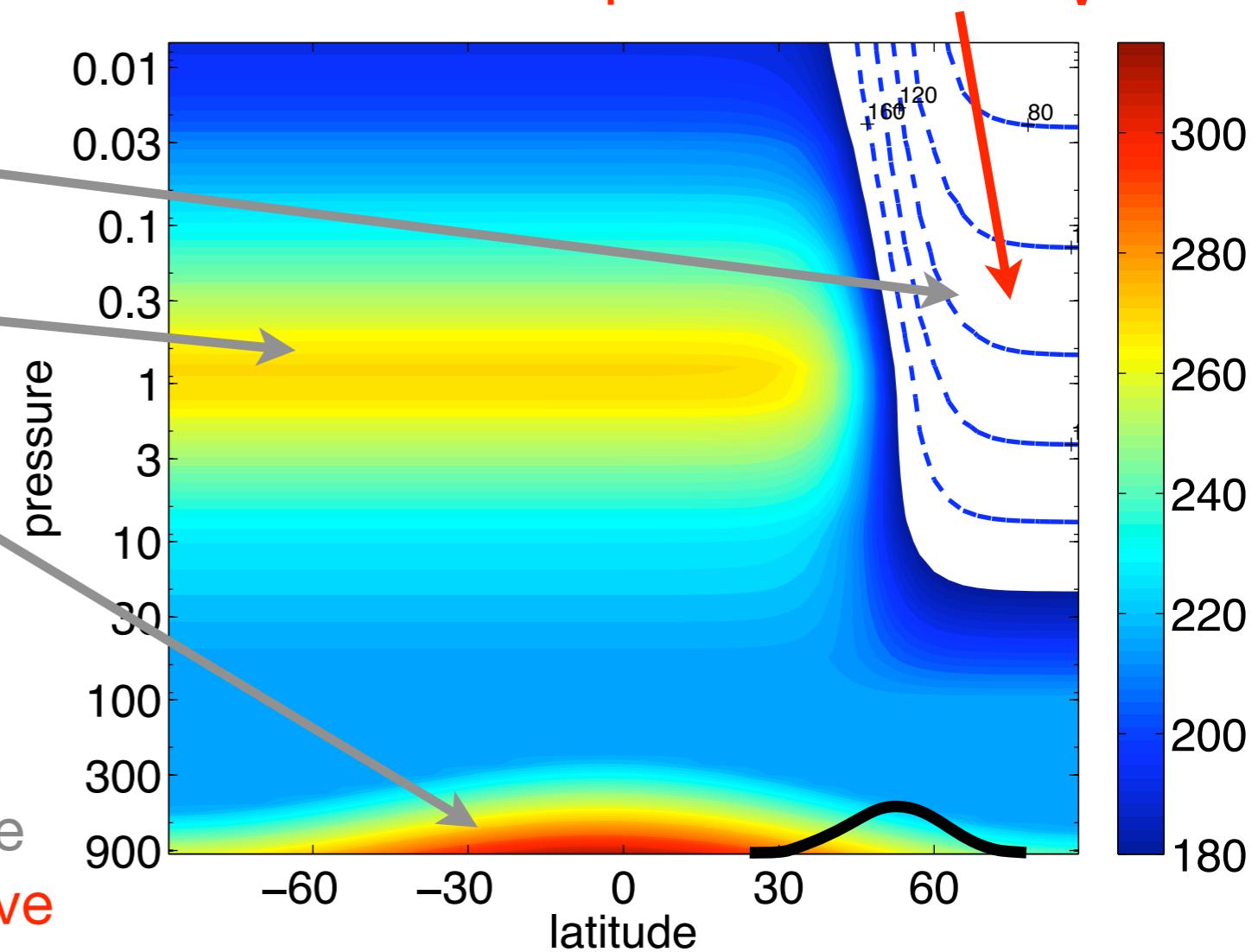
- US Standard Atmosphere

- *Held and Suarez [1994]*
troposphere

- Simple large scale topography
[Gerber and Polvani, 2009]

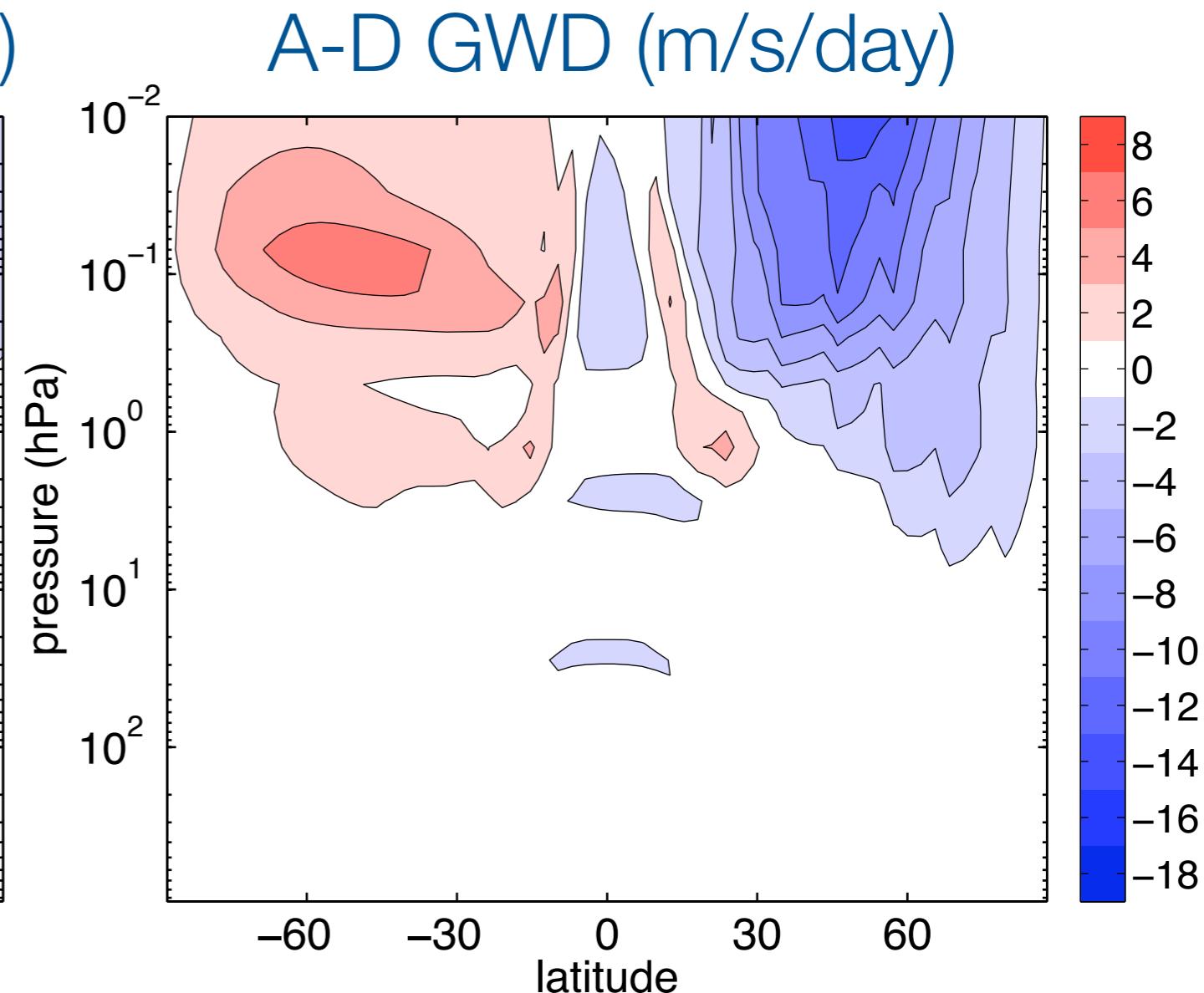
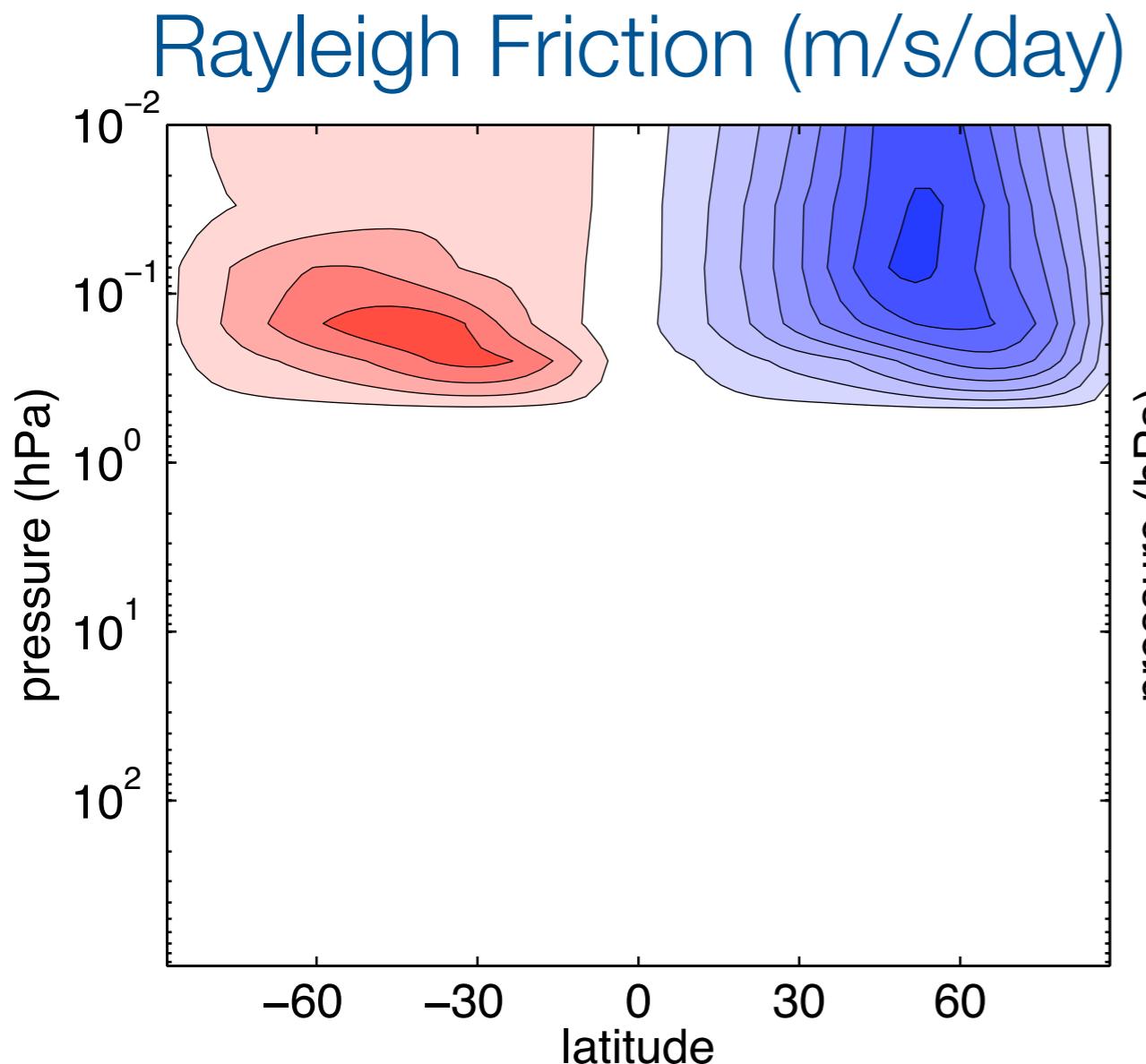
- **Rayleigh friction** at bottom (surface drag) and **at top (crude gravity wave parameterization)**

vary cooling w/
parameter γ



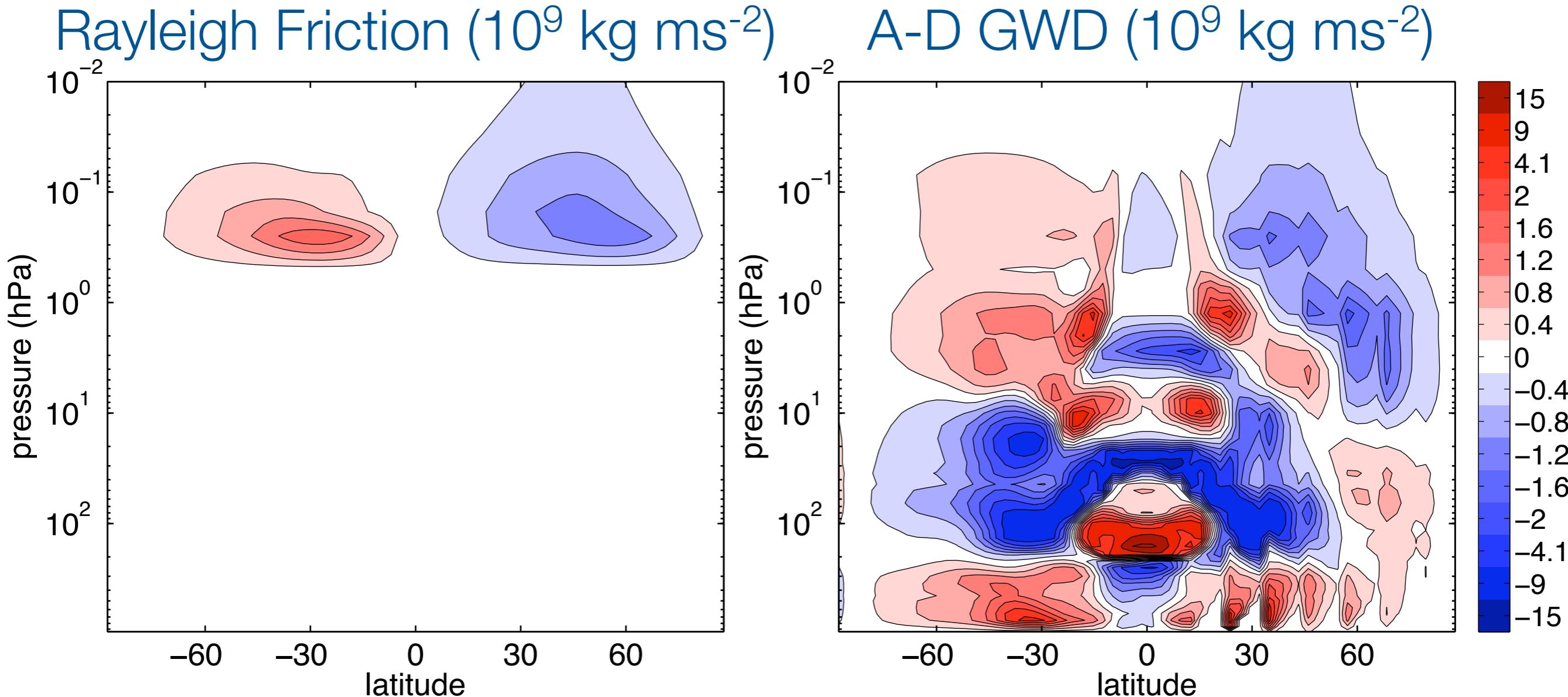
Replace Rayleigh Friction with Alexander-Dunkerton (1999) GWP

- Spectral source, approximates effect of intermittency on wave breaking
- Used to parameterized non-orographic GWD in GFDL coupled climate model



Replace Rayleigh Friction with Alexander-Dunkerton (1999) GWP

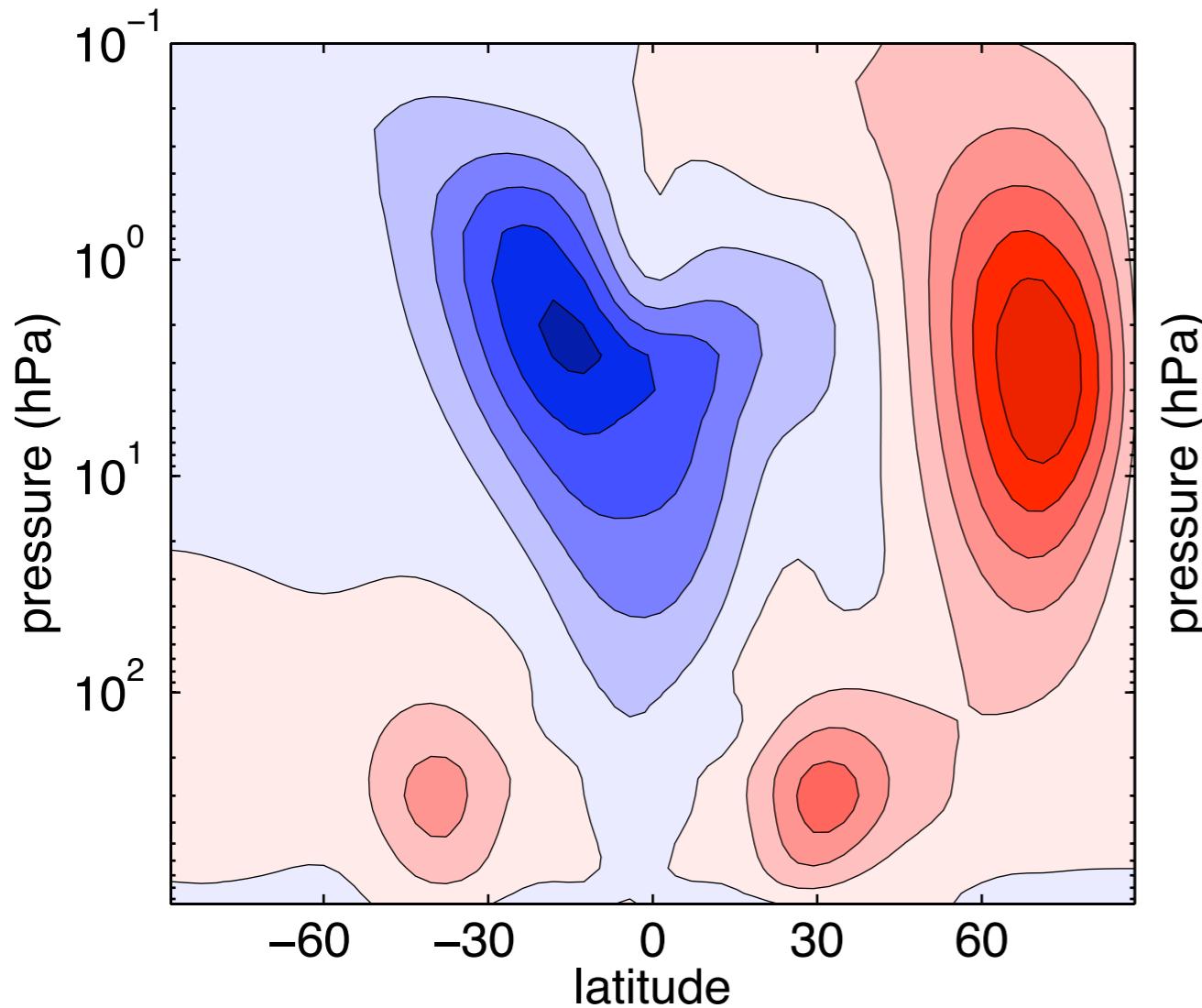
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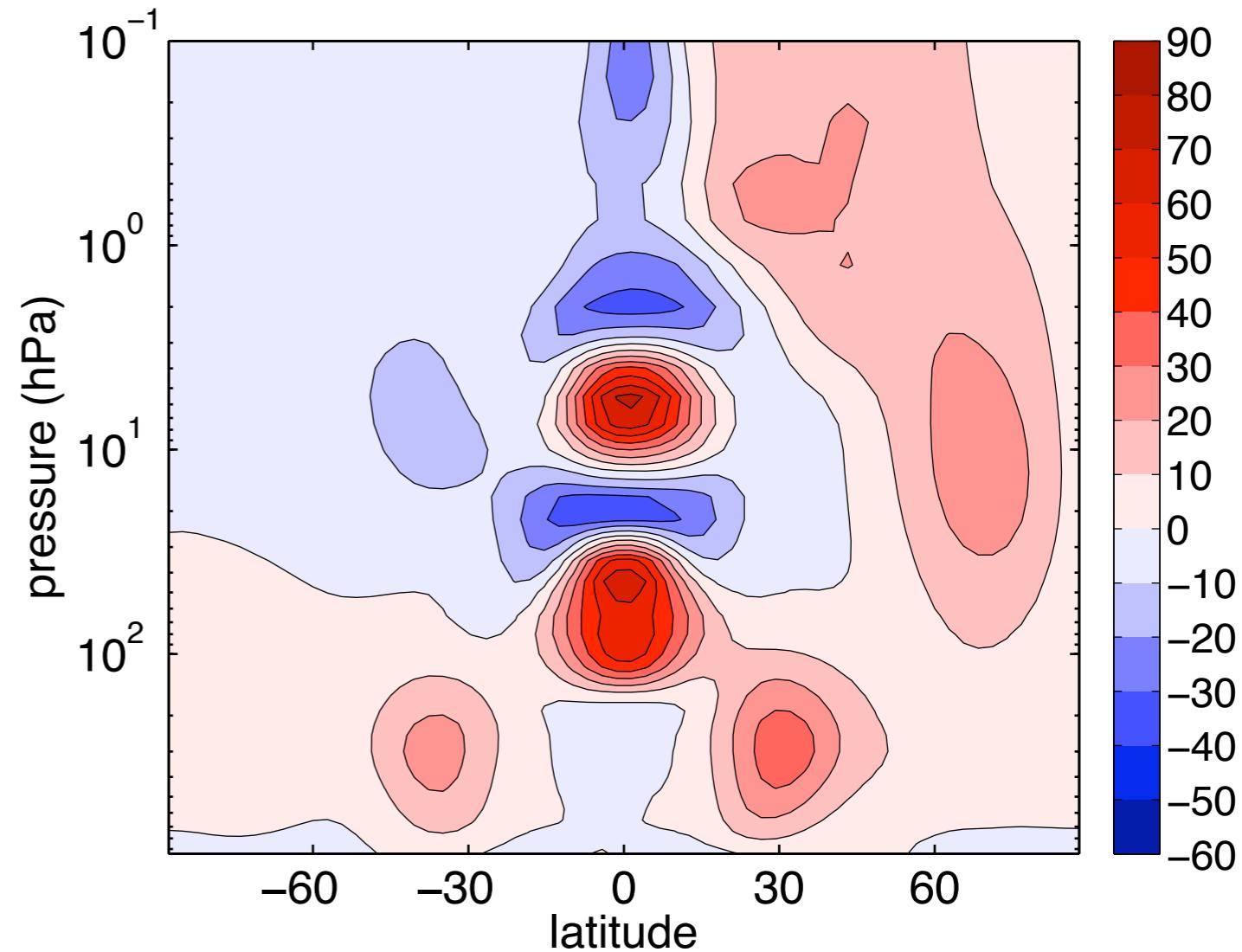
Retuning the model ...

- For same equilibrium temperature forcing, vortex is too weak in new model

\bar{u} , $\gamma=4$ (Rayleigh Friction)



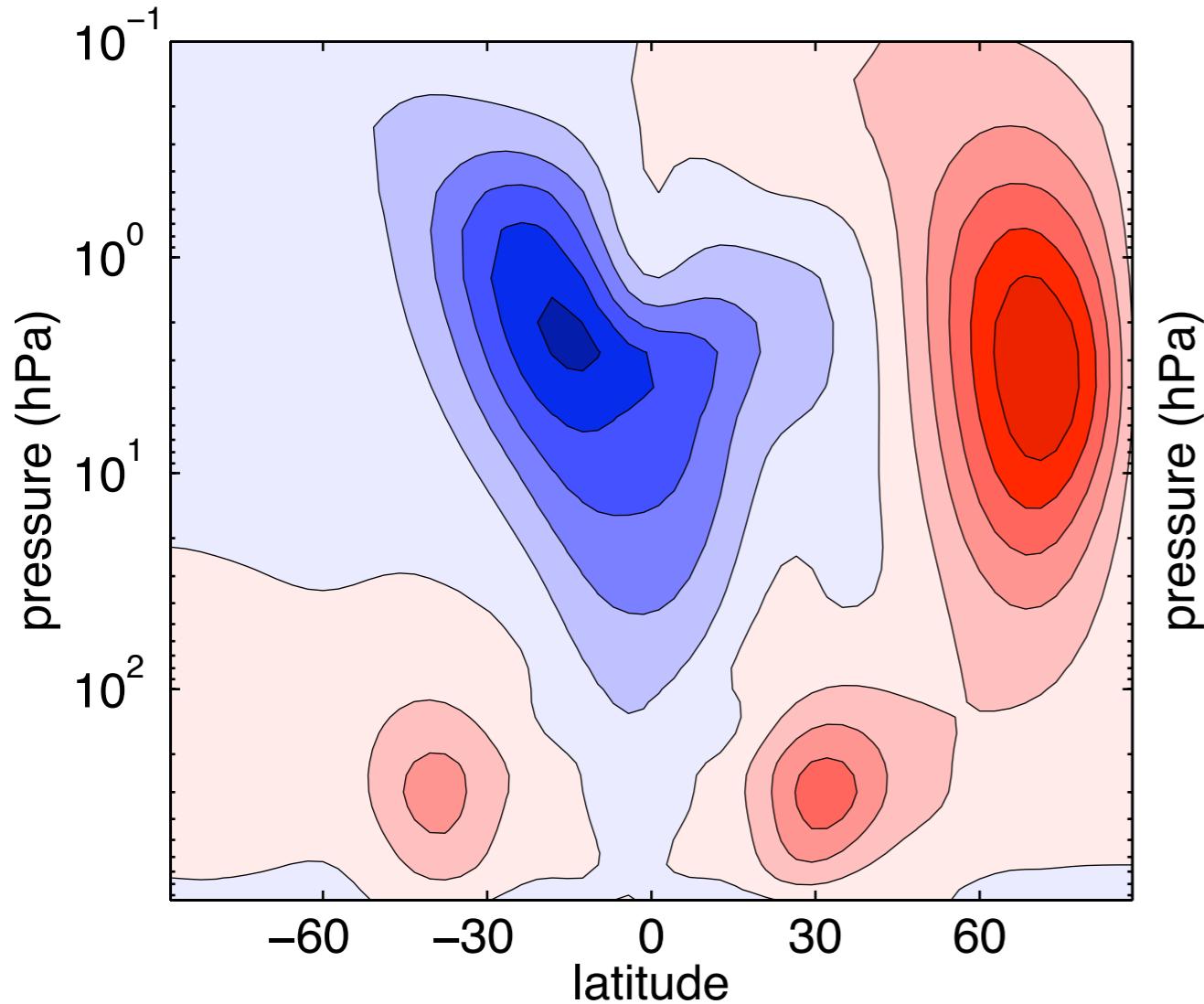
\bar{u} , $\gamma=4$ (A-D GWP)



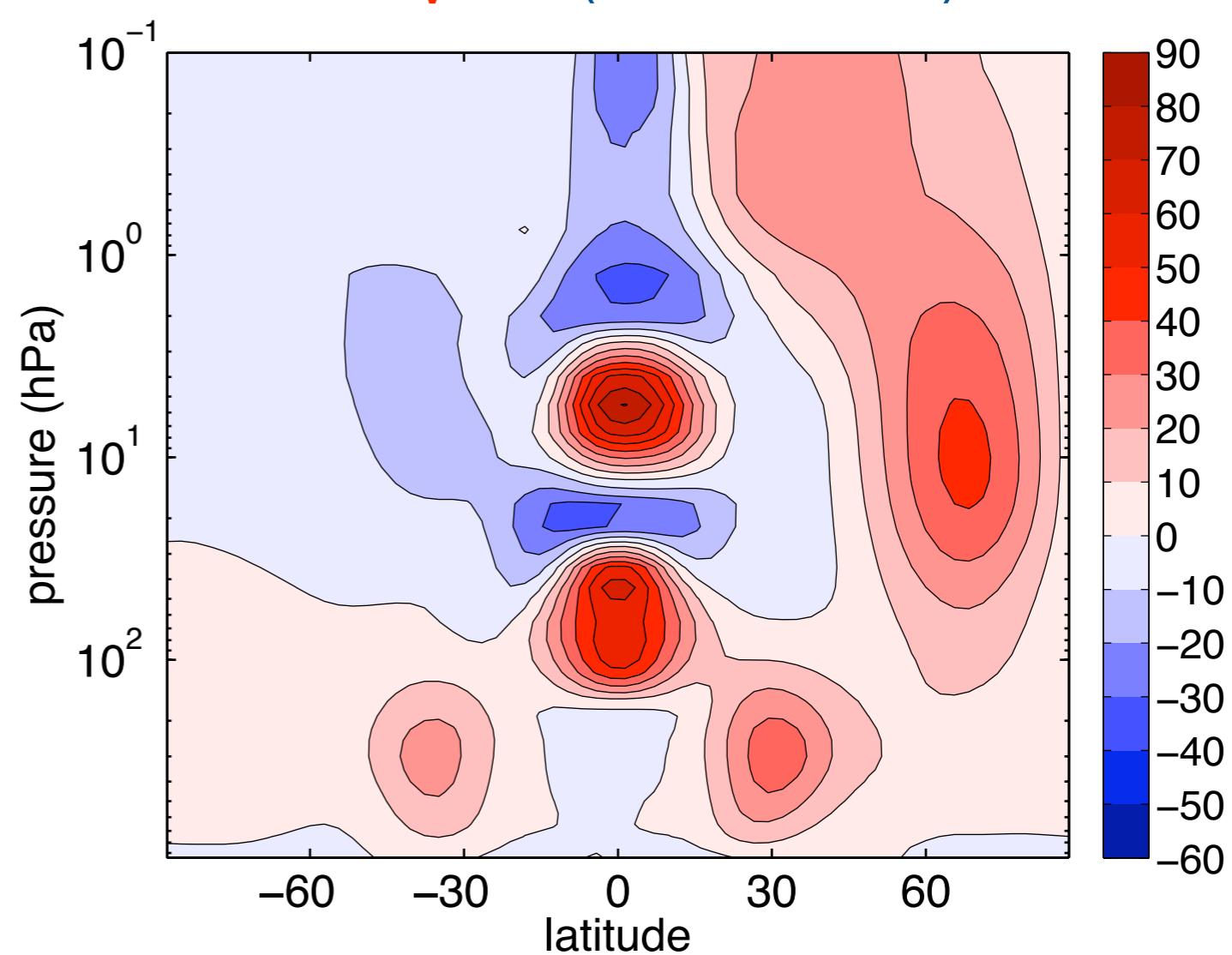
Retuning the model ...

- For same equilibrium temperature forcing, vortex is too weak in new model
- Adjust radiative profile to get stratospheric jets correct

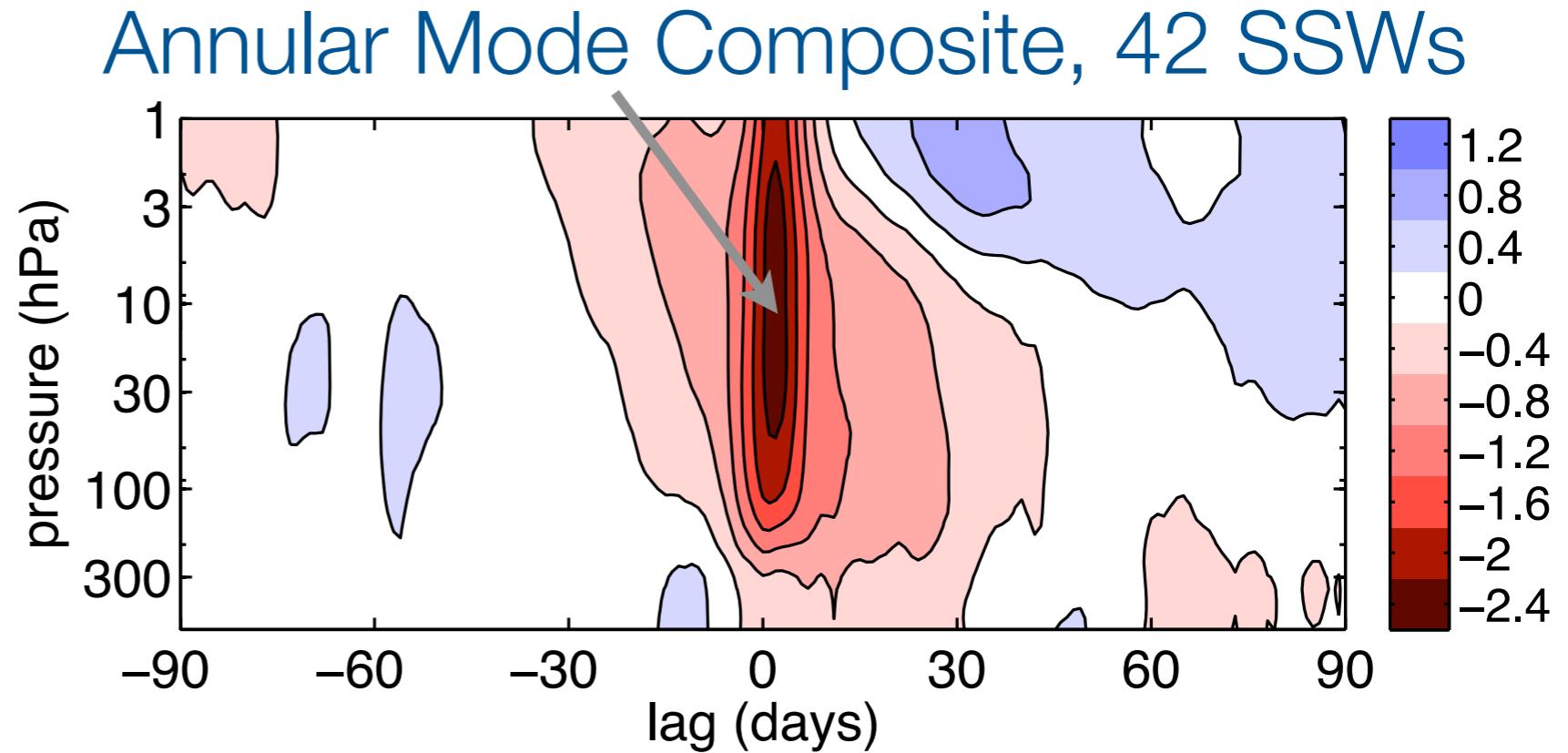
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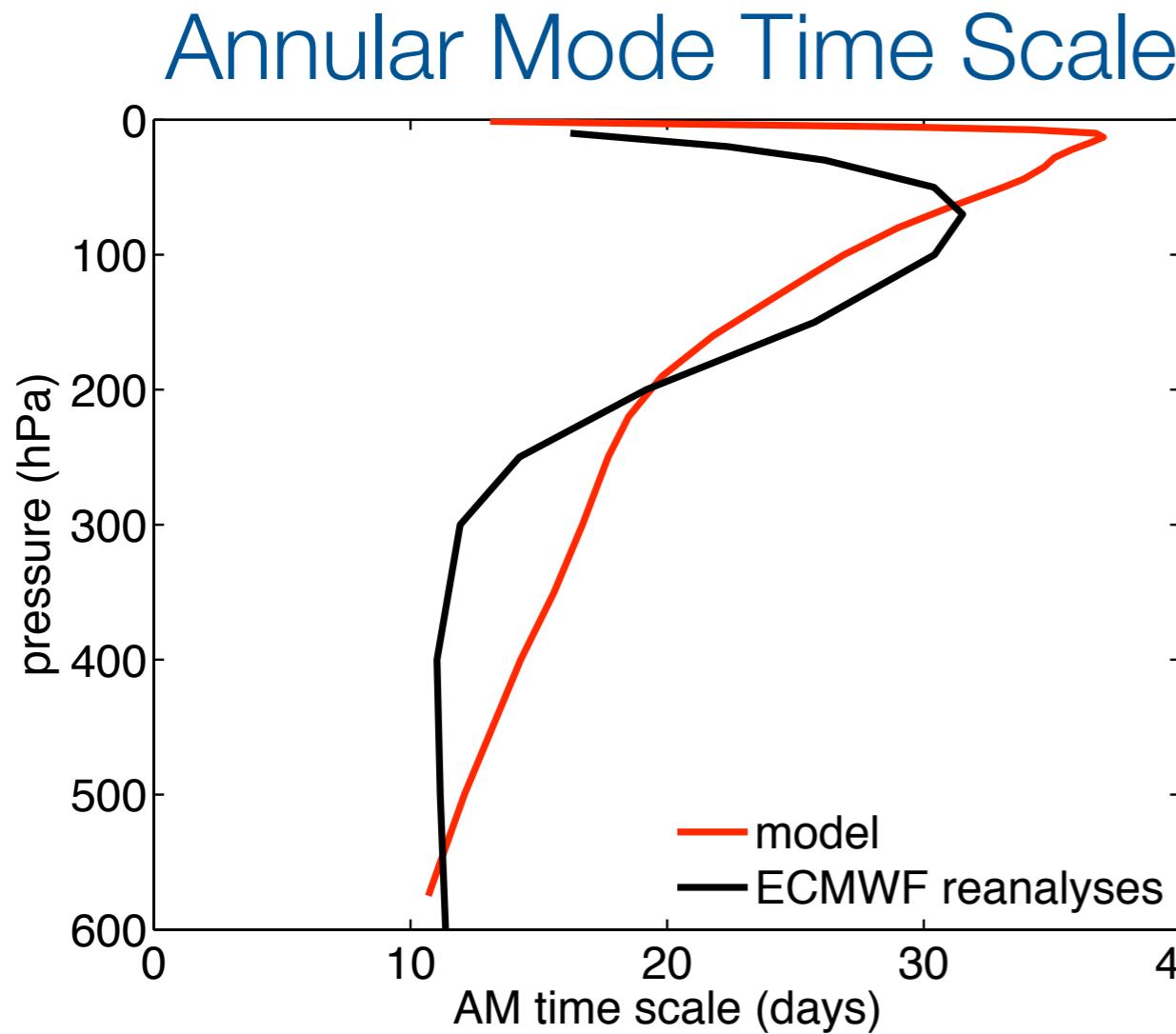
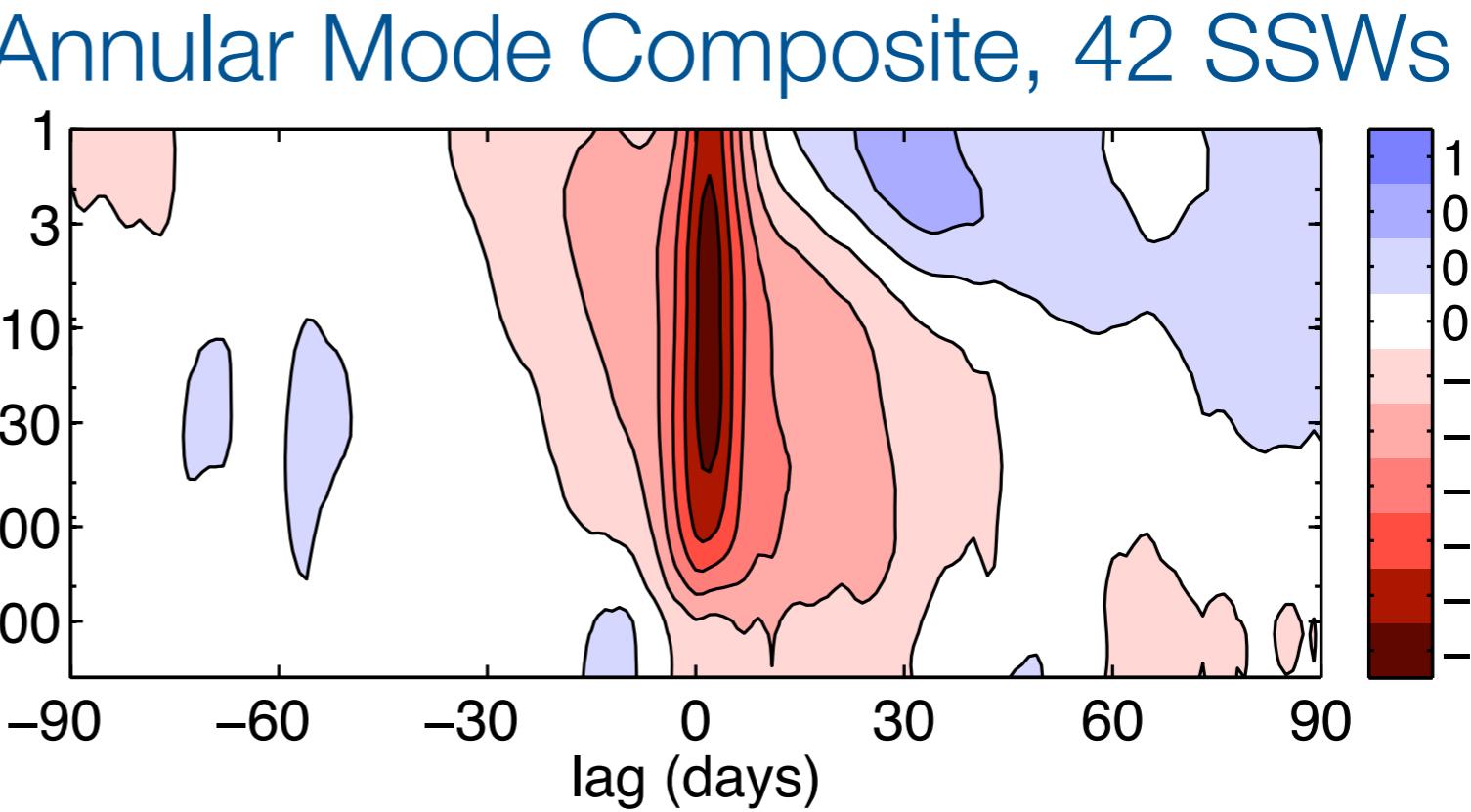
\bar{u} , $\gamma=6$ (A-D GWP)



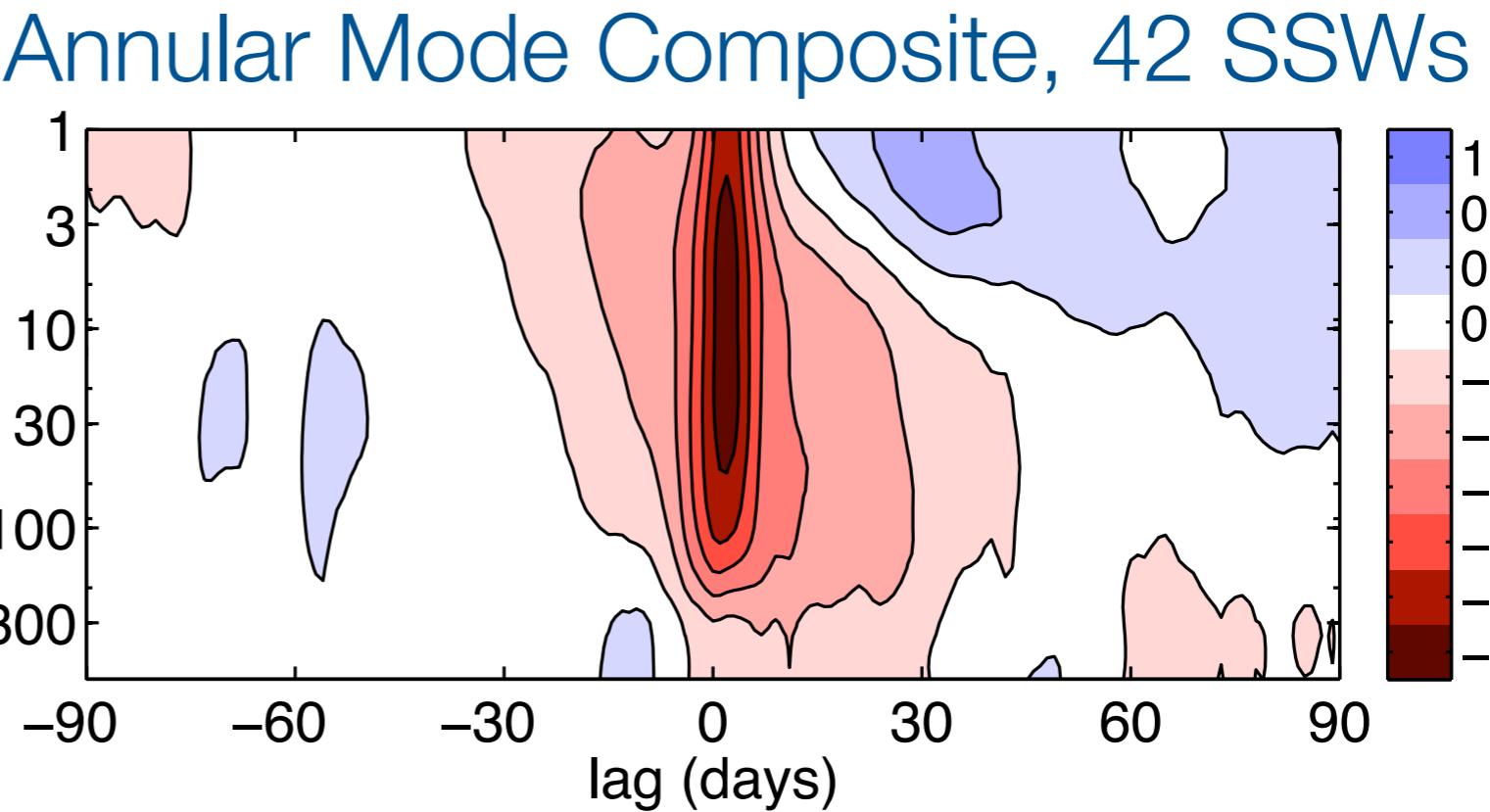
Model captures
Stratosphere-
Troposphere
Coupling



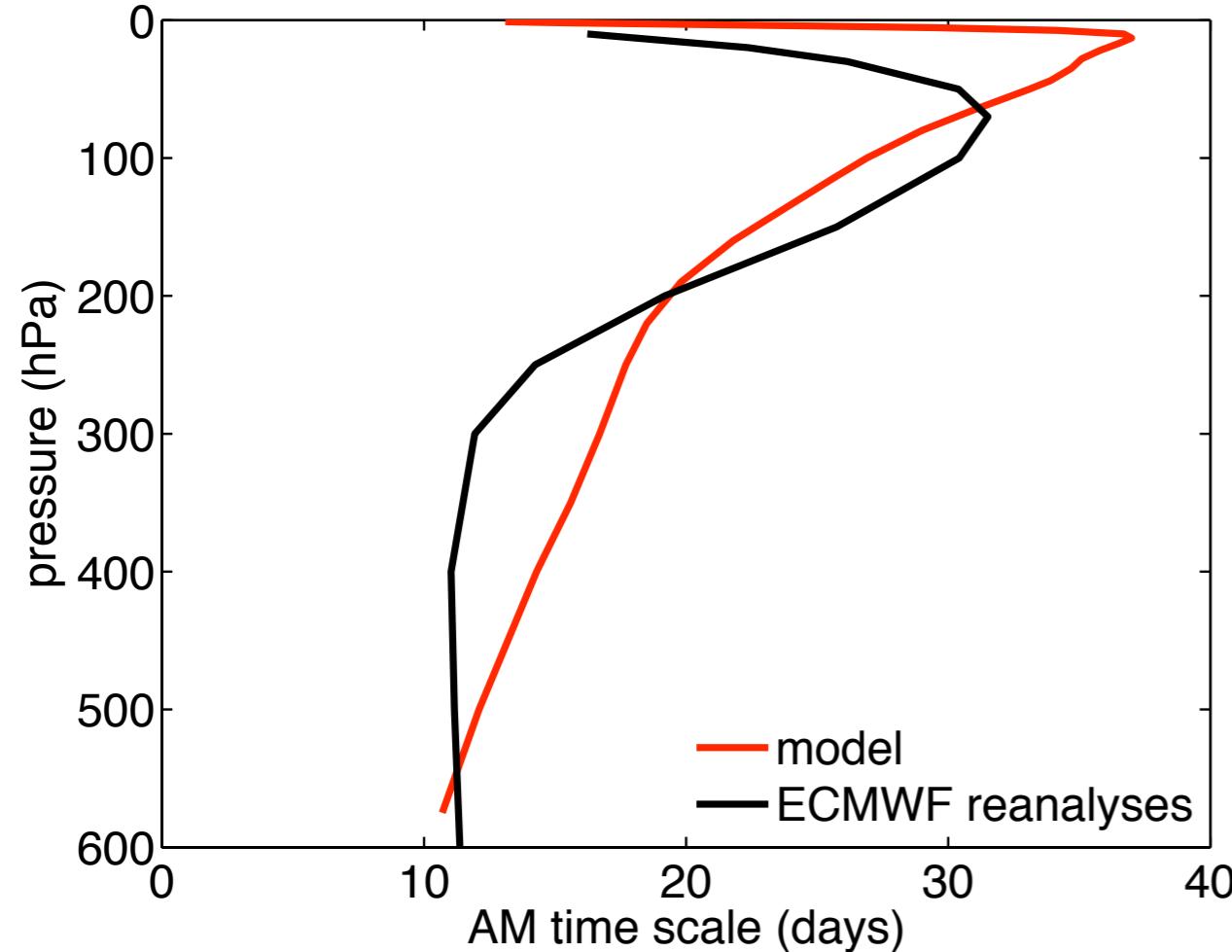
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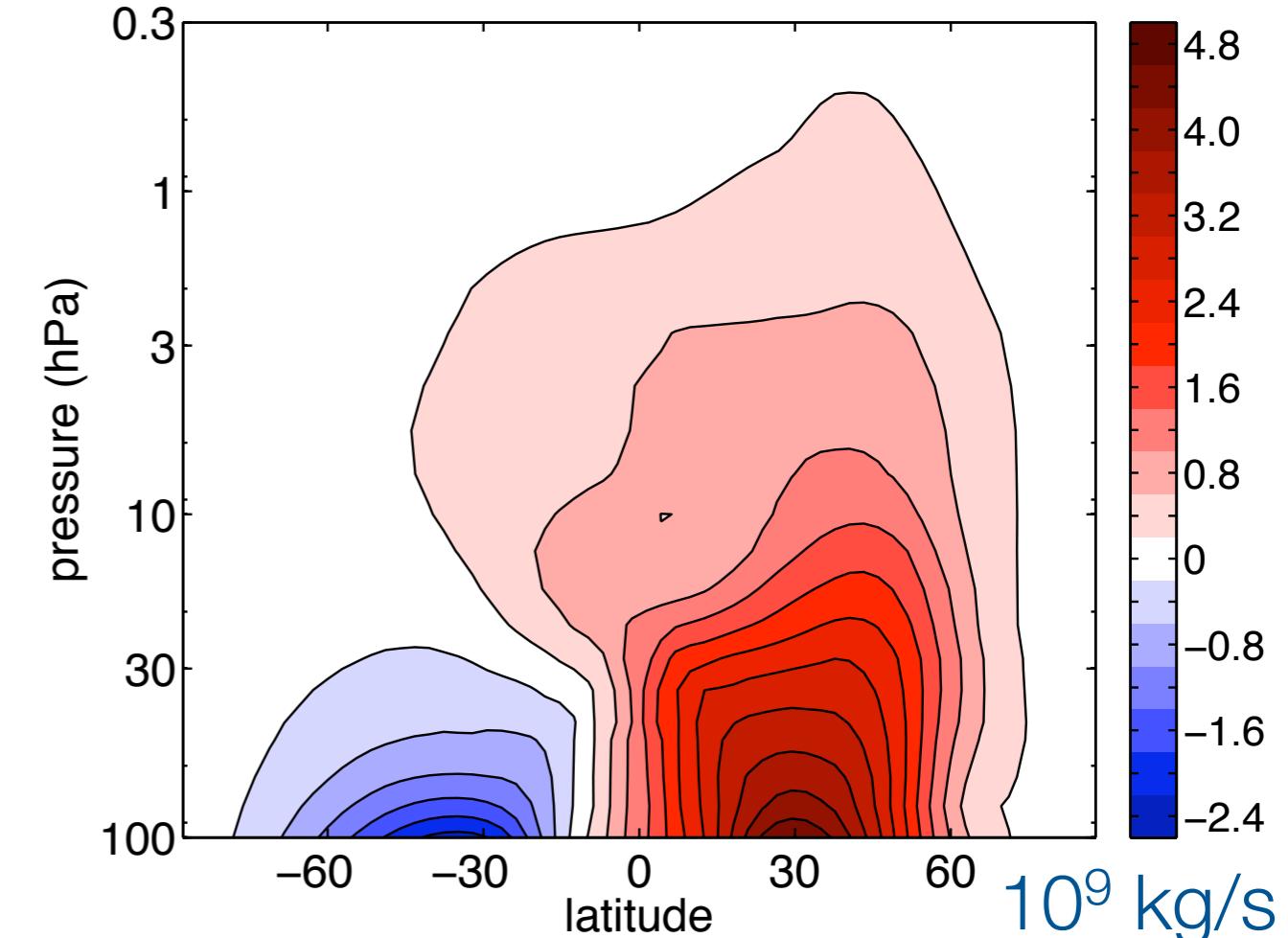
Model captures
Stratosphere-
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Annular Mode Time Scale



Residual Mean Circulation



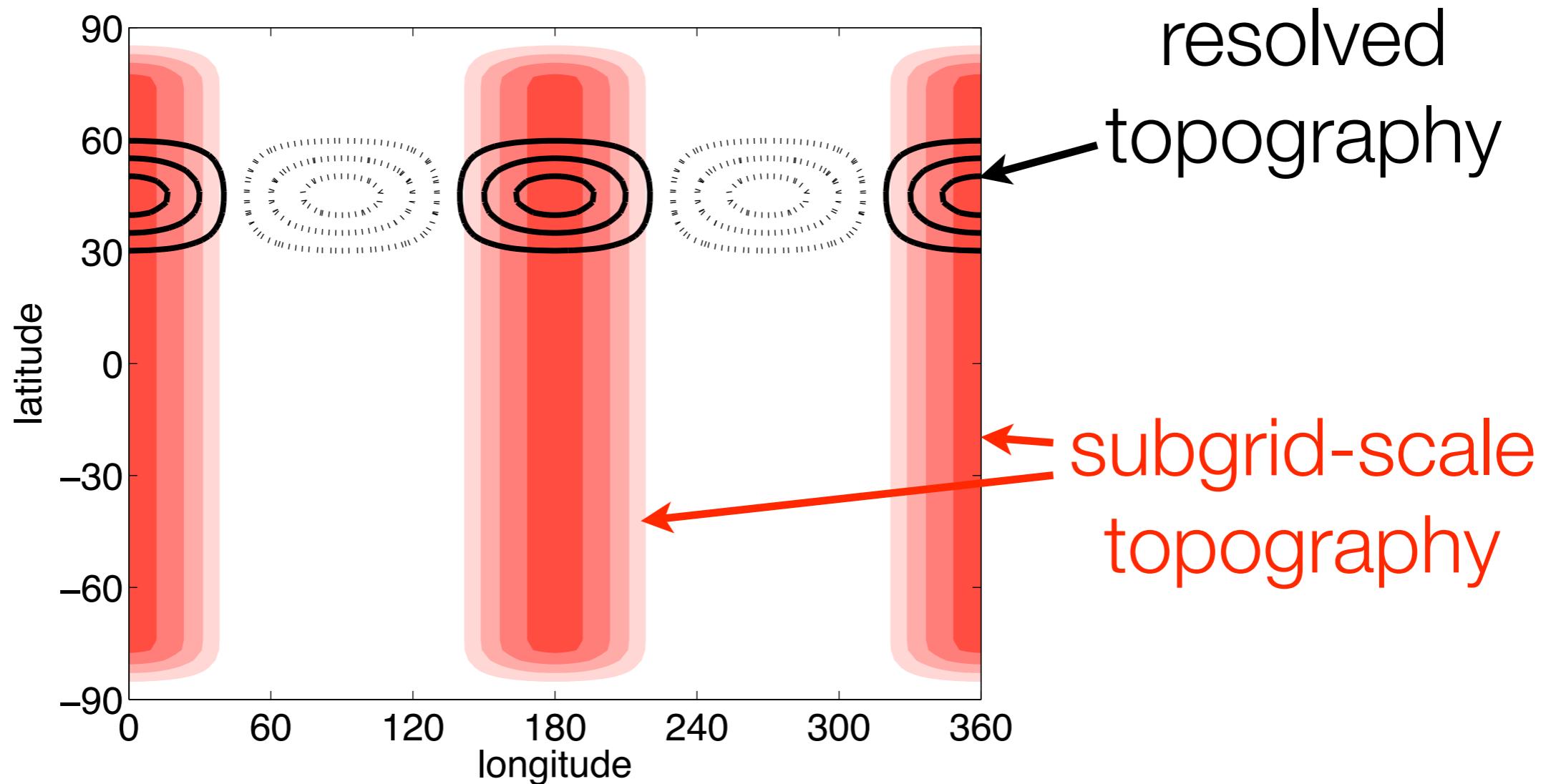
Orographic Gravity Waves: Pierrehumbert Scheme

- Used in GFDL's coupled climate models CM2 and CM3.
- Lindzen-like scheme
 - Estimate wave flux from low level flow, subgrid-scale topography parameter
 - Wave breaking determined by comparing flux against a saturation flux. Momentum is deposited slowly, as saturation flux decays with height.
- Tends to allow small amount of flux reach near model lid; trivial amount of momentum, but leads to numerical issues.

Experimental Setup:

Maximize interaction with resolved planetary waves

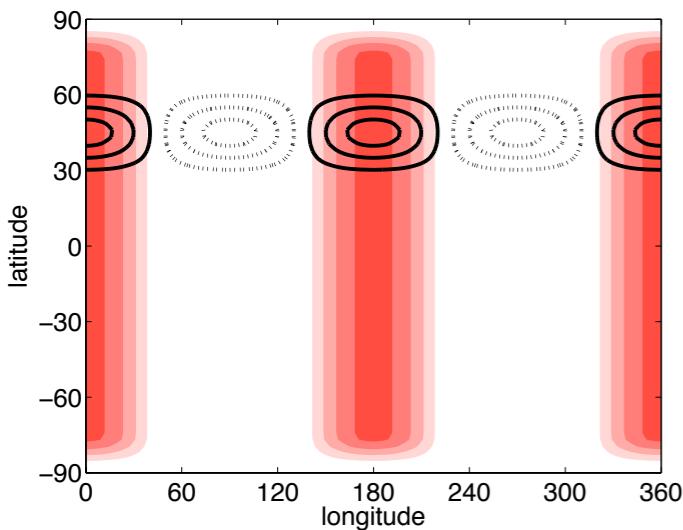
- Wavenumber 2 topography creates a standing waves in stratosphere
- Subgrid-scale topography with large scale wave number 2 pattern



Experimental Setup:

Maximize interaction with resolved planetary waves

- Wavenumber 2 topography creates a standing waves in stratosphere
- Subgrid-scale topography with large scale wave number 2 pattern
- Vary phase between the subgrid-scale pattern and the actual topography: differences in surface winds and upper level structure will perturb orographic wave flux and breaking criteria

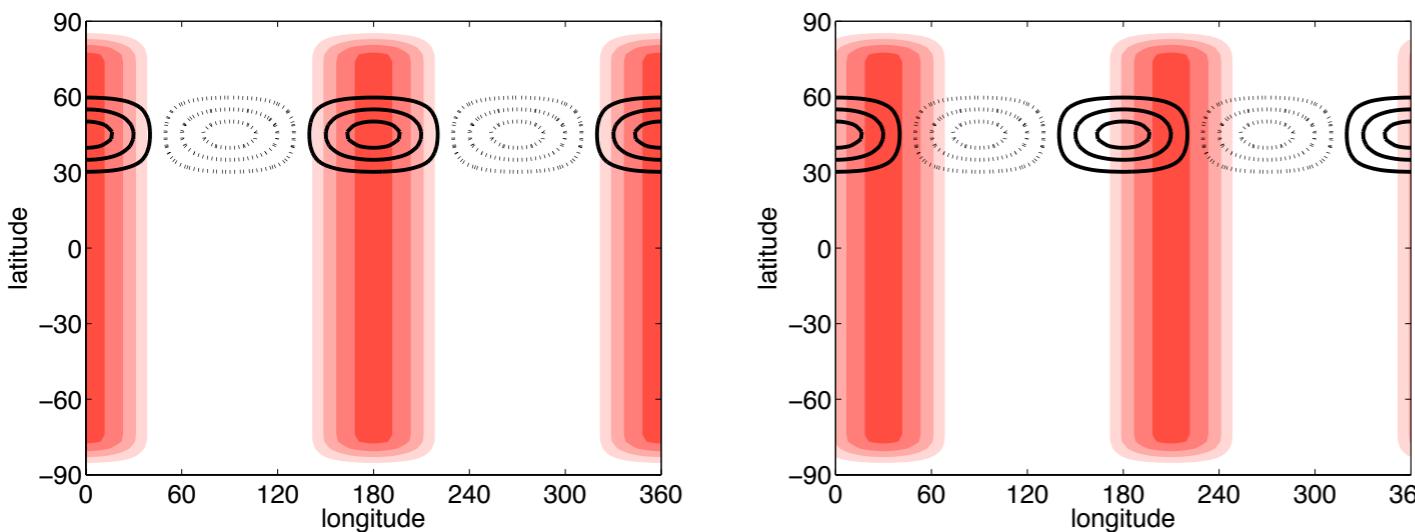


phase shift = 0°

Experimental Setup:

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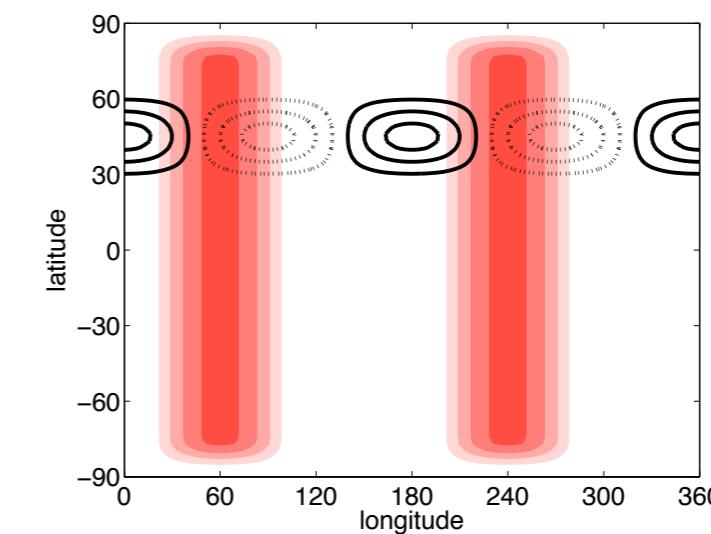
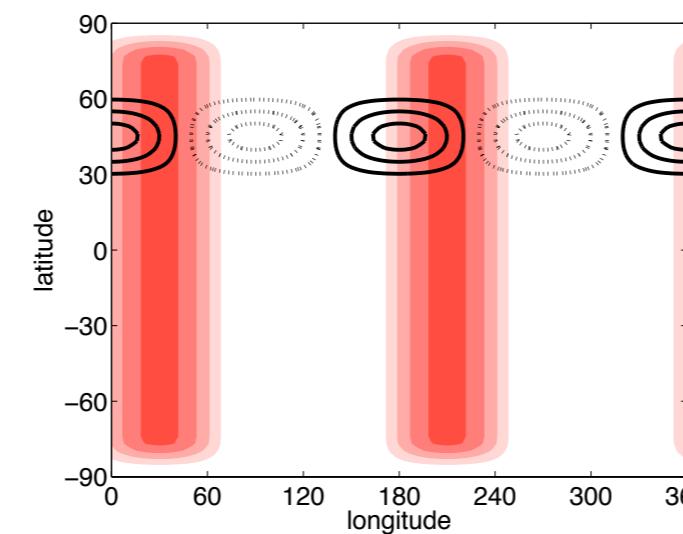
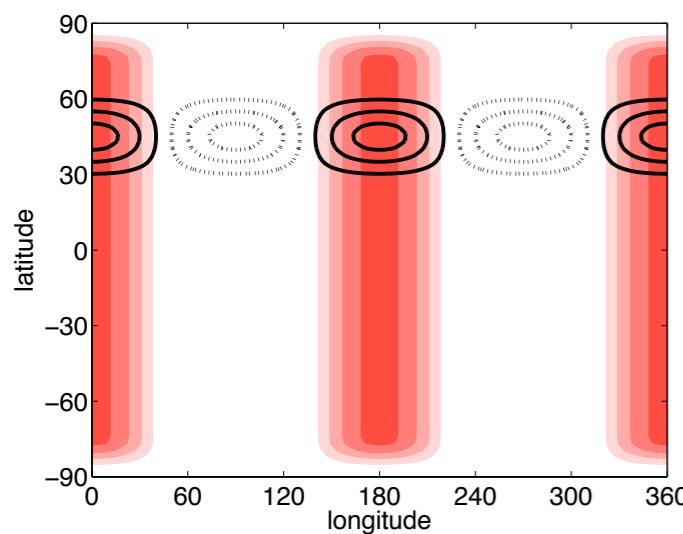


phase shift = 0° phase shift = 30°

Experimental Setup:

Maximize interaction with resolved planetary waves

- Wavenumber 2 topography creates a standing waves in stratosphere
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phase shift = 0°

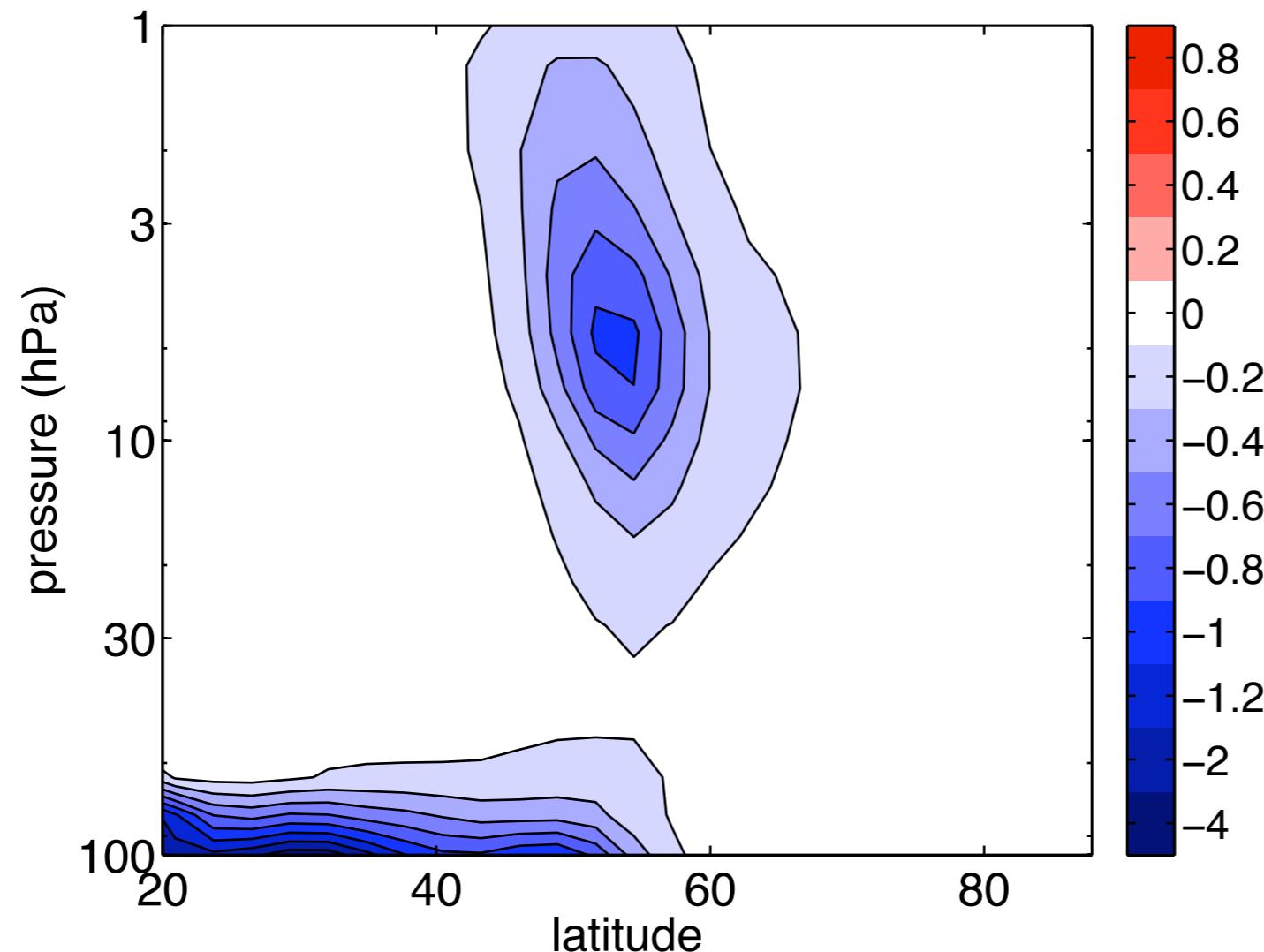
phase shift = 30°

phase shift= 60°

...

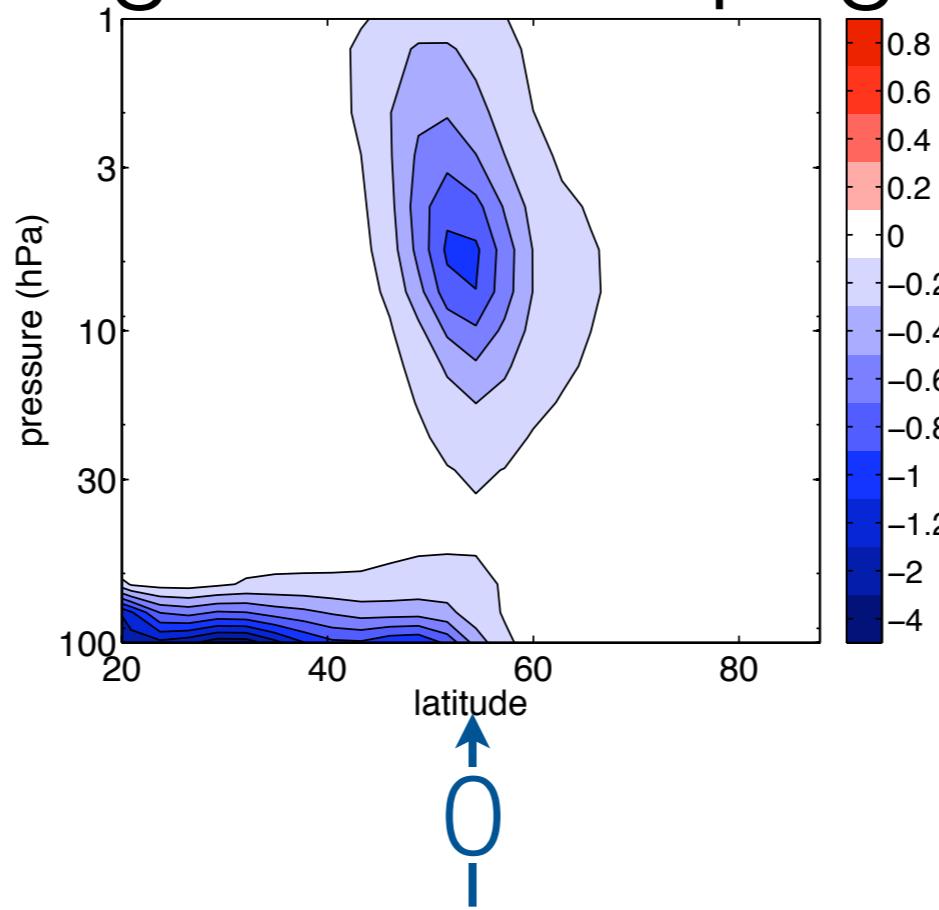
Impact of orographic gravity wave parameterization

OGW Forcing ($10^{10} \text{ kg ms}^{-2}$)



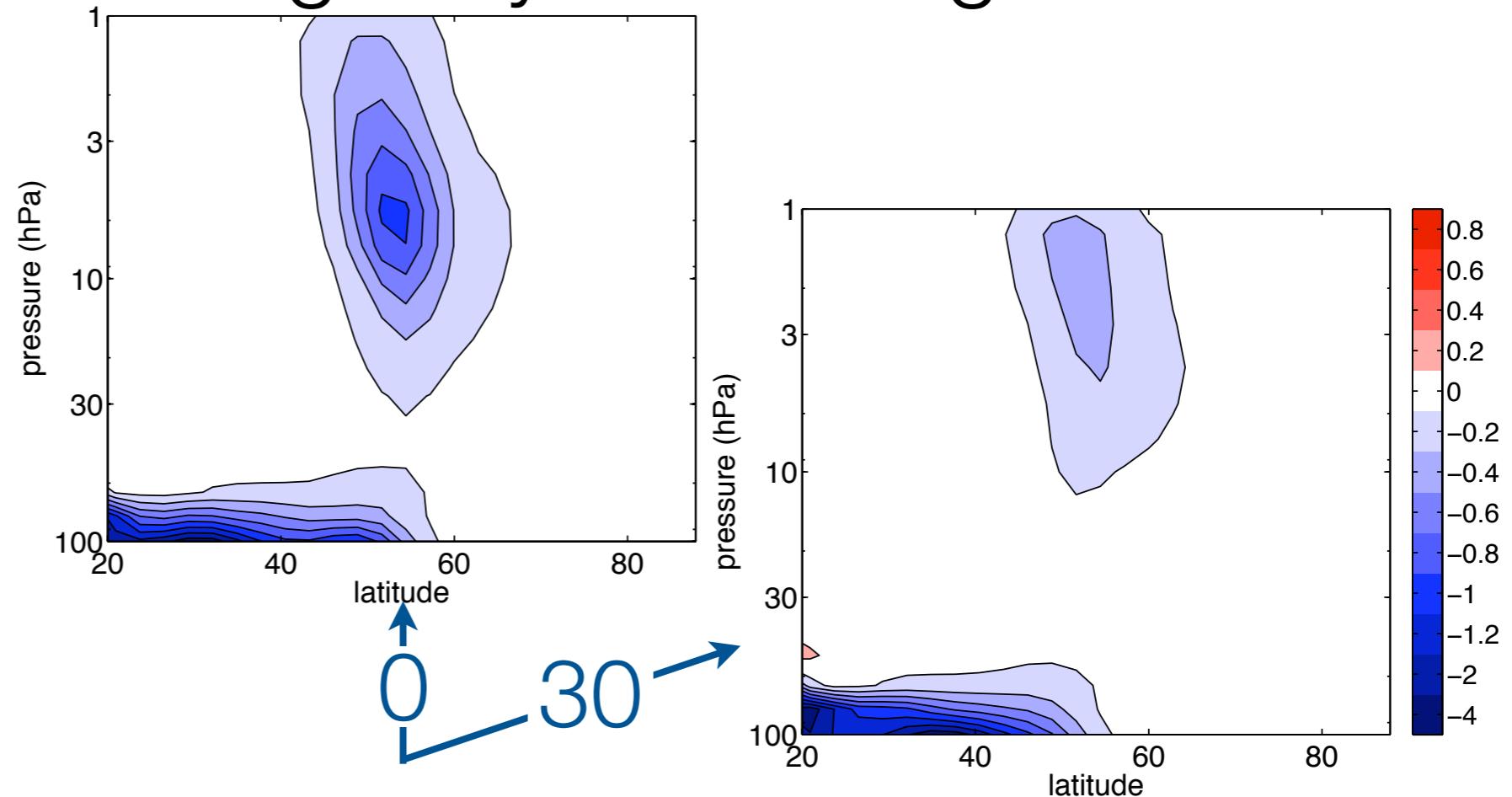
(for phase shift = 0 integration,
10,000 day mean, perpetual January)

Vary phase of subgrid-scale topography



phase angle between
subgrid-scale and
resolved topography

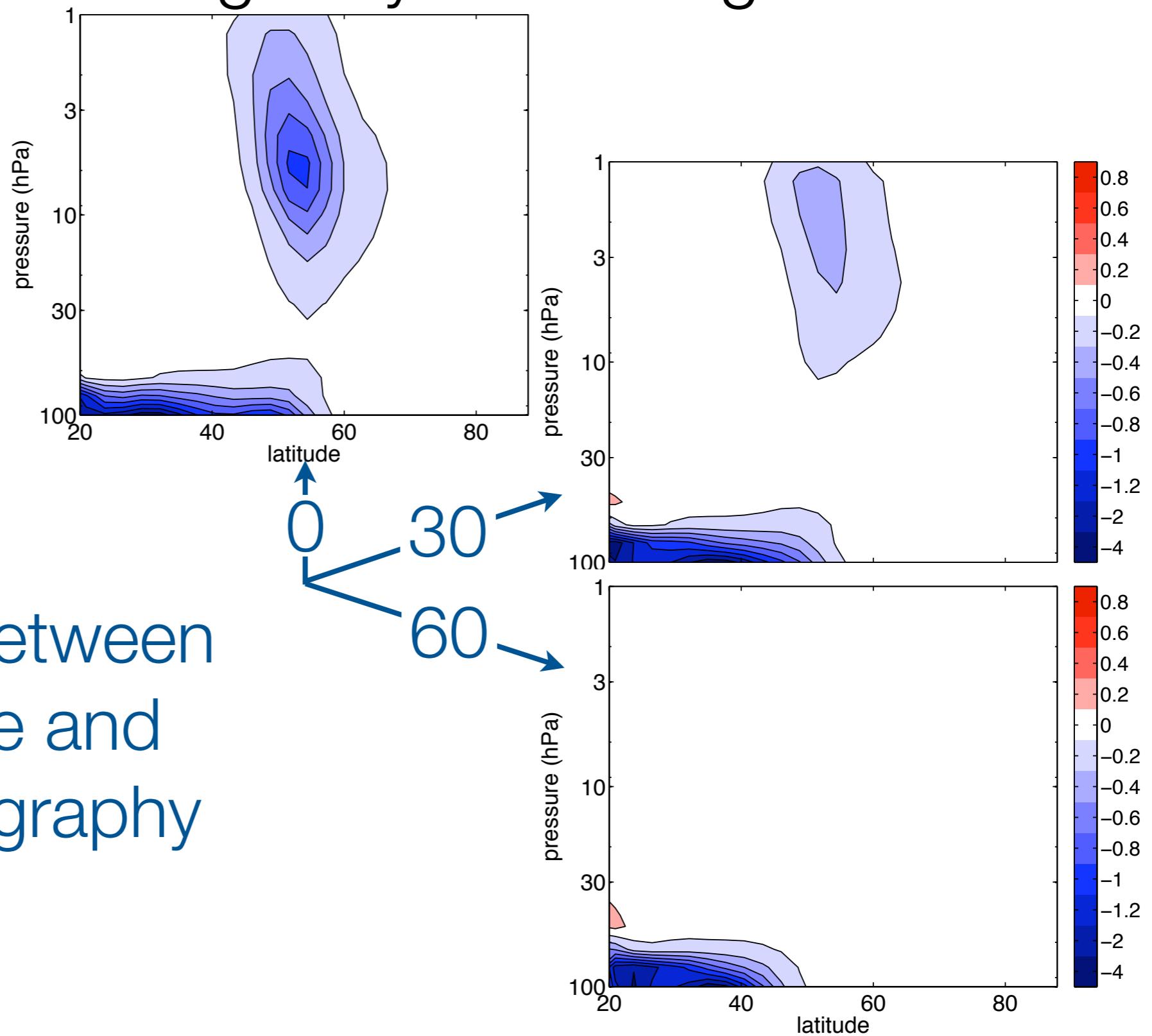
Large perturbations in gravity wave drag ...



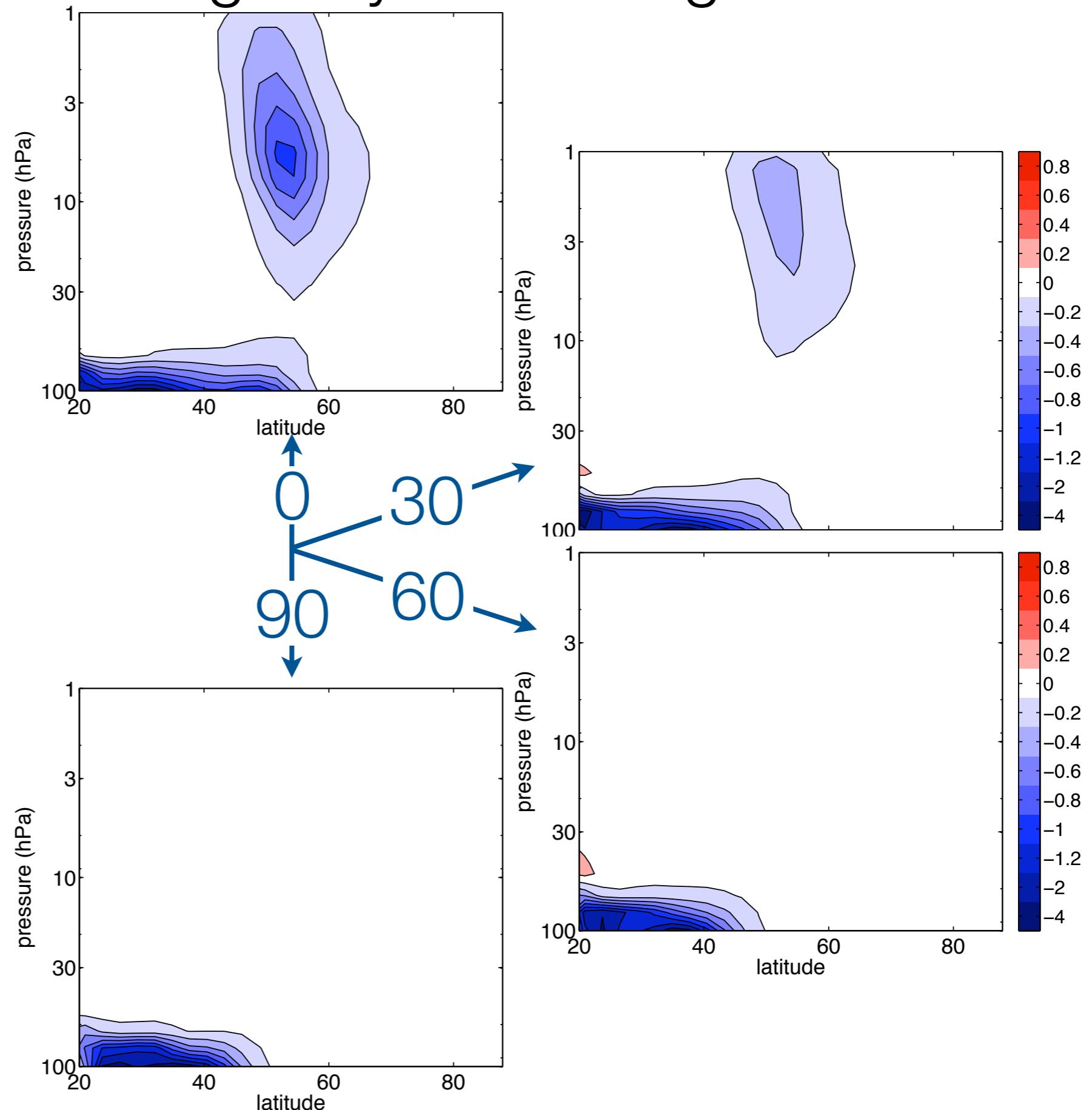
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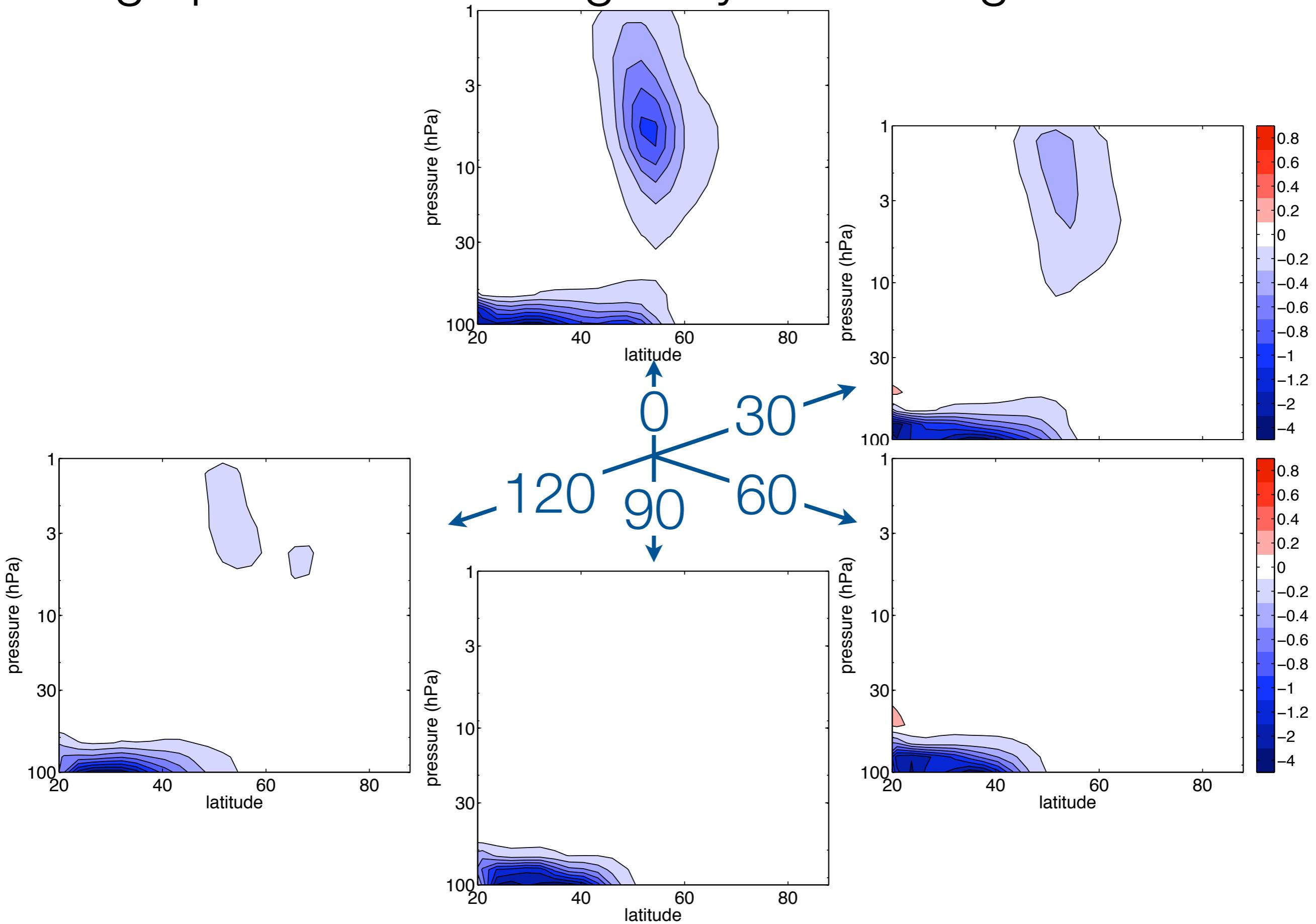
phase angle between
subgrid-scale and
resolved topography



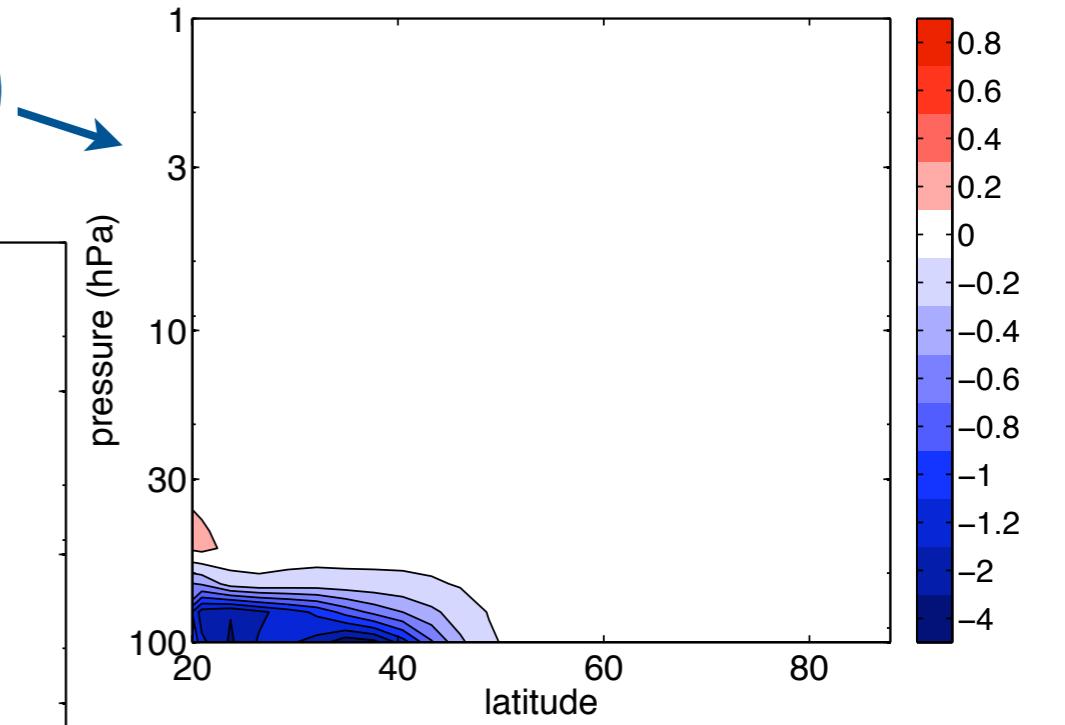
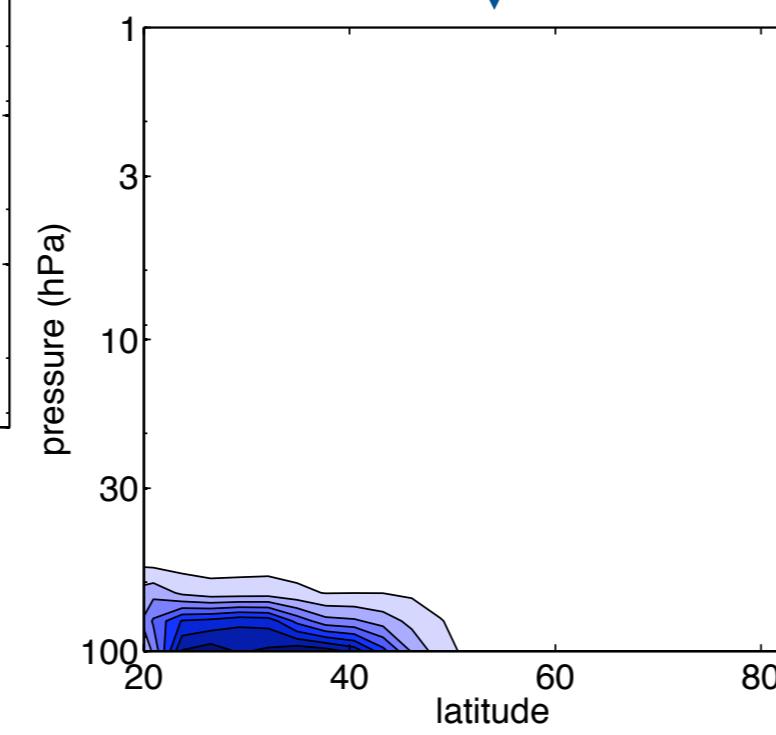
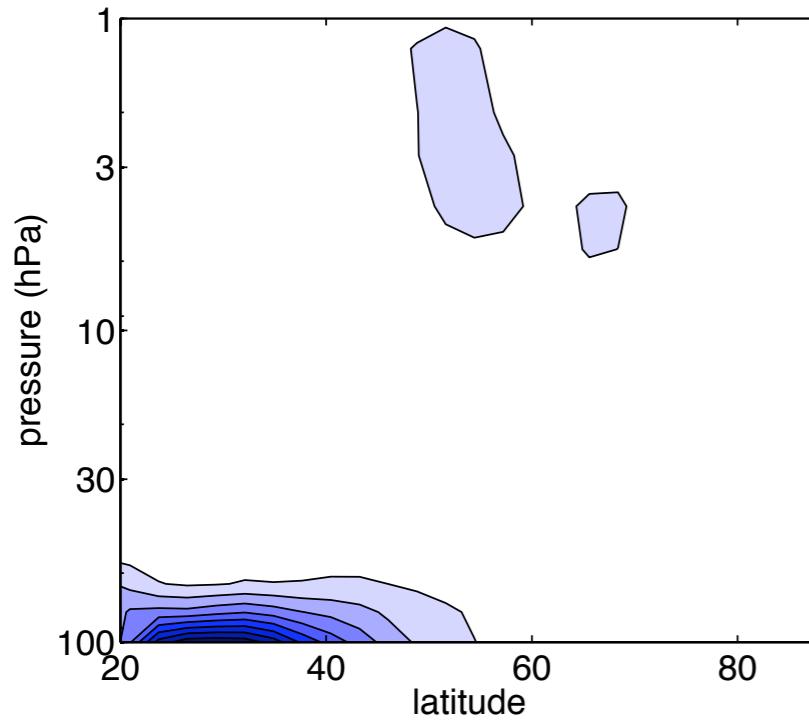
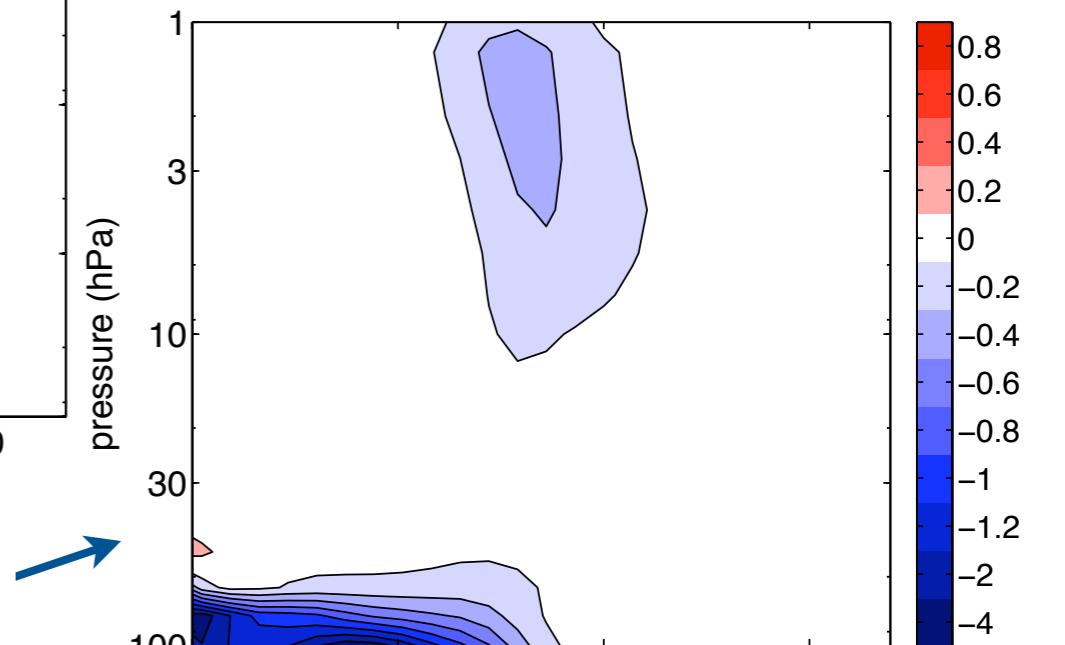
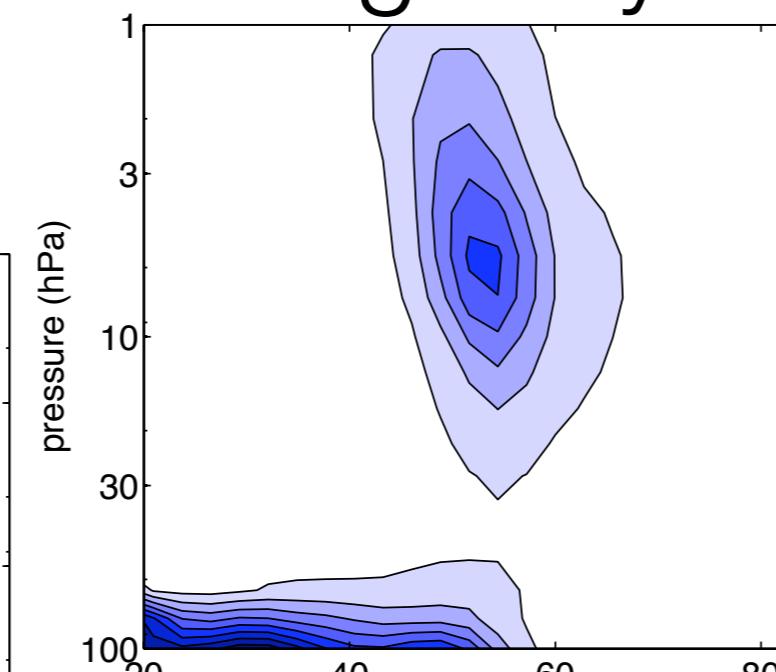
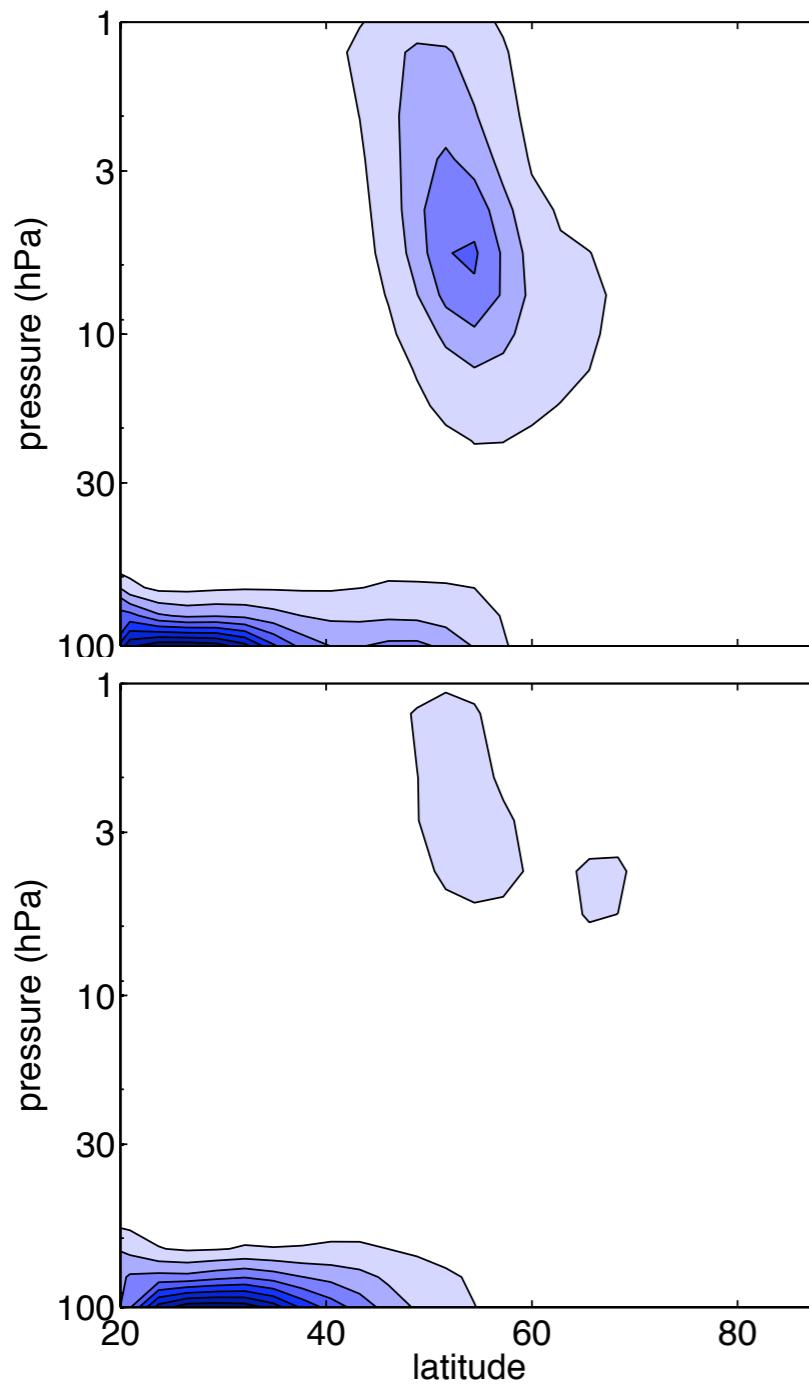
Large perturbations in gravity wave drag ...



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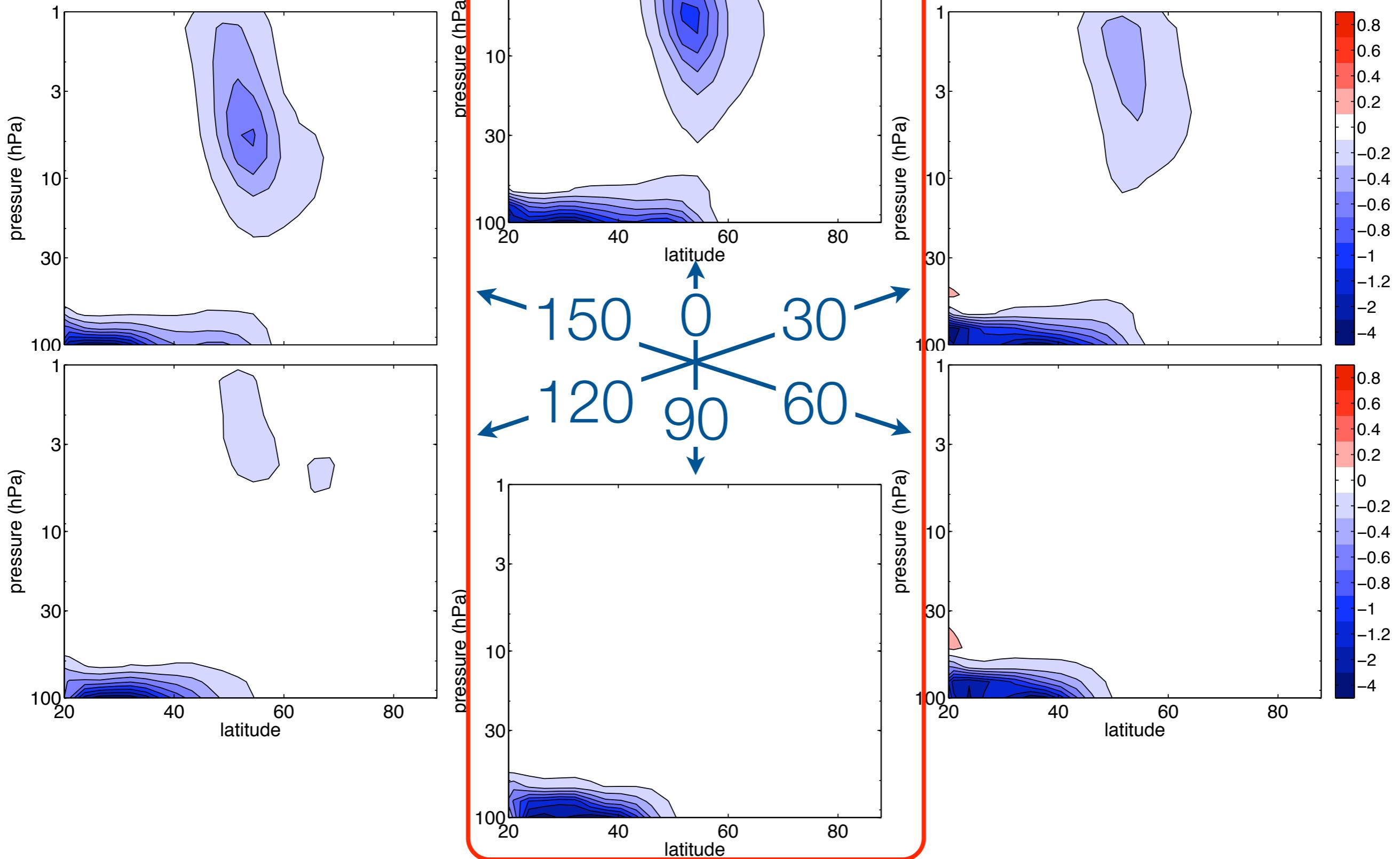


150
120
90
60
30

0

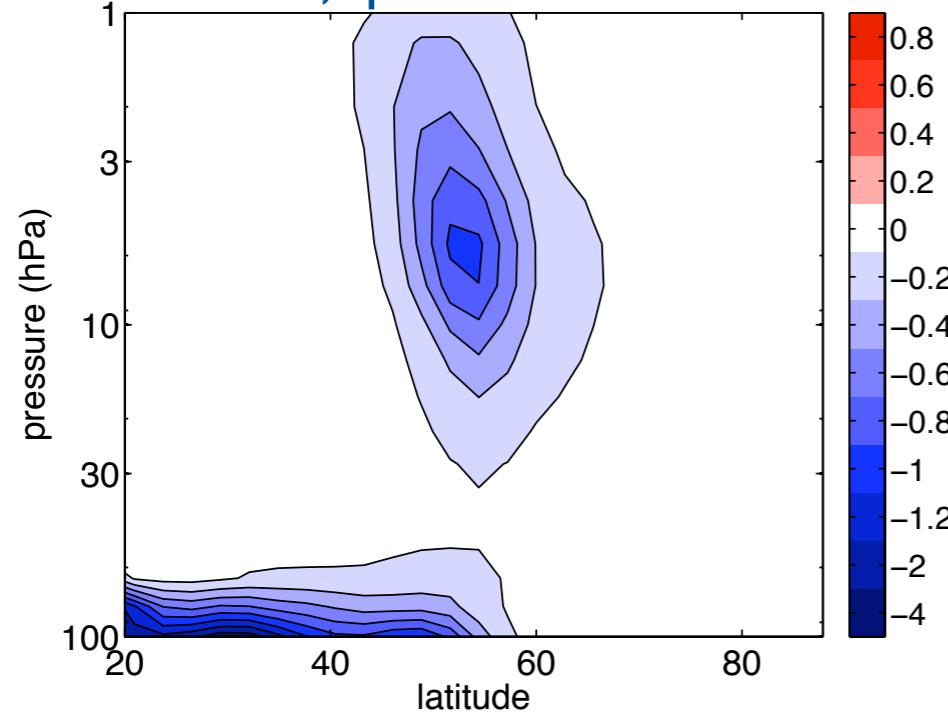
latitude

Large perturbations in gravity wave drag ...

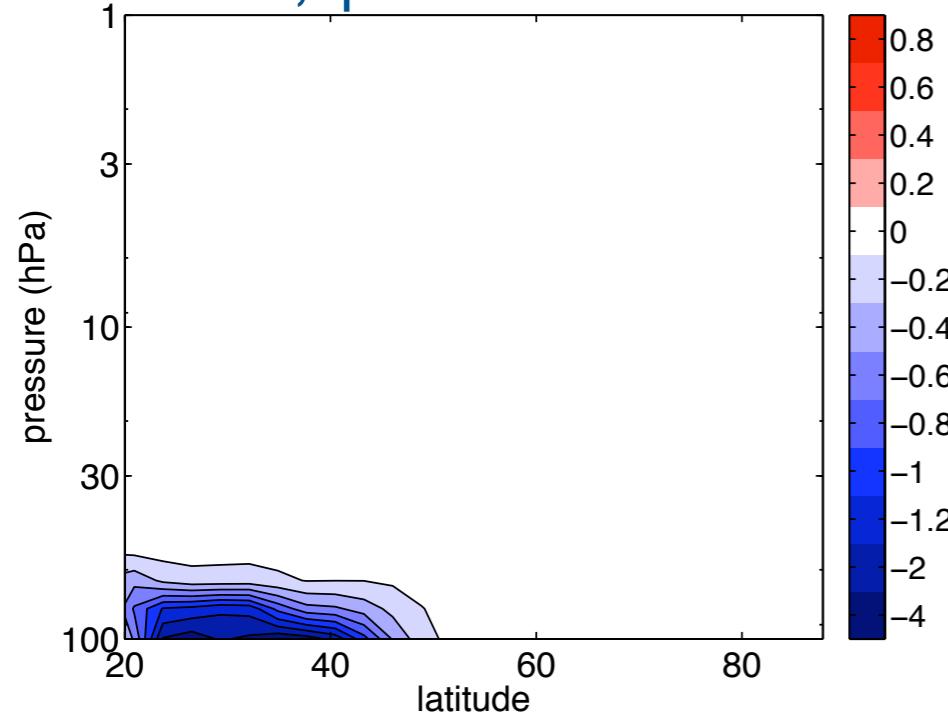


Difference in gravity wave drag

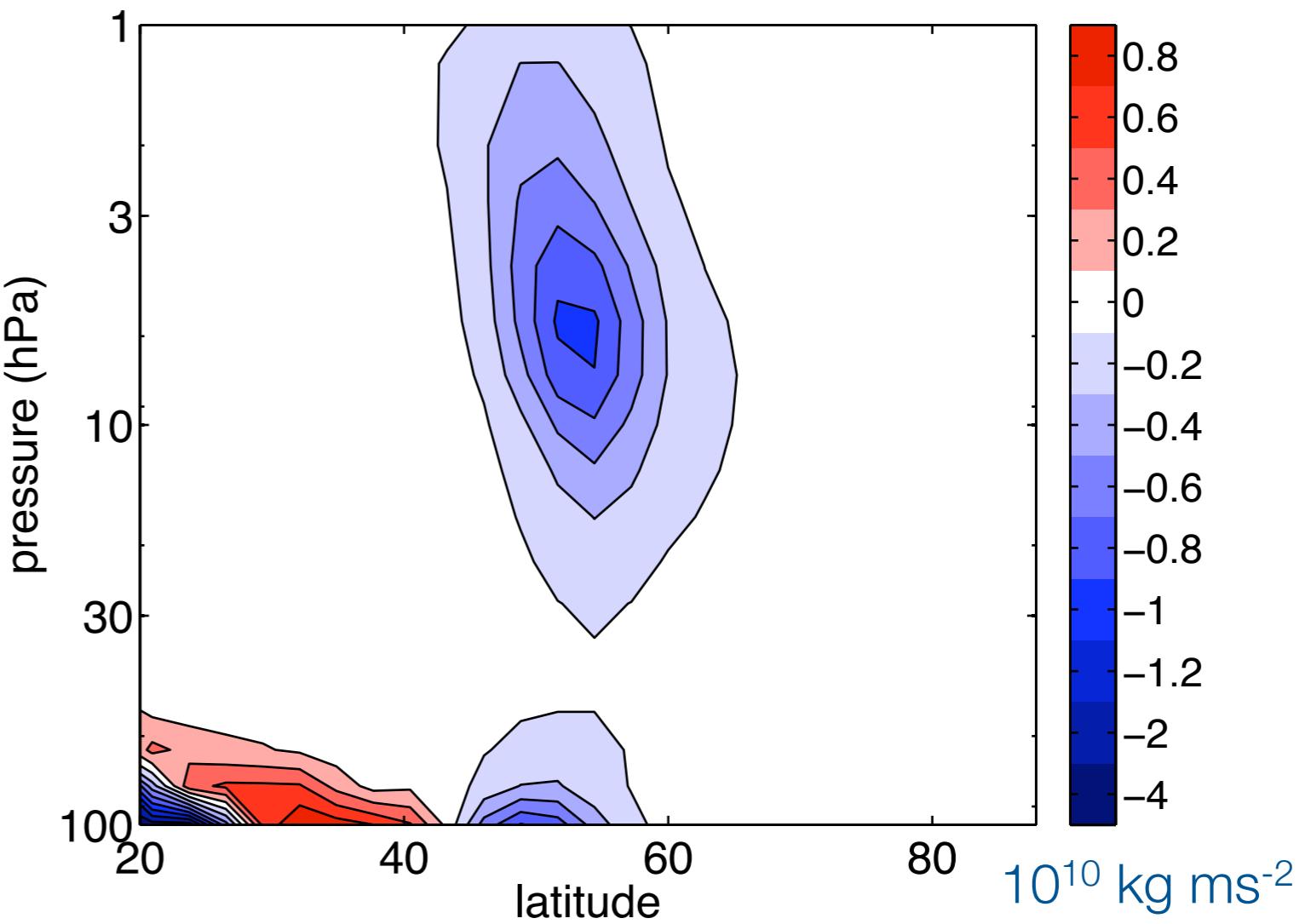
OGWD, phase shift = 0



OGWD, phase shift = 90

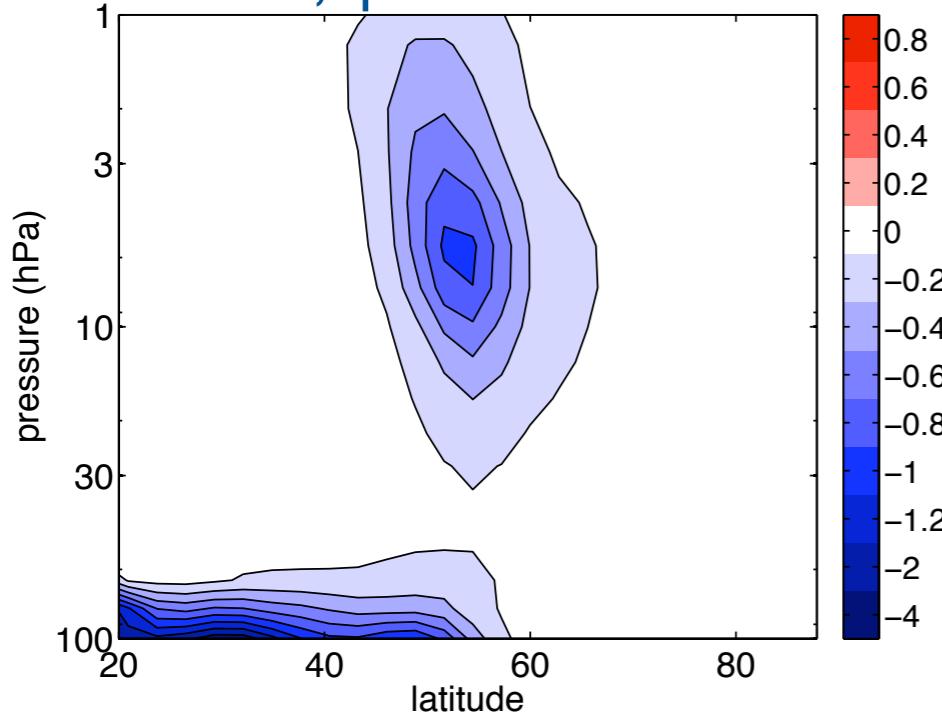


Difference in OGWP Torque

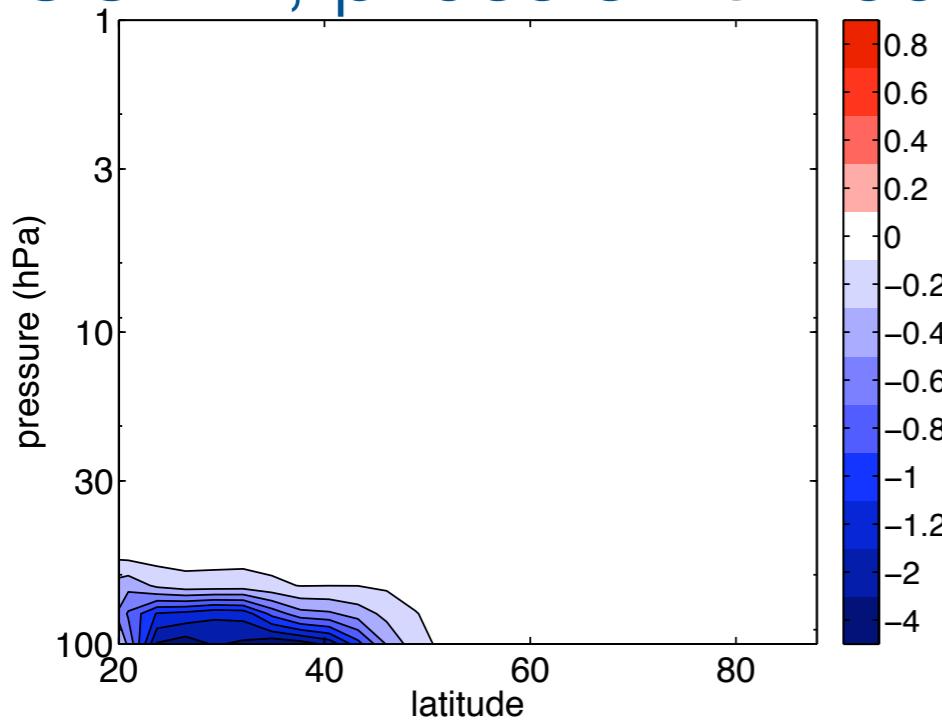


Difference in gravity wave drag

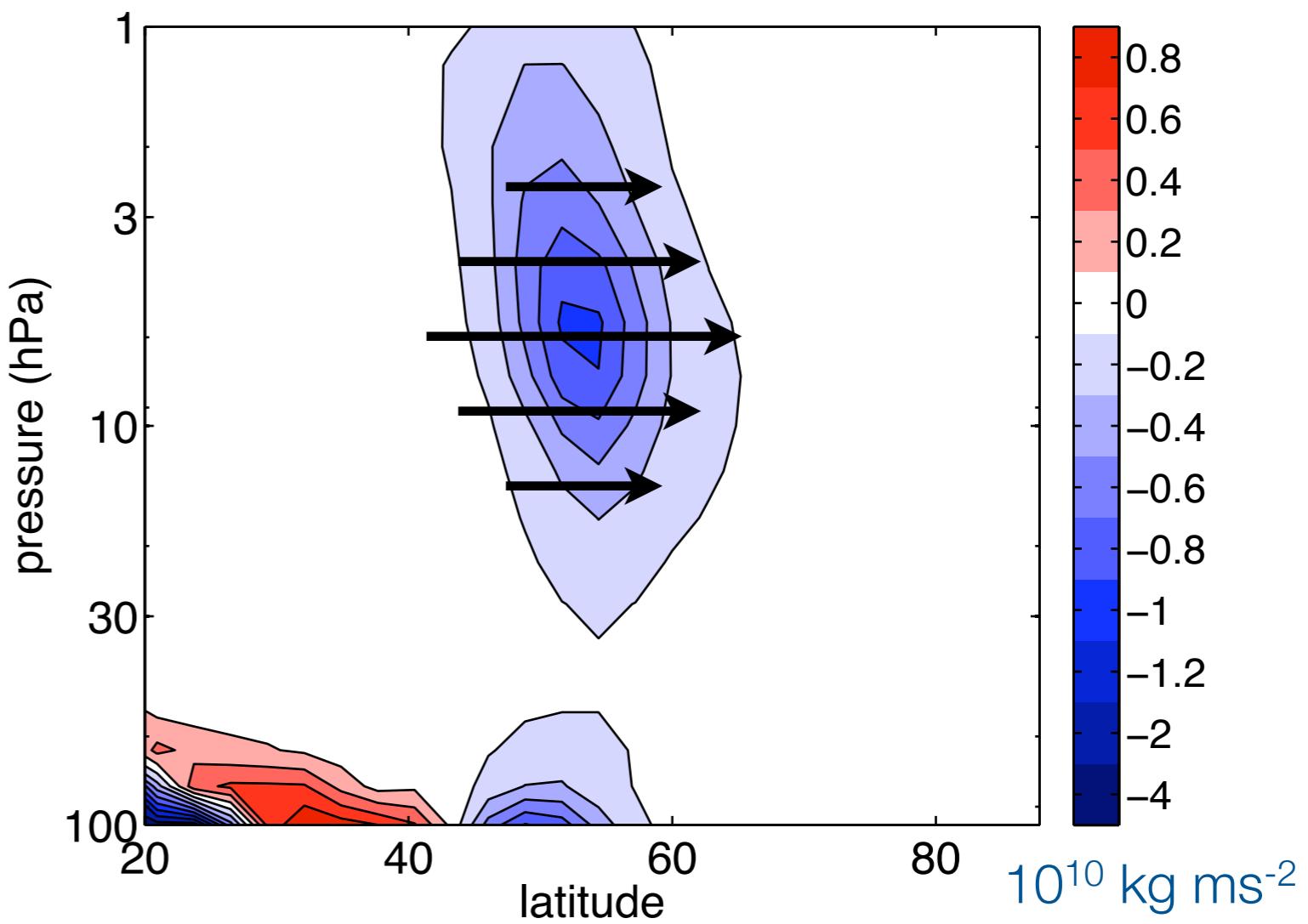
OGWD, phase shift = 0



OGWD, phase shift = 90

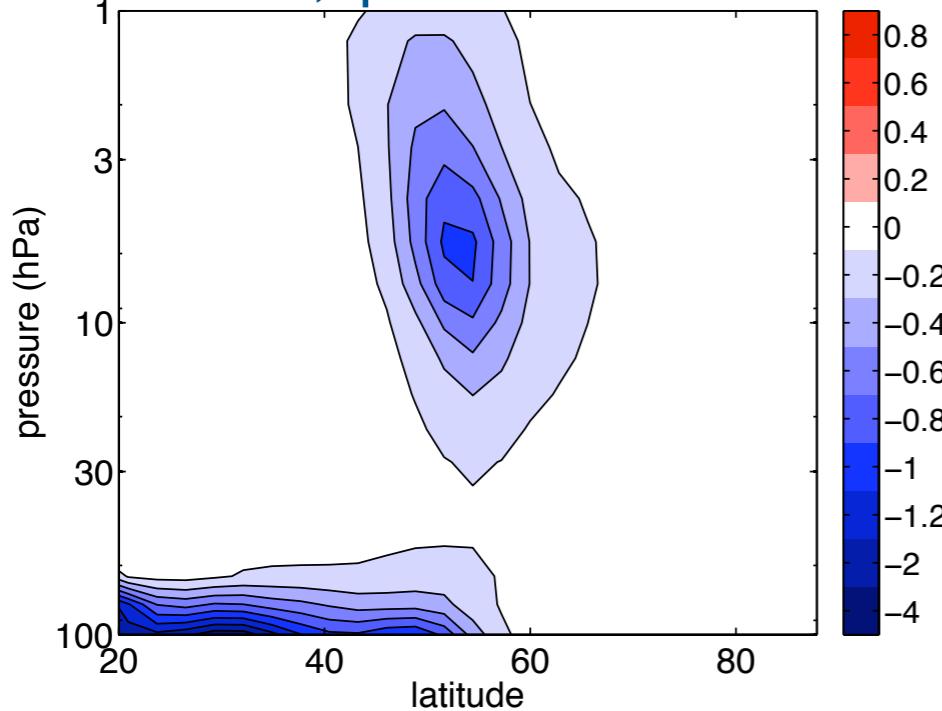


In steady state, torque
must be balanced by
meridional flow ...

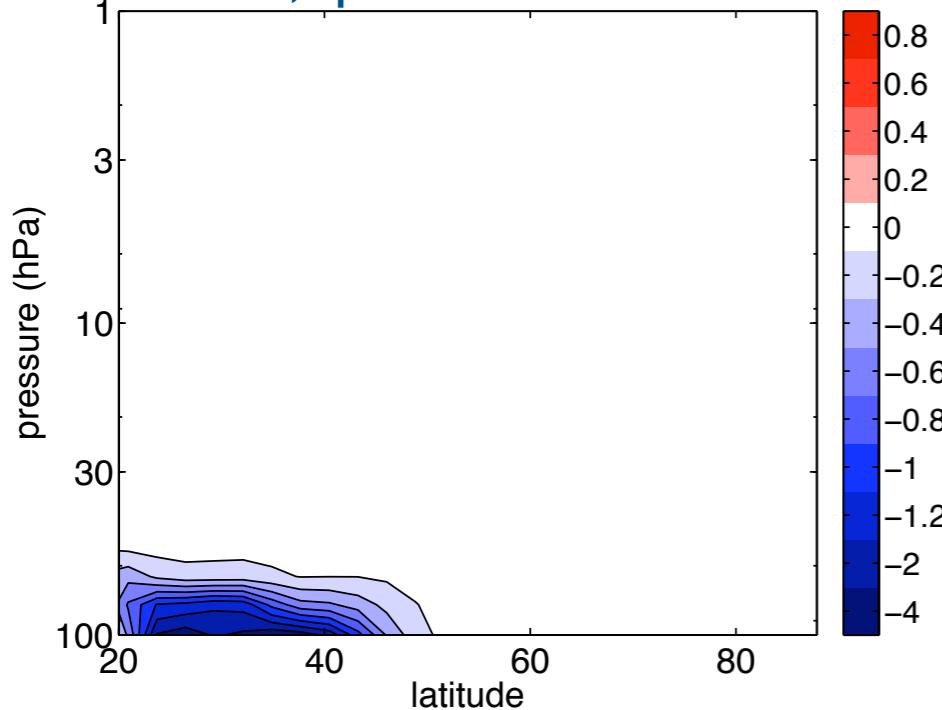


Difference in gravity wave drag
should imply significant change in circulation

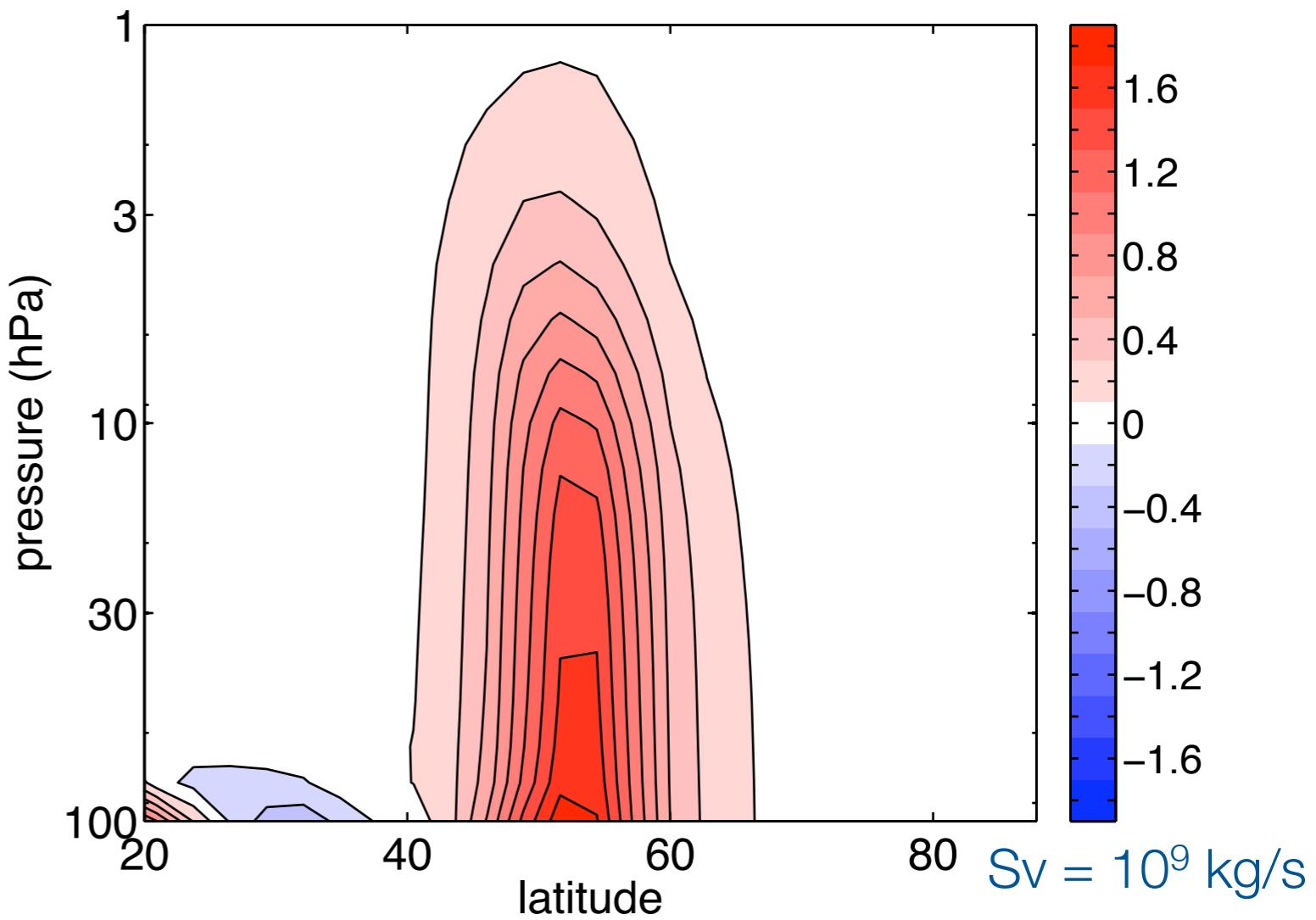
OGWD, phase shift = 0



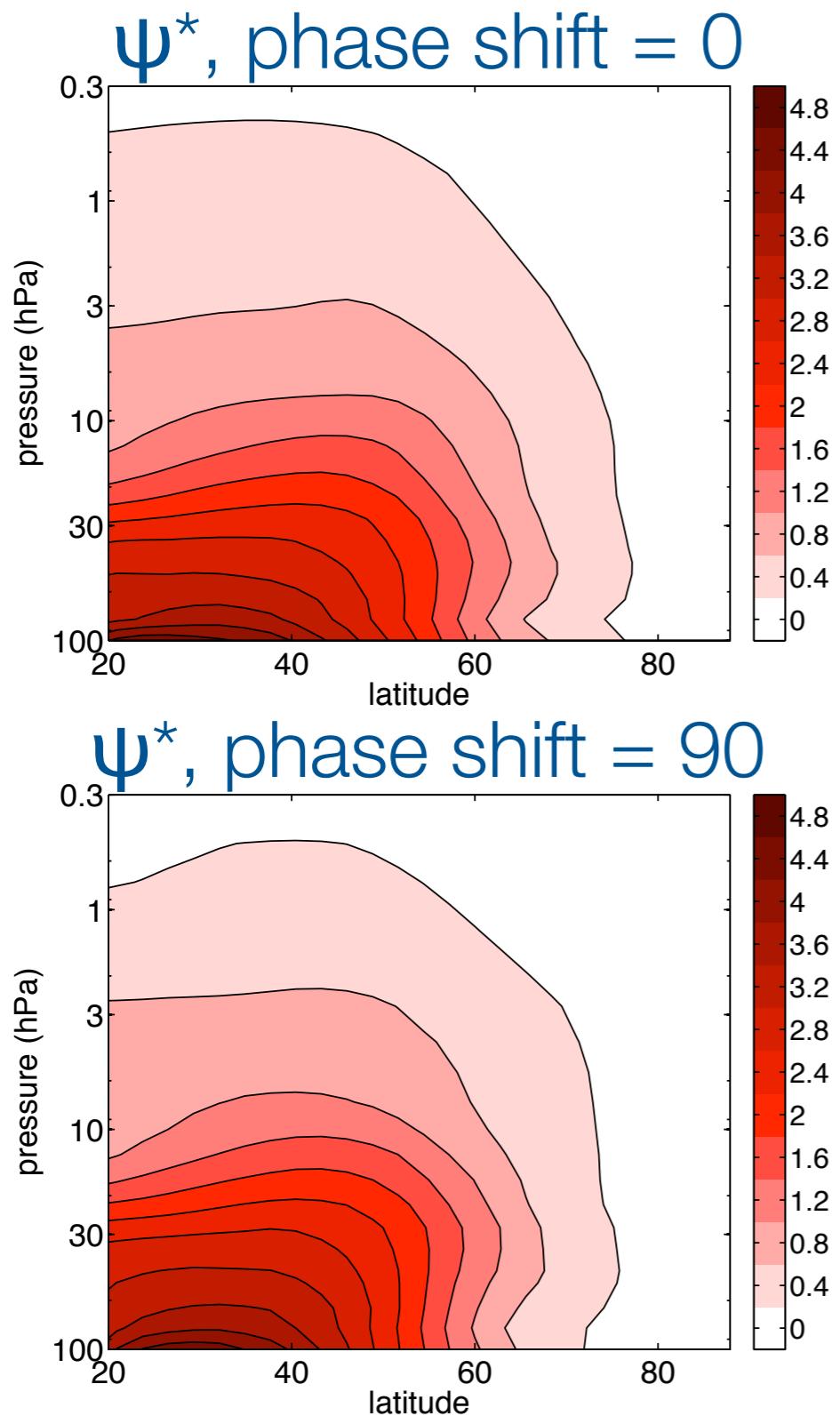
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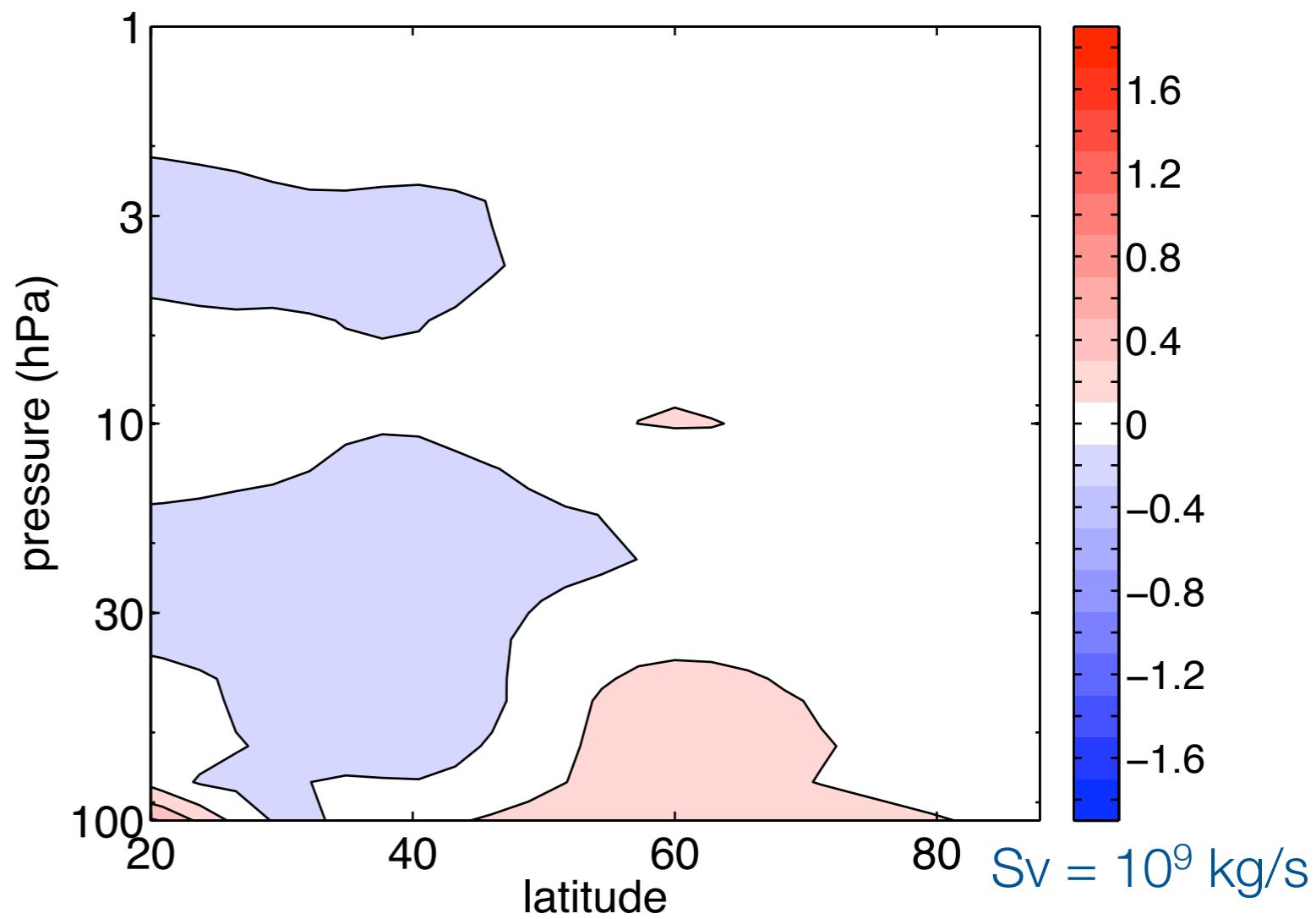
$\Delta\Phi_{DC}^*$,
suggested by DWC



... but, there's hardly any difference in the actual residual mean circulation!

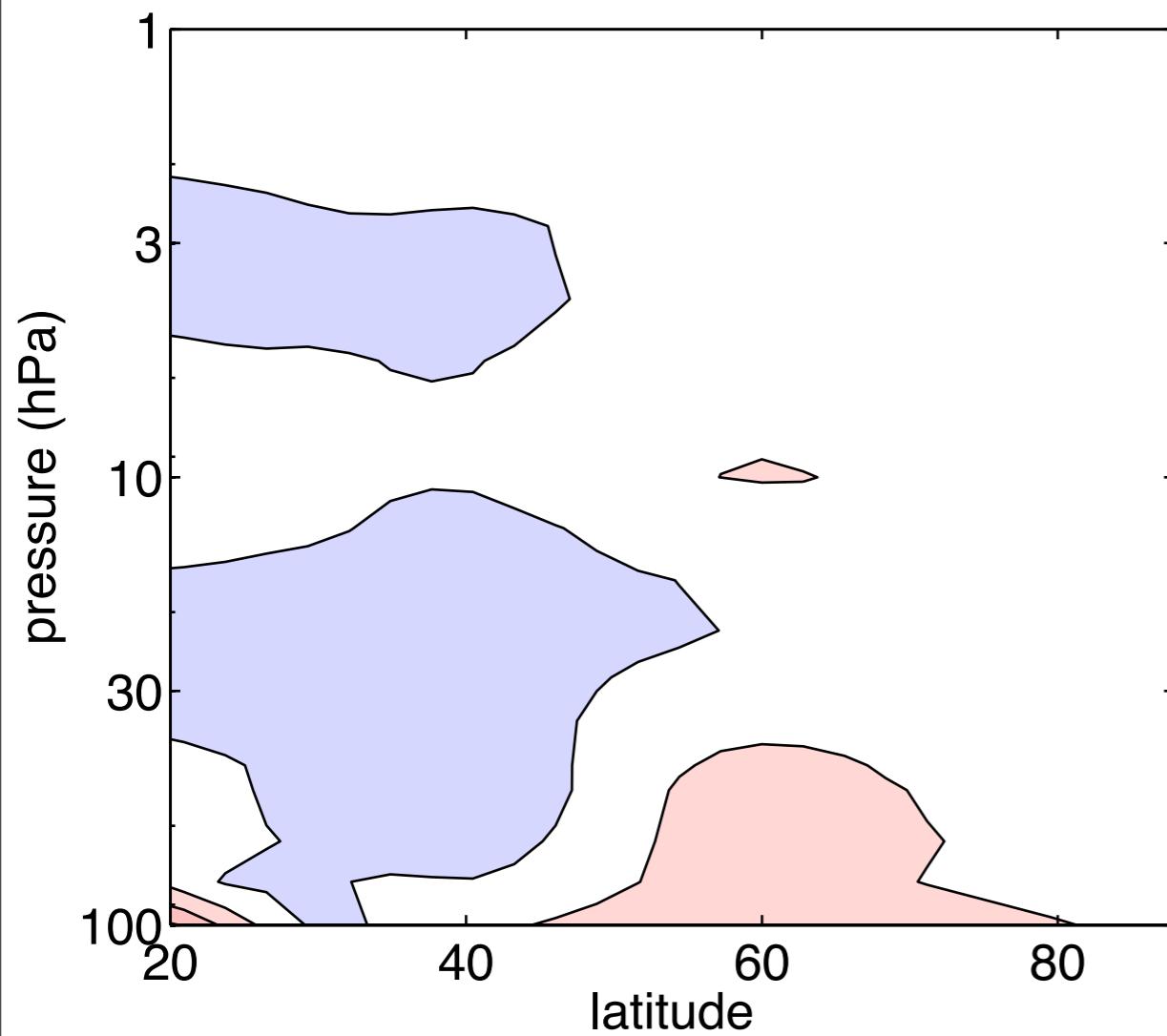


$\Delta\Psi^*$, the difference in the TEM circulation

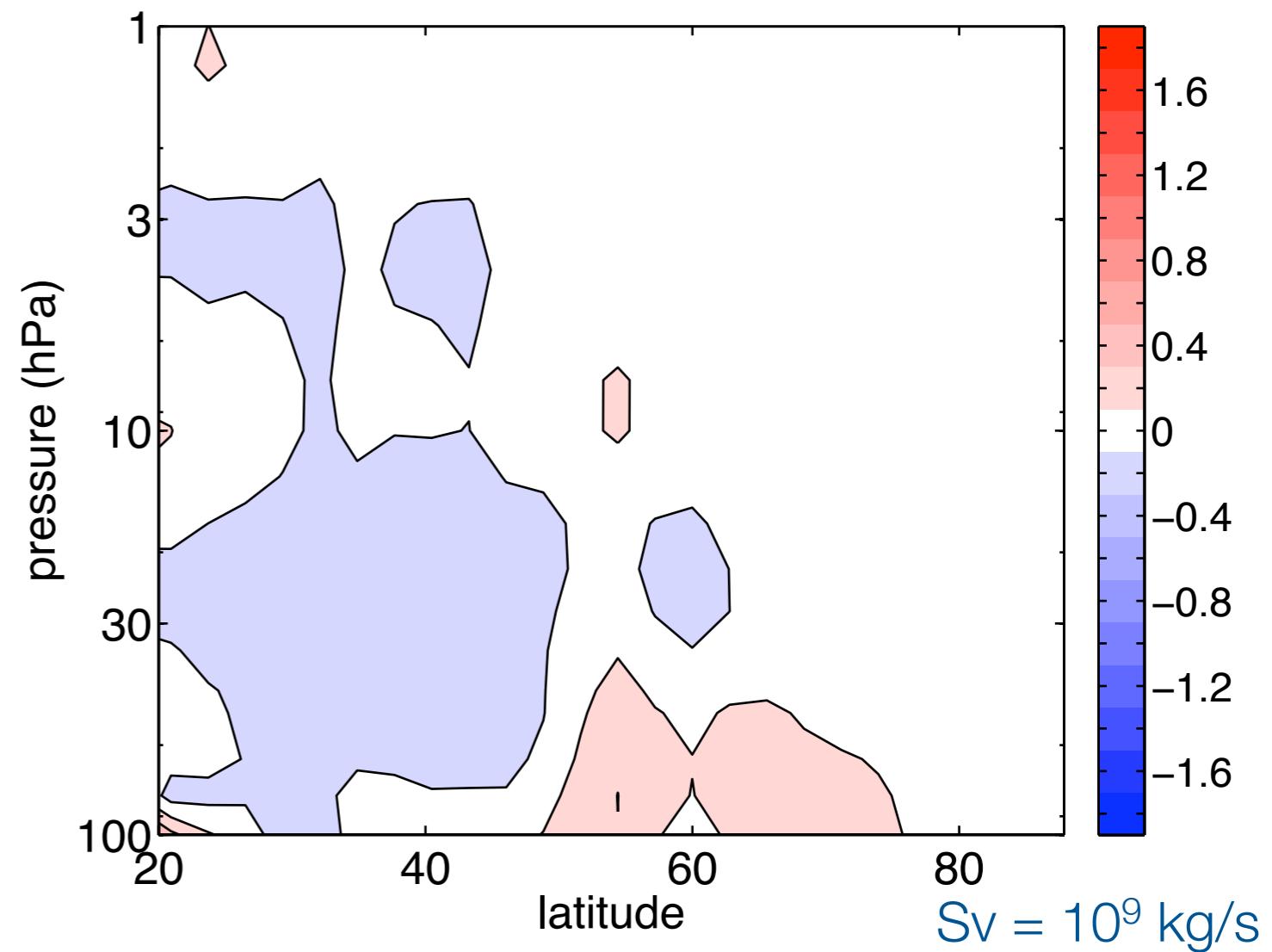


Downward control is not the problem ...

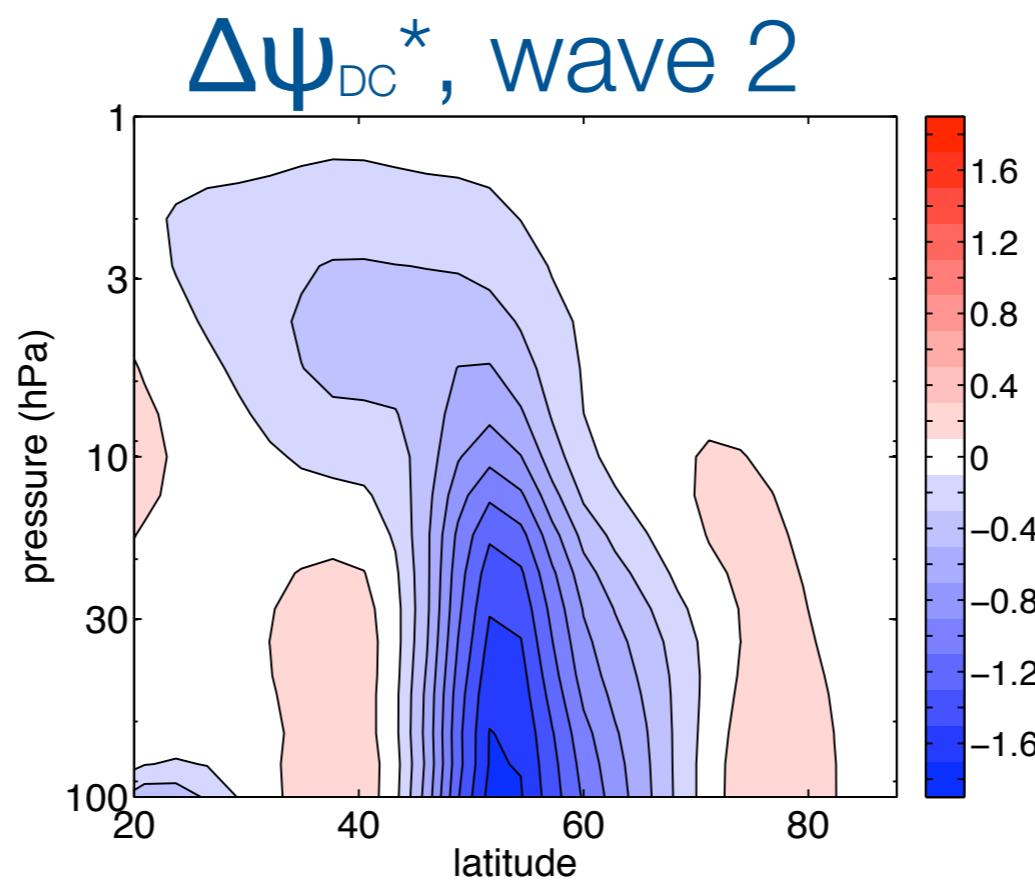
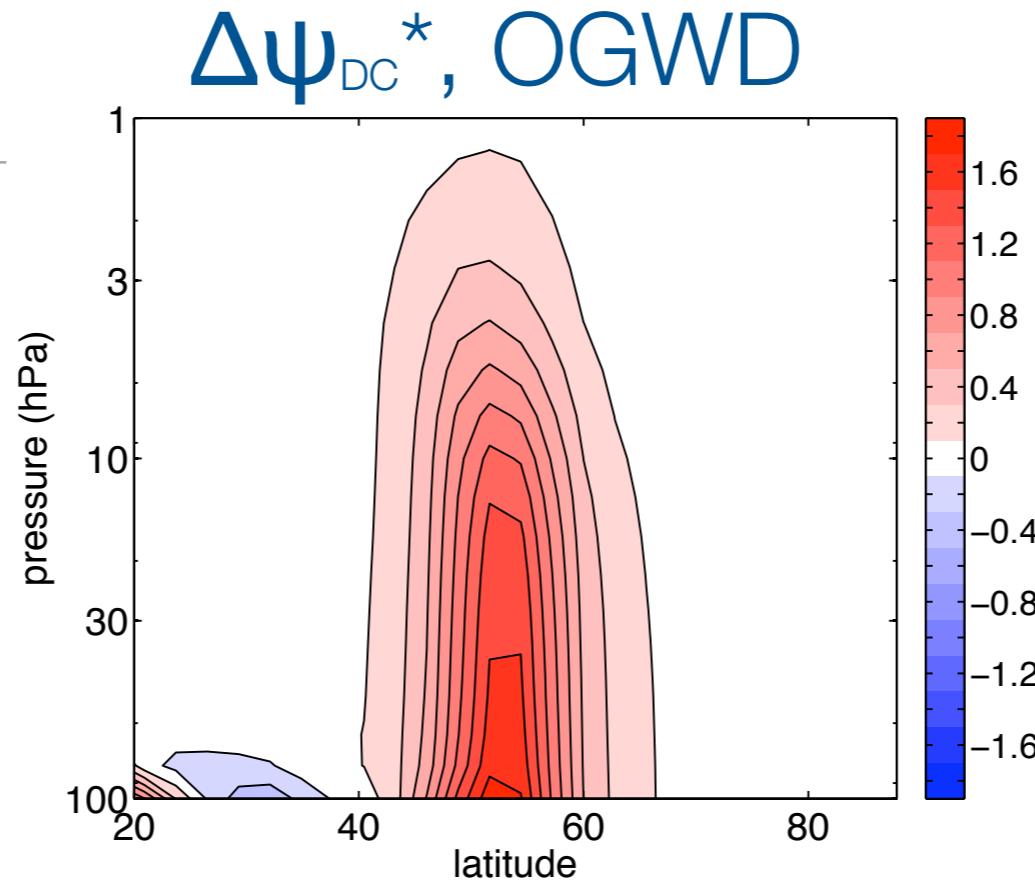
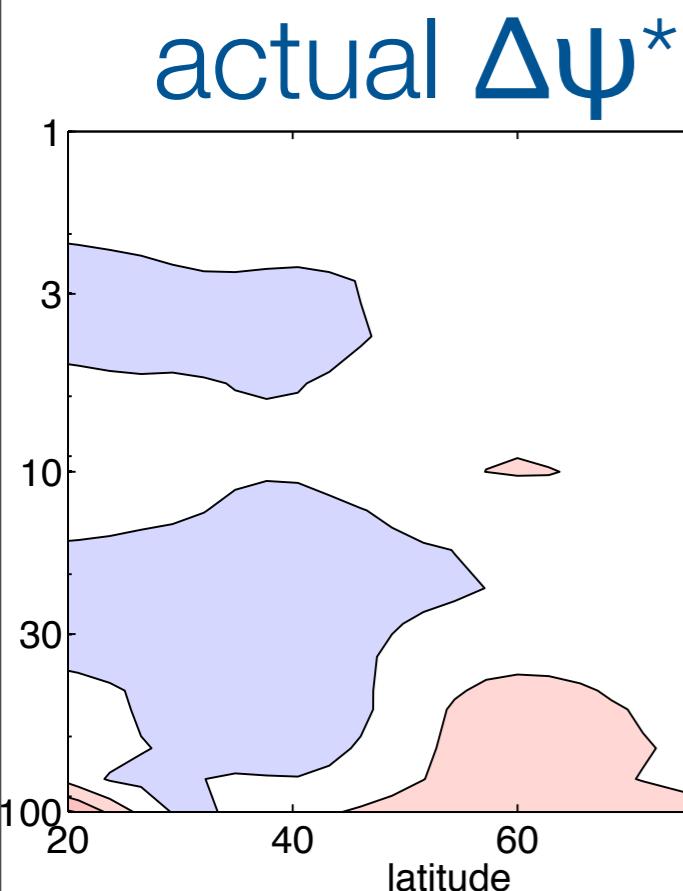
Observed $\Delta\psi^*$



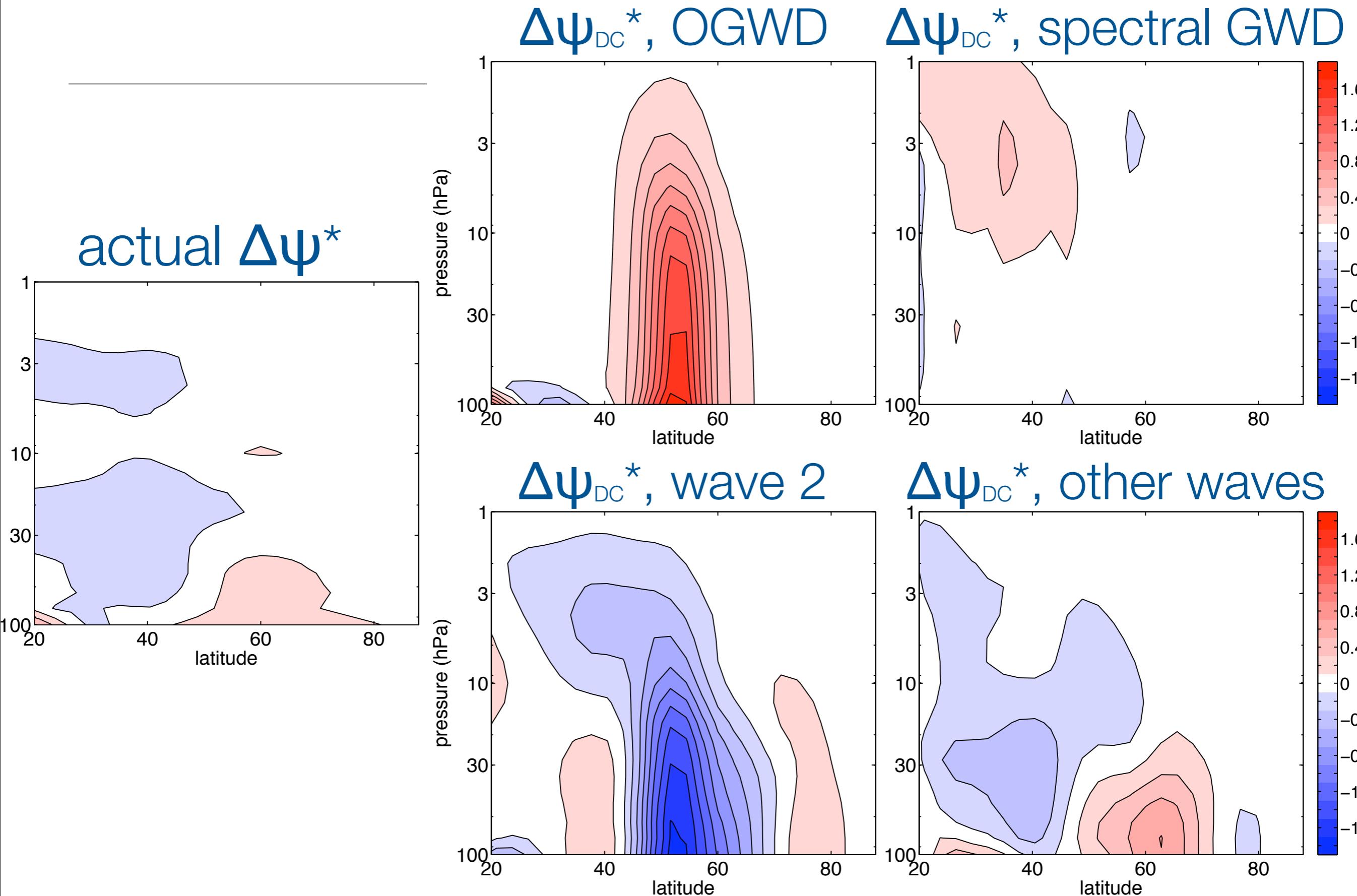
$\Delta\psi_{DC}^*$, computed
with downward control



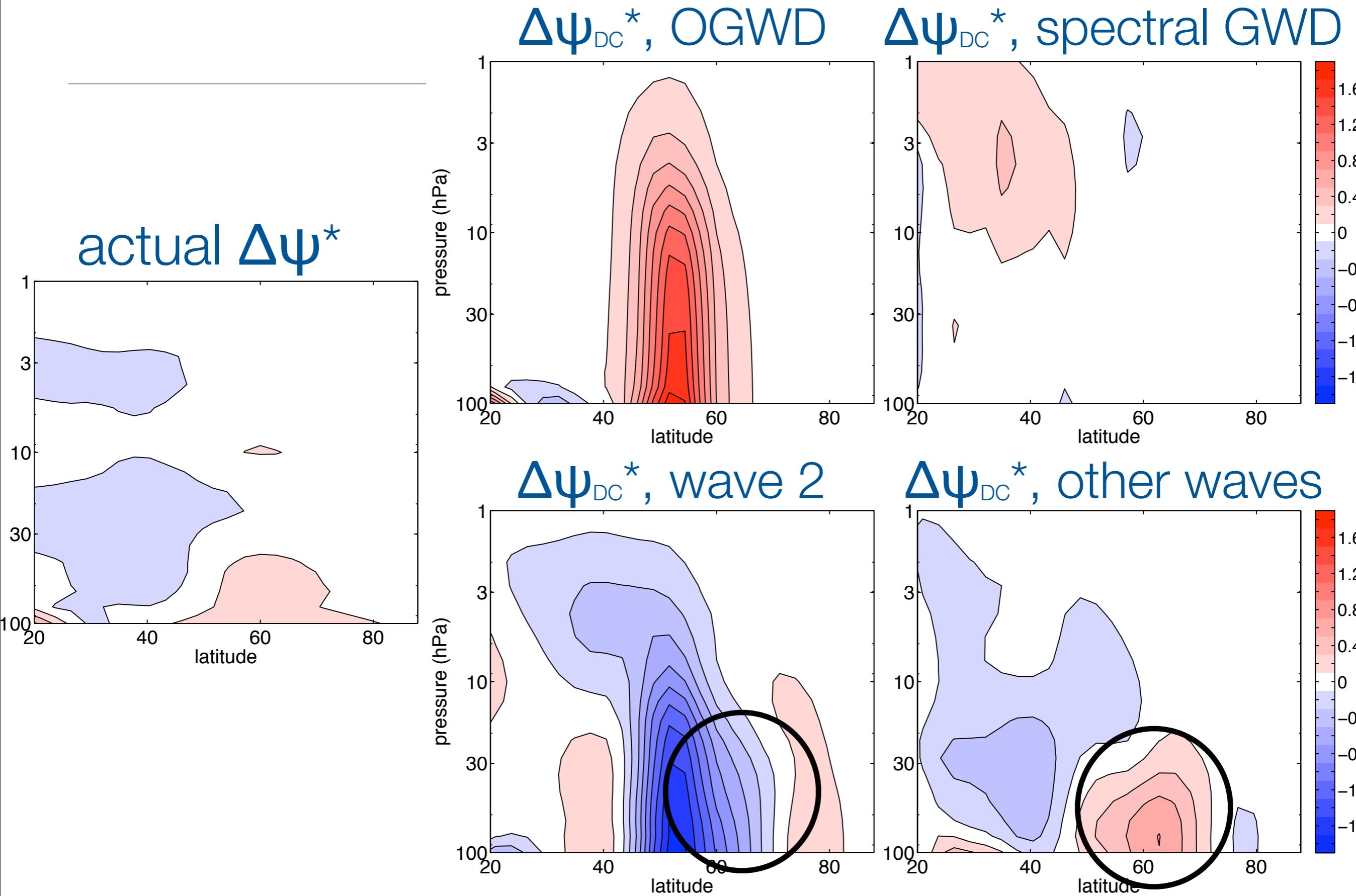
Rather, rest of circulation compensates for OGWD



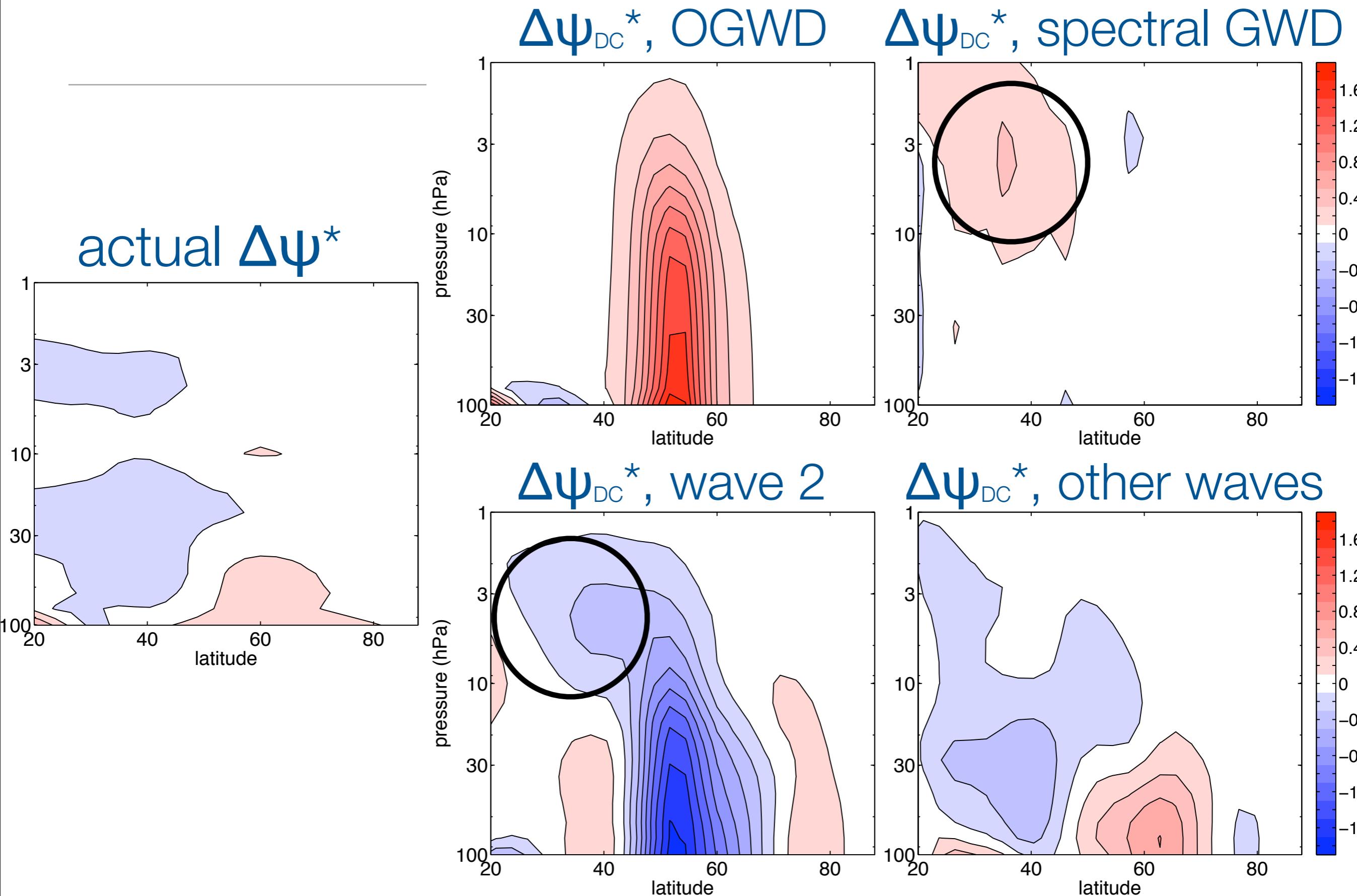
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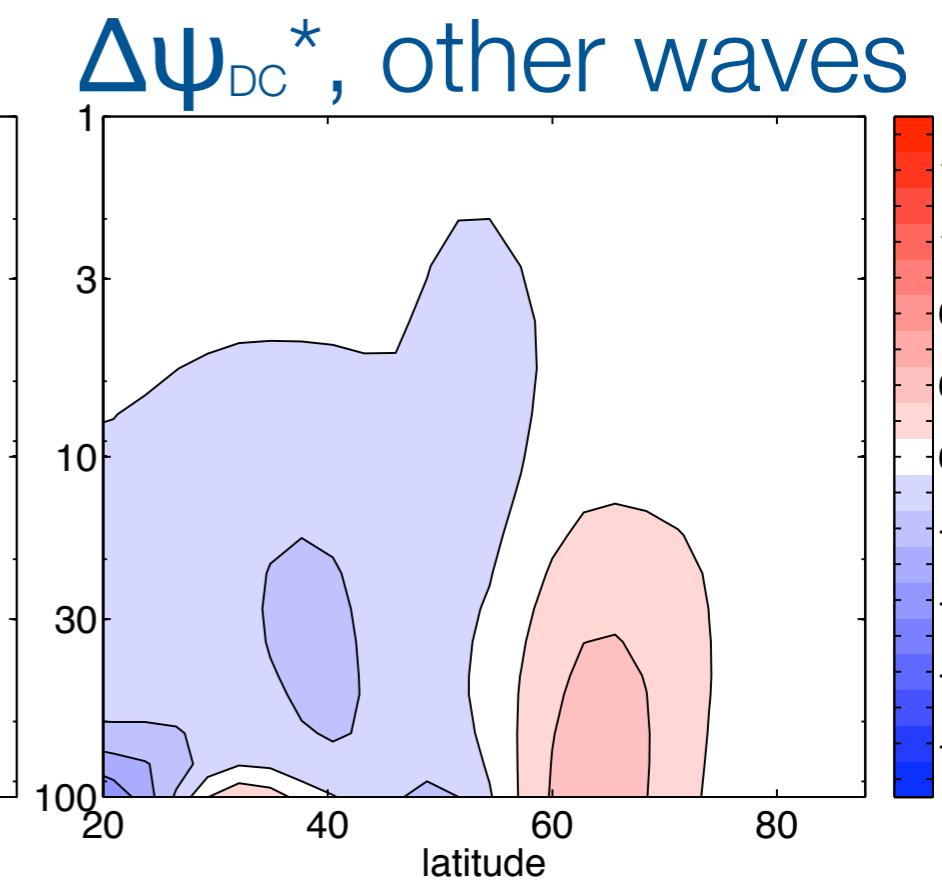
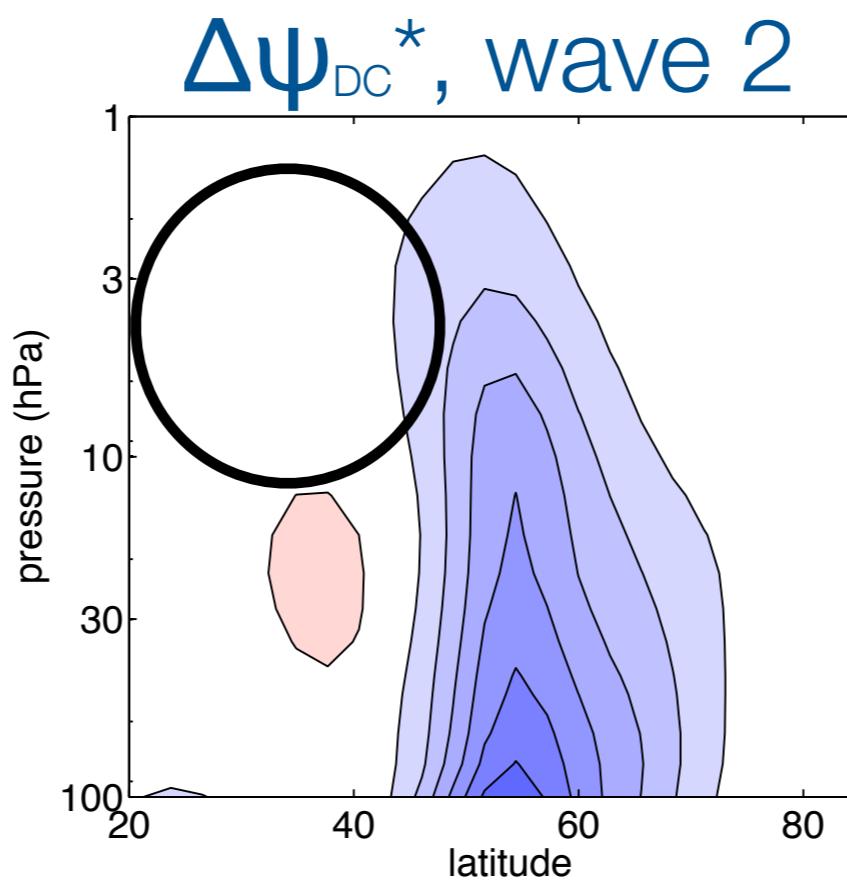
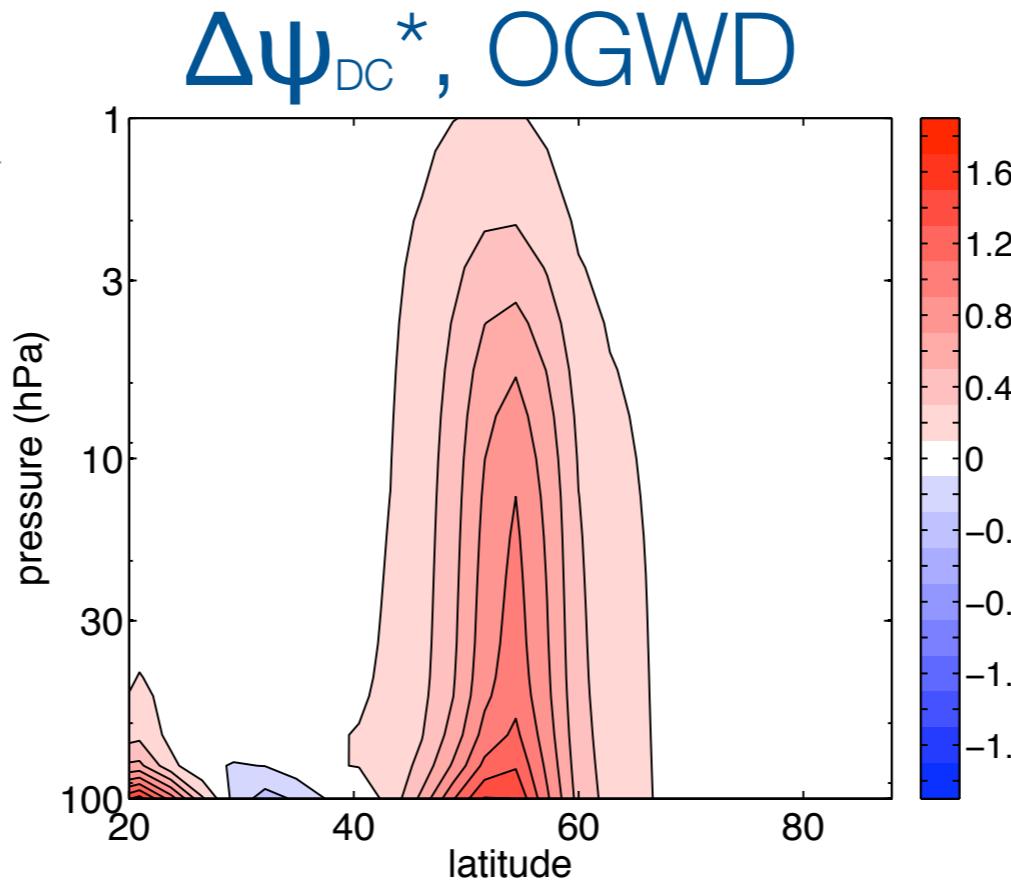
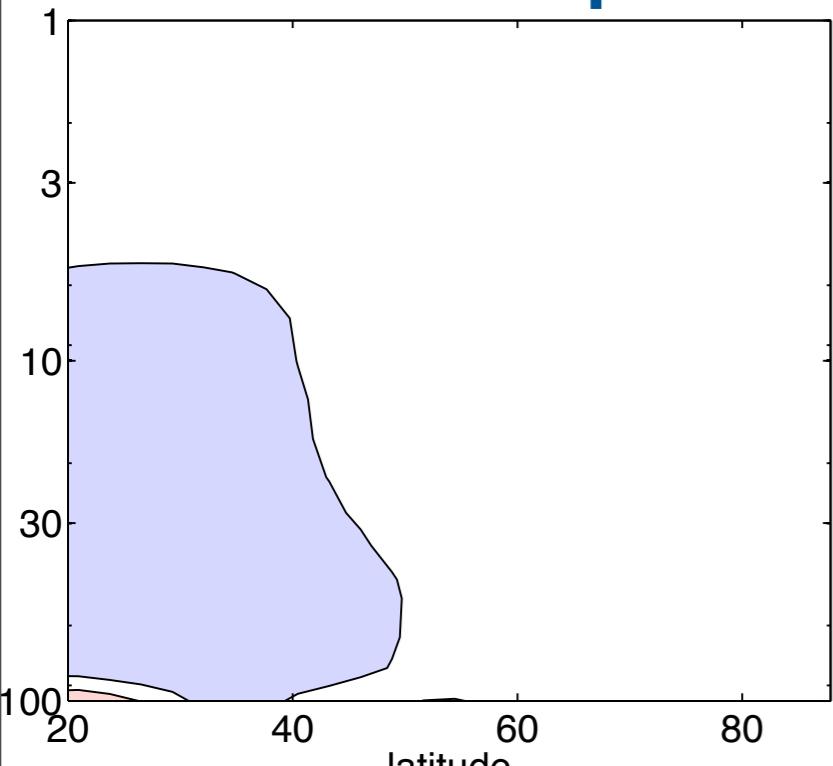
Rather, rest of circulation compensates for OGWD



Aside: same experiment w/o spectral GWP

OGWD

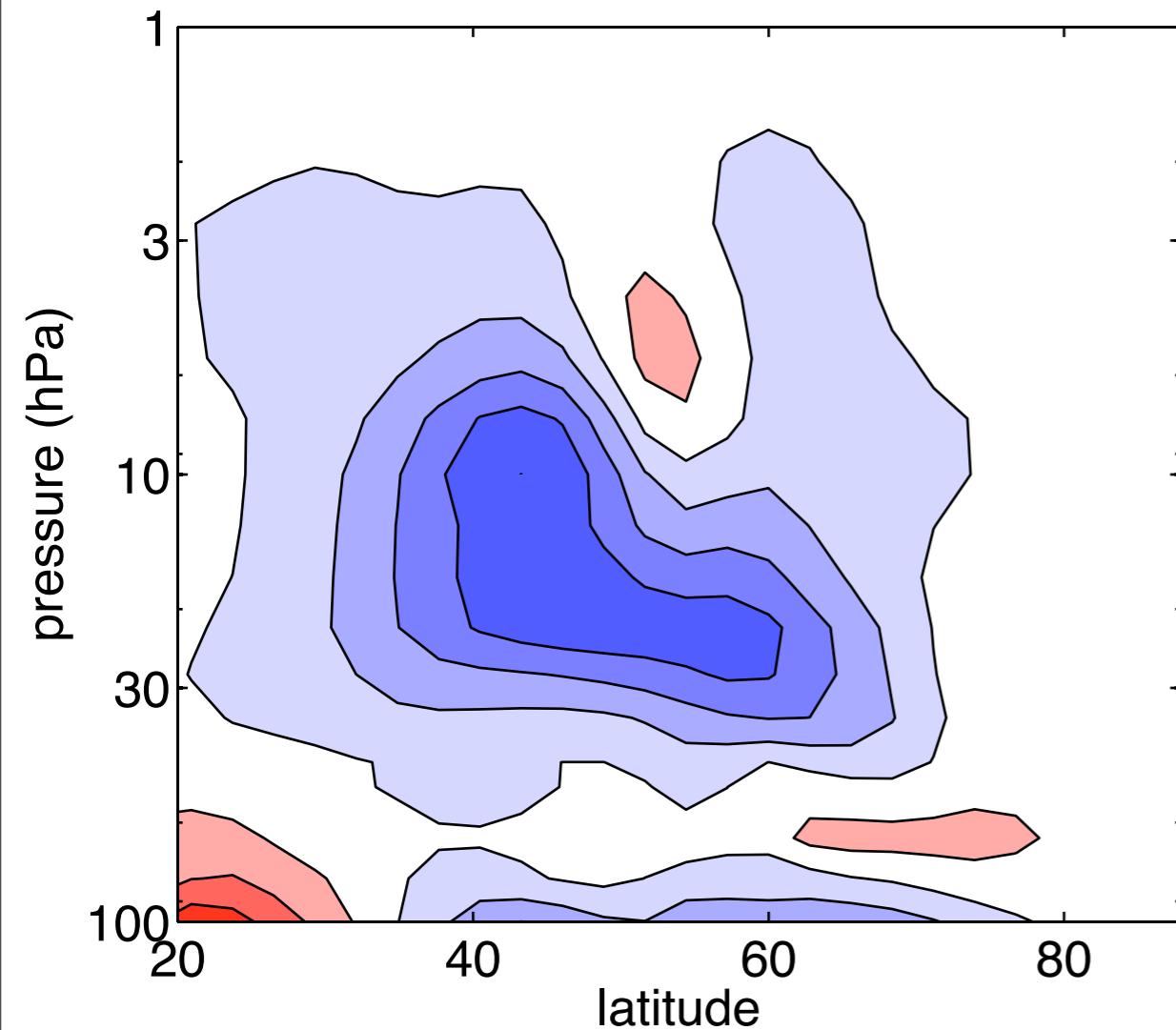
actual $\Delta\psi^*$



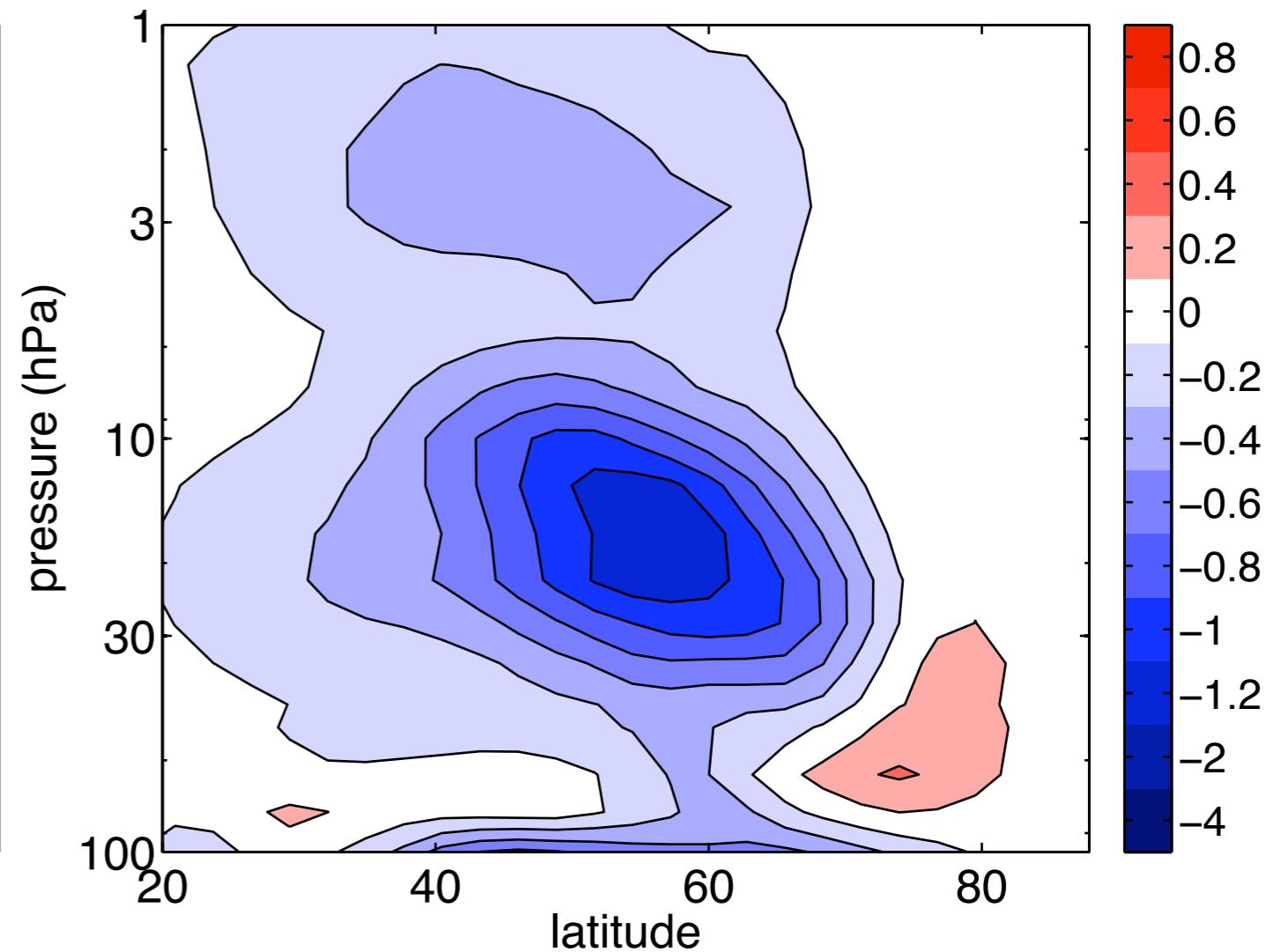
Possible Mechanisms

$\nabla \cdot F$, wavenumber 2 ($10^{10} \text{ kg ms}^{-2}$)

phase shift = 0°



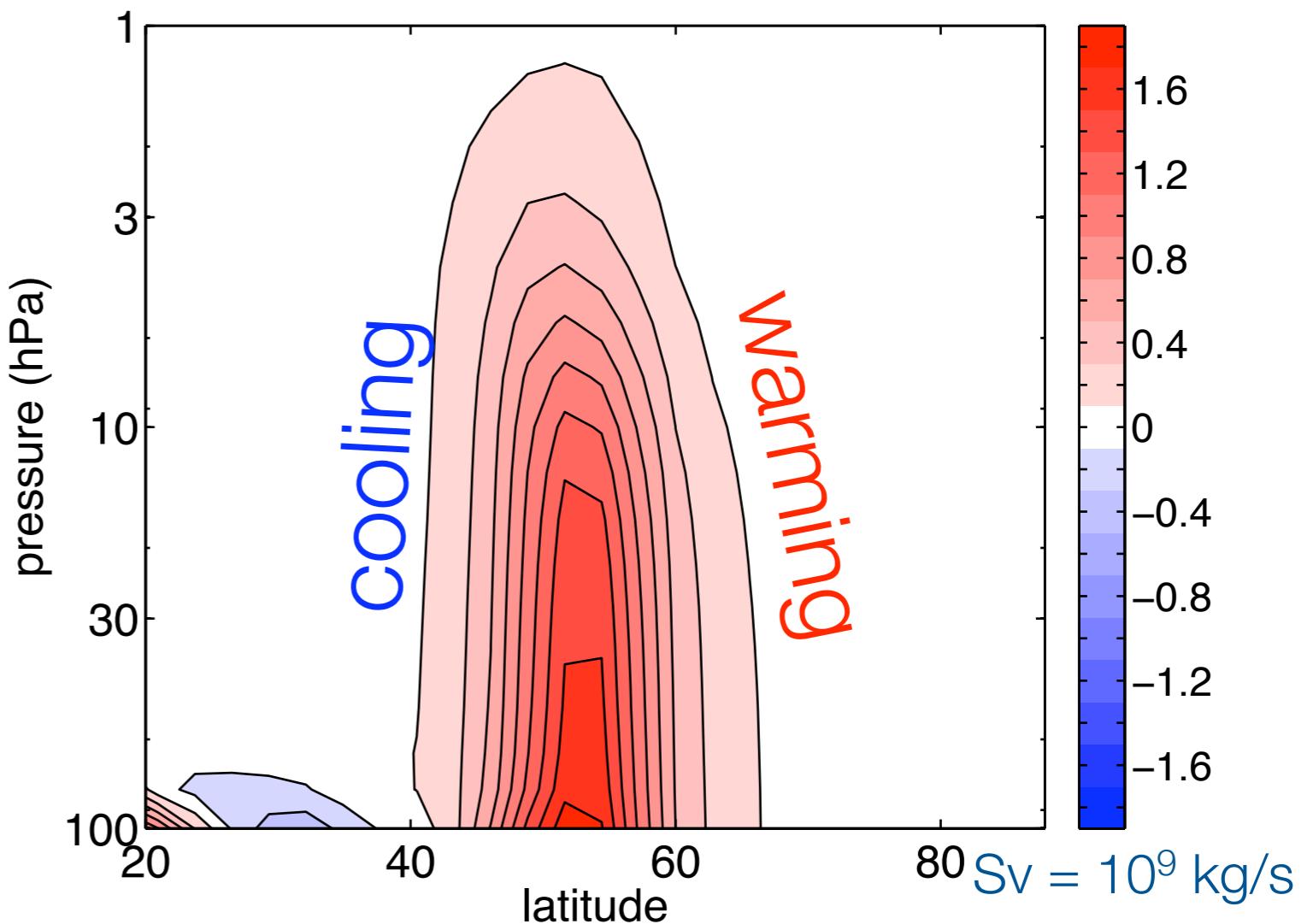
phase shift = 90°



Possible Mechanisms

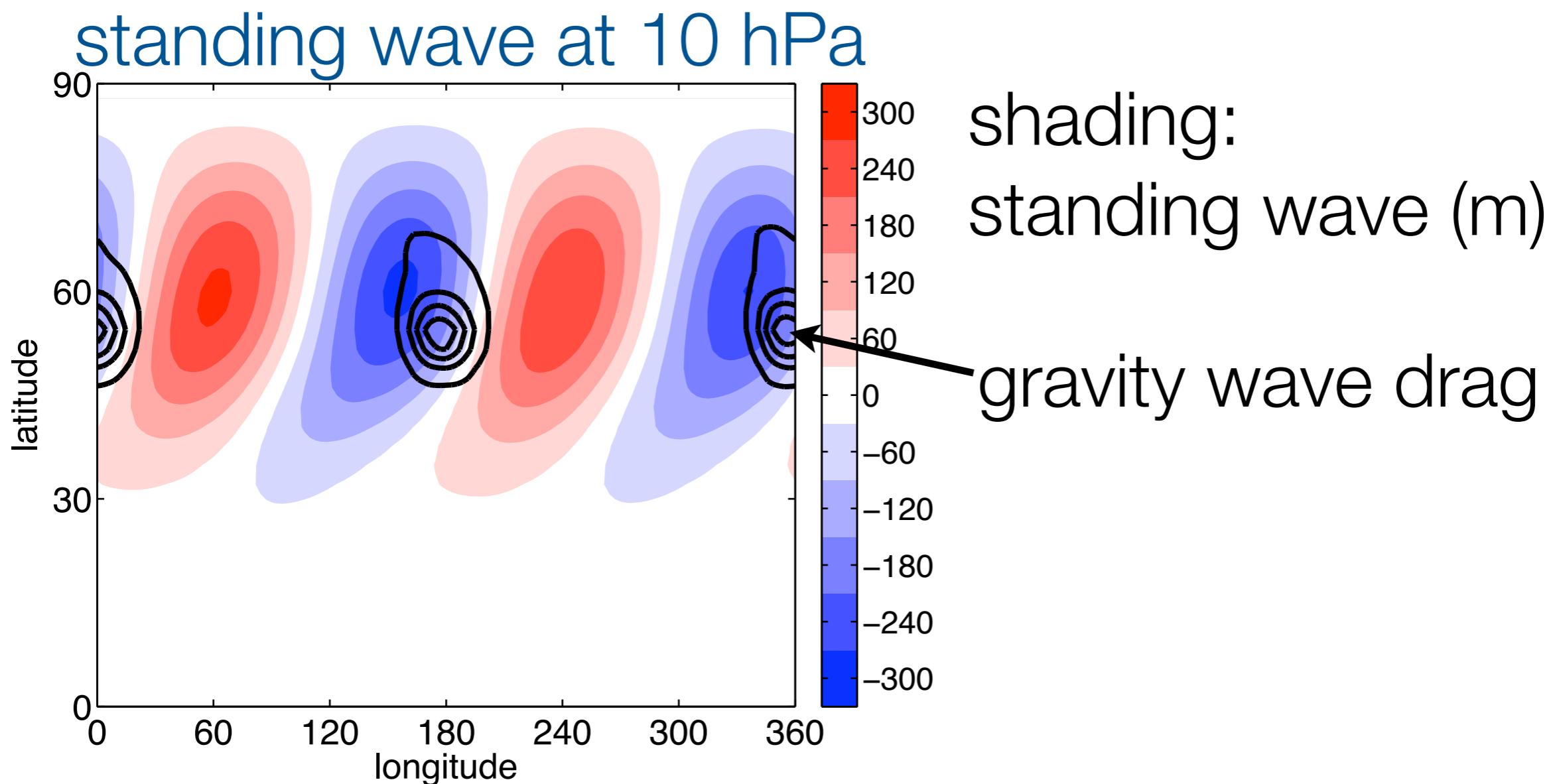
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Possible Mechanisms

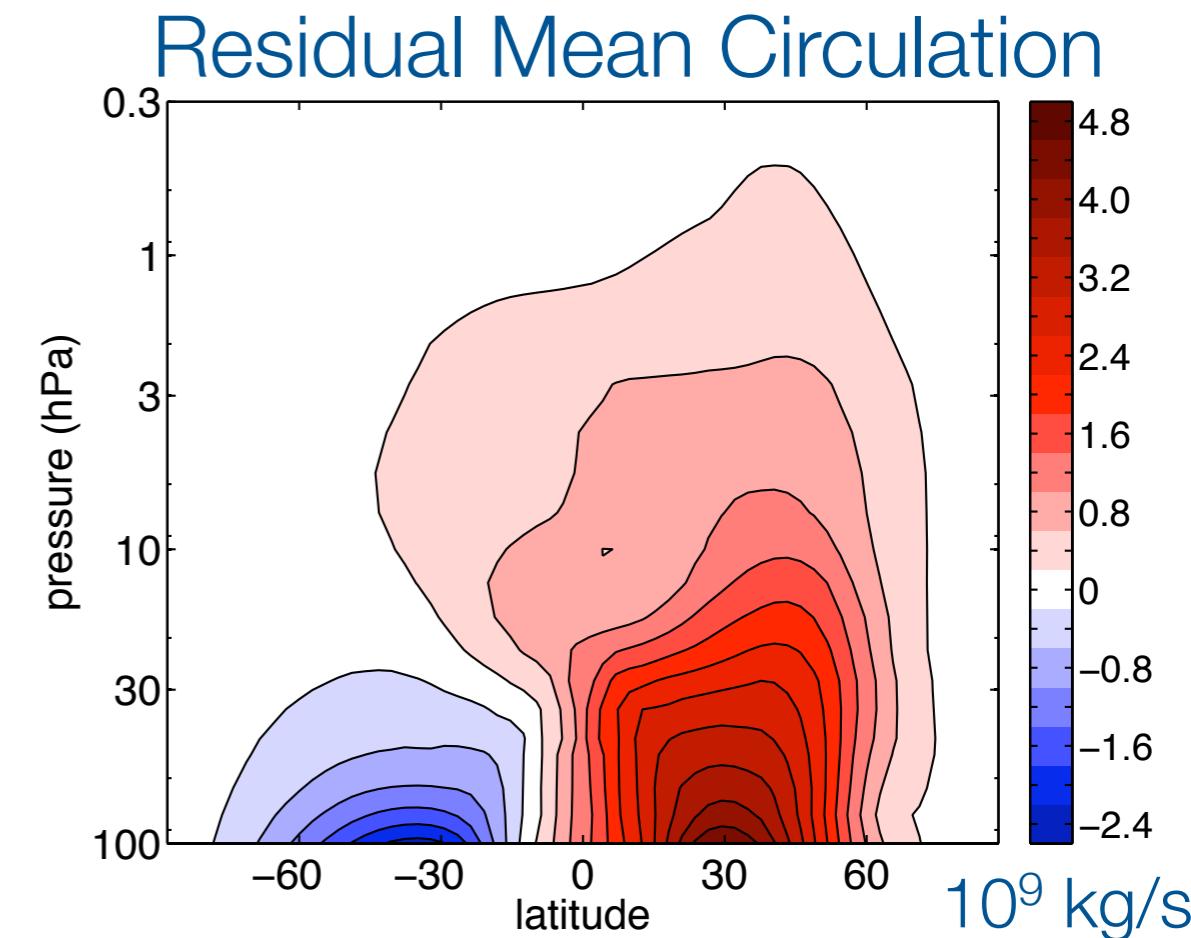
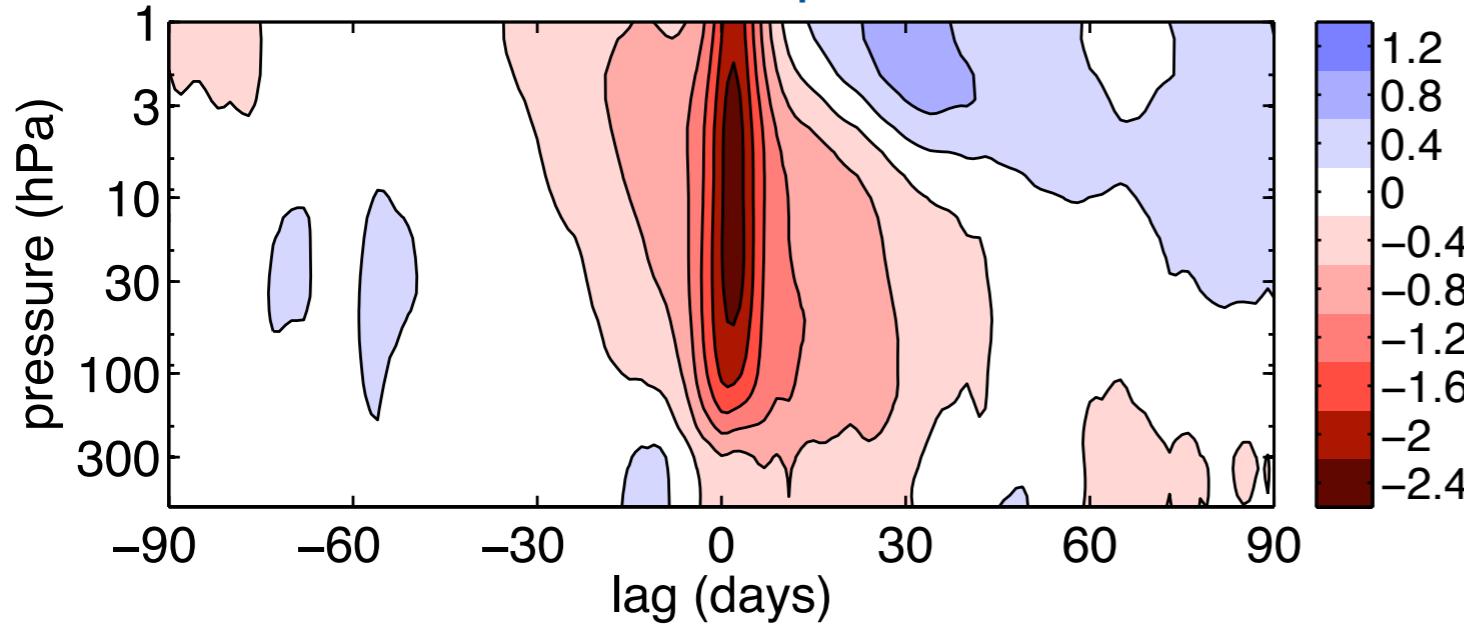
- GWD perturbation would modify temperature and wind structure, redirecting resolved waves
- GWD has zonal structure, can directly interact with resolved waves



Conclusions

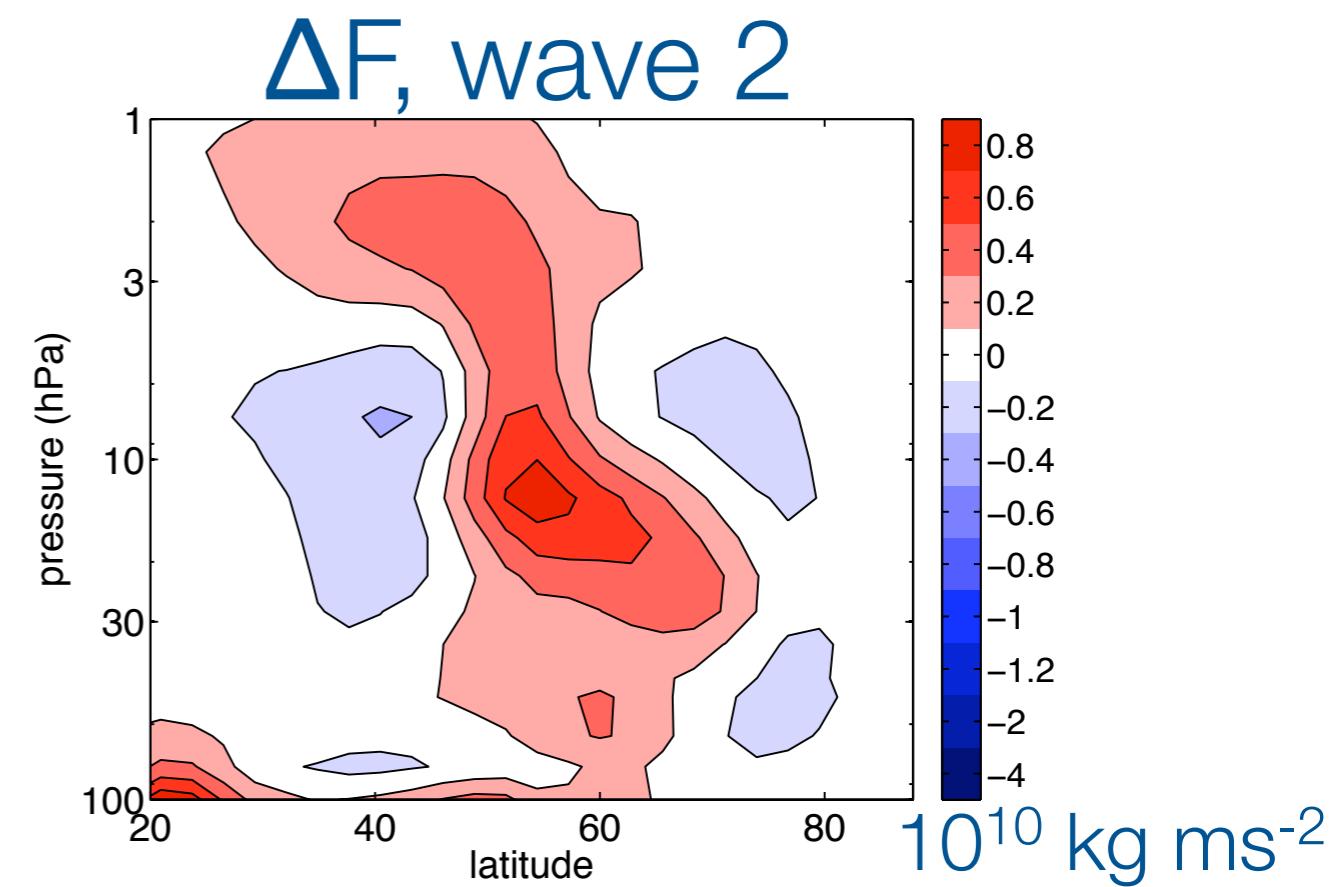
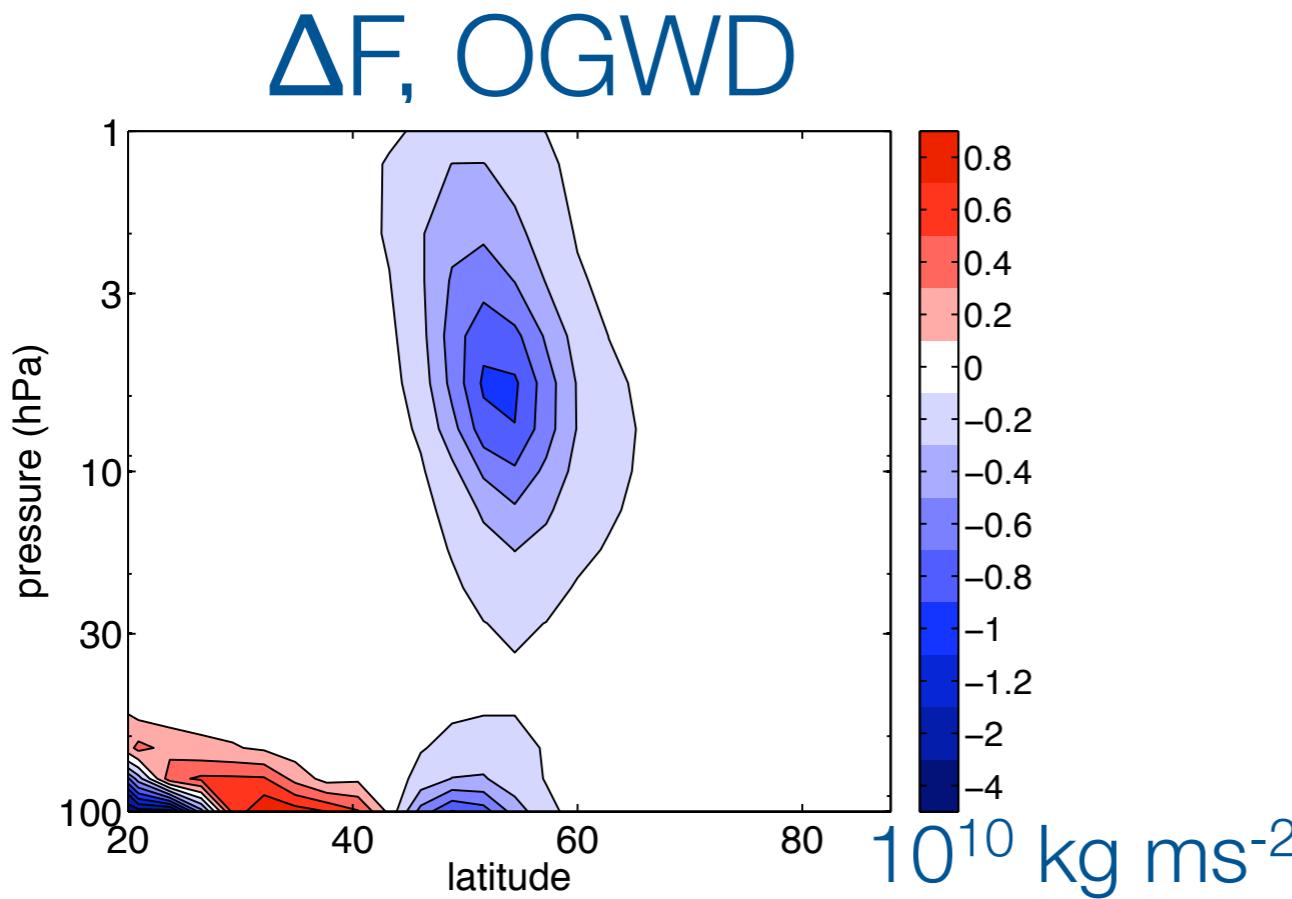
- We've developed a (relatively) idealized framework for studying the interaction between gravity wave parameterizations and the resolved circulation

Annular Mode Composite, 42 SSWs



Conclusions

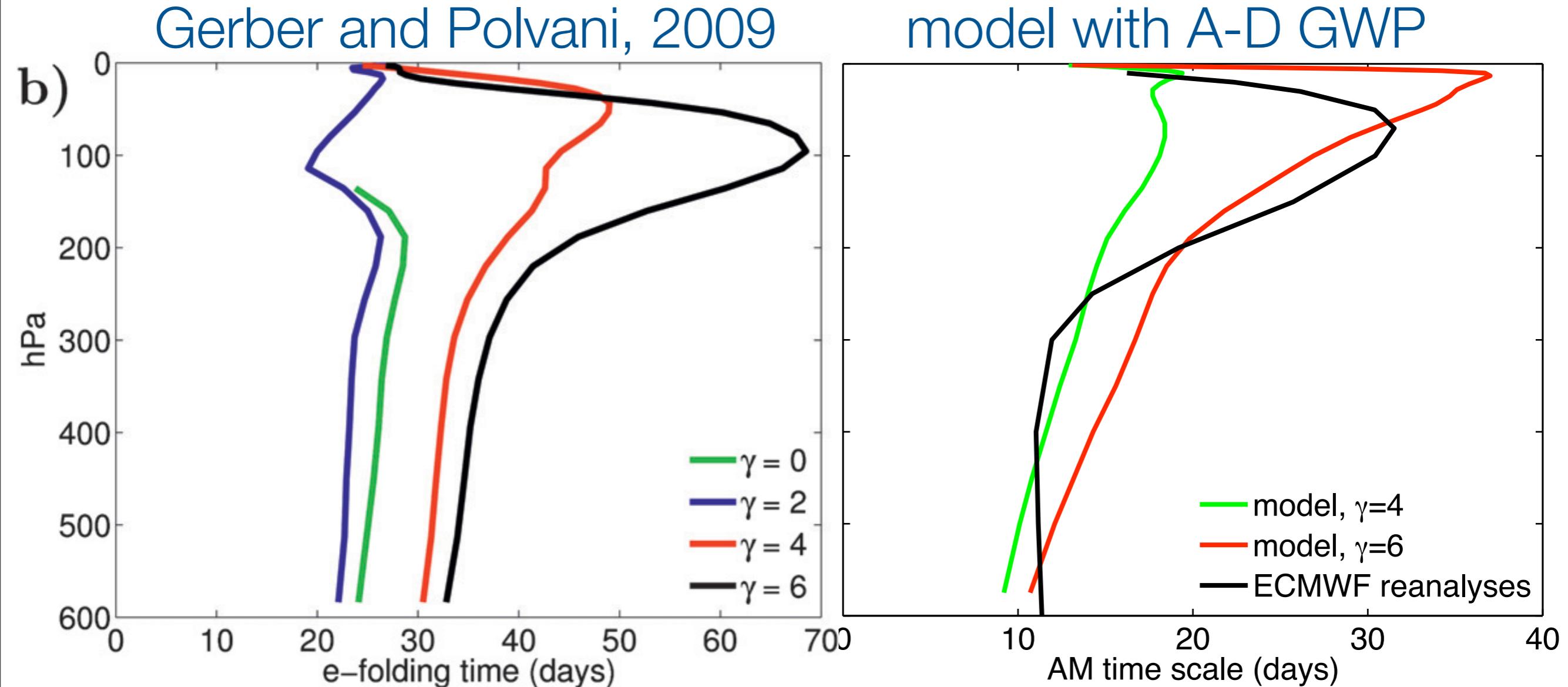
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Conclusions

- We've developed a (relatively) idealized framework for studying the interaction between gravity wave parameterizations and the resolved circulation
- Perturbation experiments show that the response to changes in gravity wave drag can be hard to predict: compensation between resolved and parameterized wave forcing
- Suggests large scale constraints on the structure of the residual mean circulation in the stratosphere

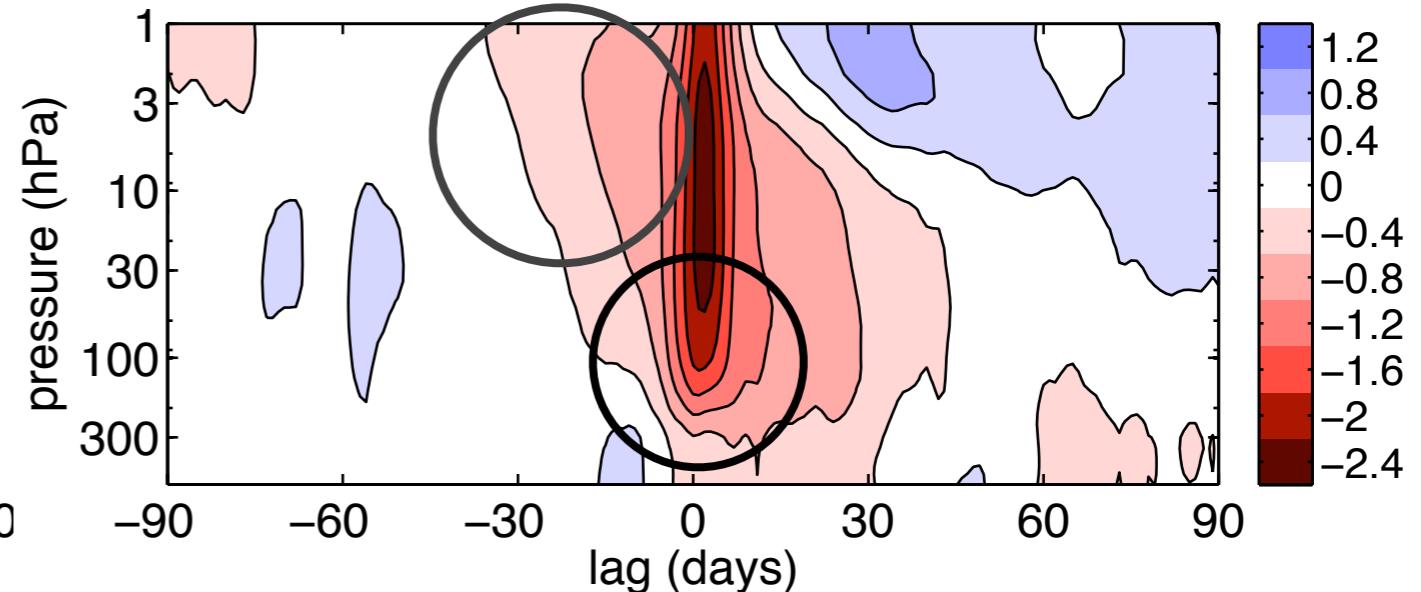
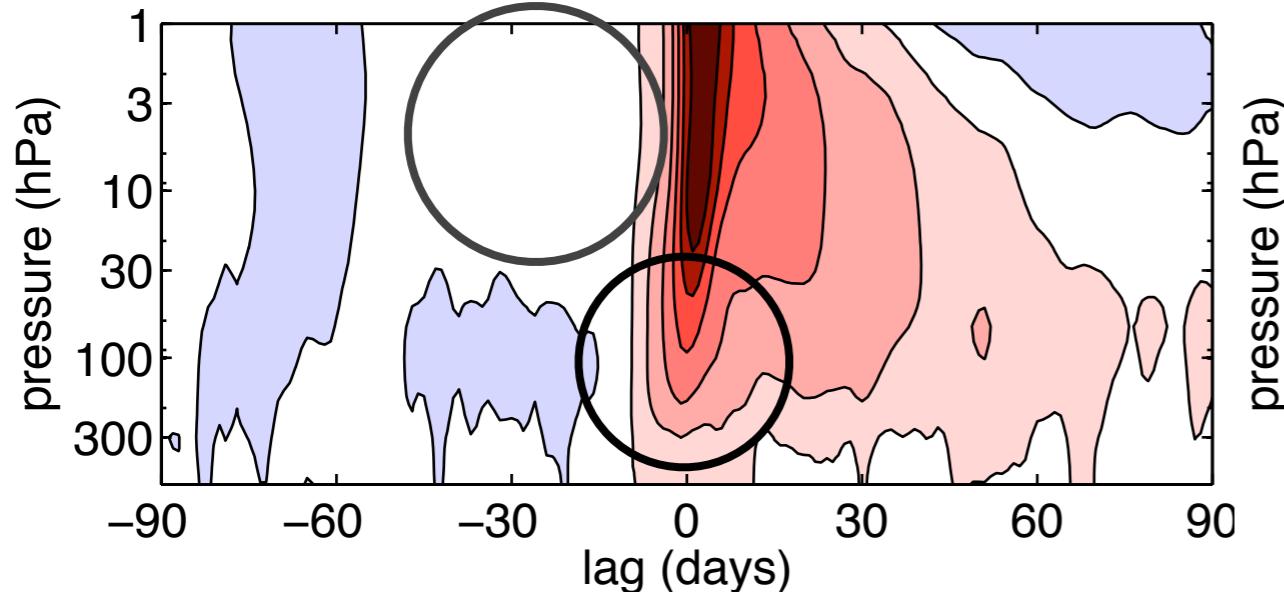
Annular mode e-folding time scale: stratospheric variability impacts troposphere



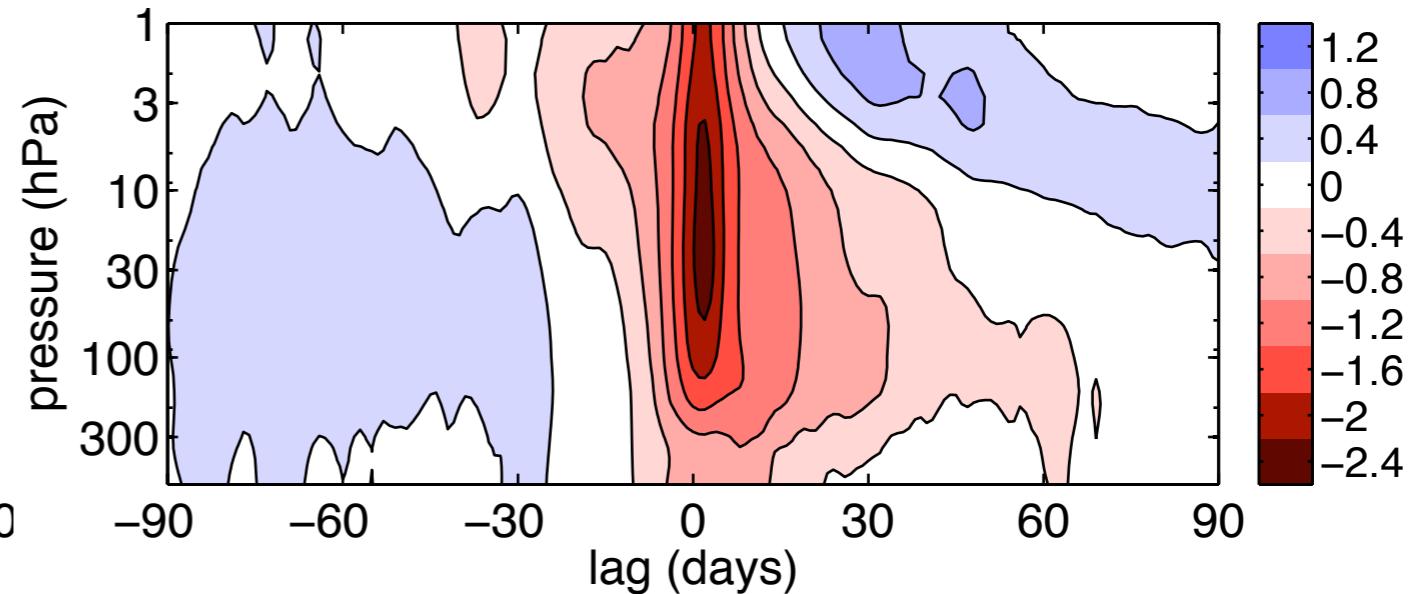
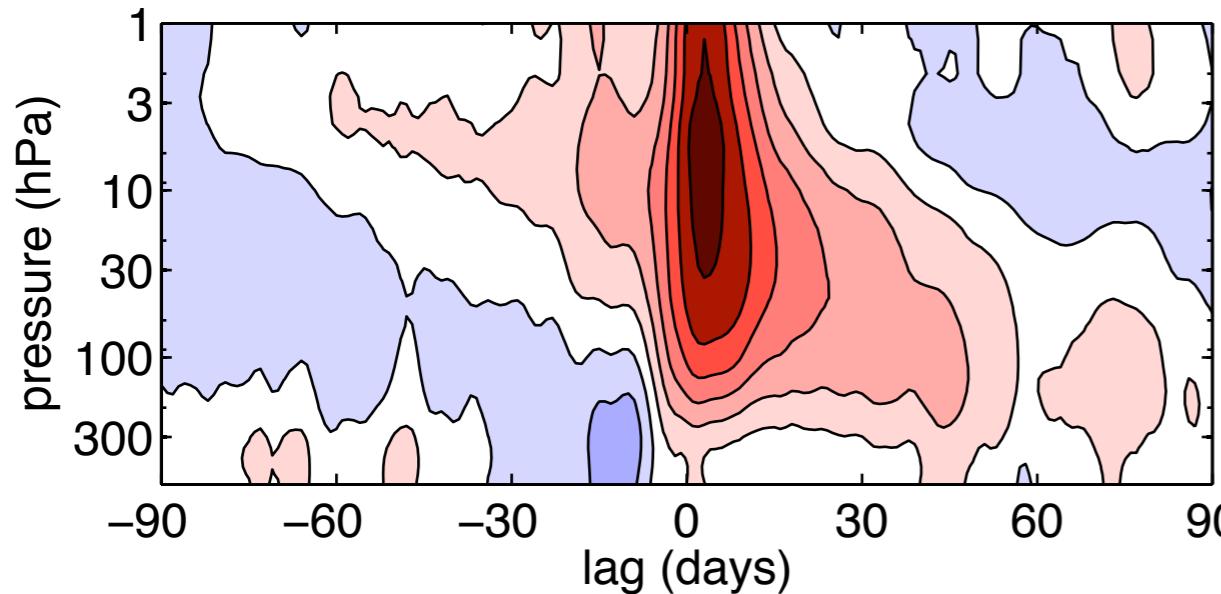
previously reported results are robust

Impact of A-D scheme on the annular mode and SSWs

Rayl. Friction, $\gamma=4$ (46 events) A-D GWP, $\gamma=6$ (42 events)



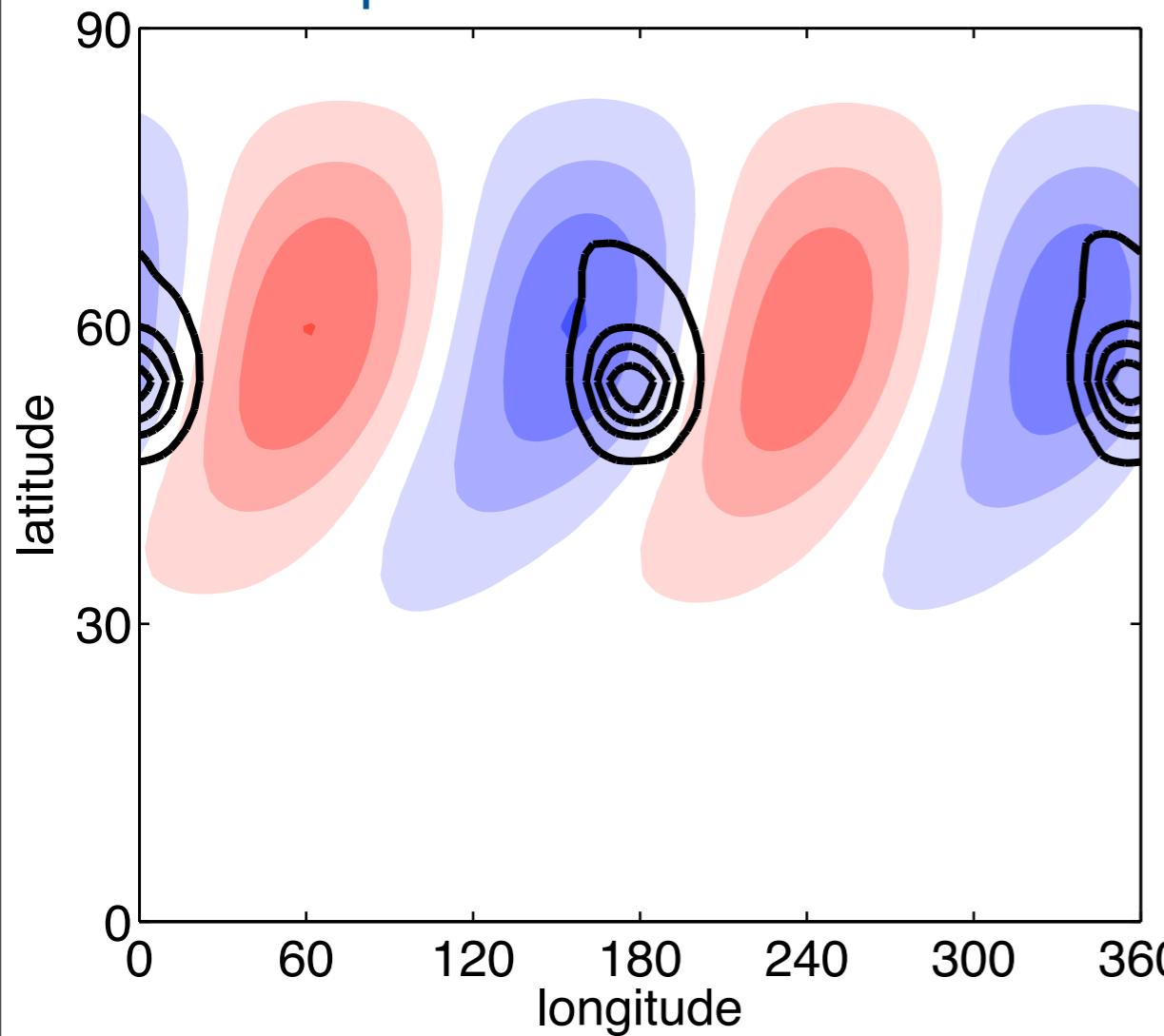
A-D spectral + Orographic GWP, $\gamma=6$
phase shift = 0° (29 events) phase shift = 90° (41 events)



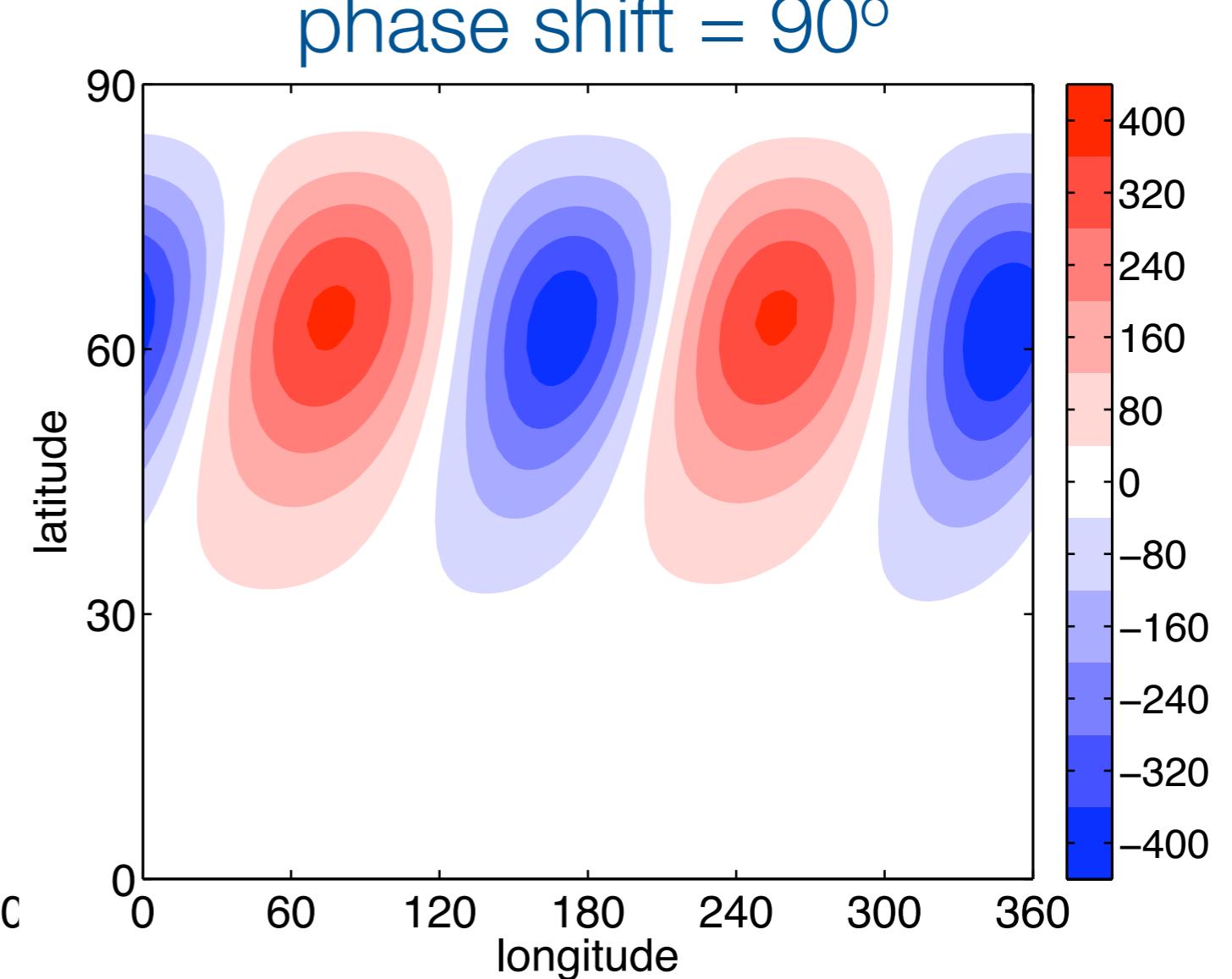
Interaction between OGWD and stationary waves

z' (shading) and OGWD at 10 hPa

phase shift = 0°



phase shift = 90°



Residual Mean Streamfunction

